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Supporting Information for

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**Larger aftershocks happen farther away: non-separability of magnitude and spatial distribution of aftershocks**

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10 **Contents of this File**

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15 **Introduction**

16 This document contains the following supporting information: (1) Maps and  
17 descriptions of the subset of aftershock sequences for which the largest aftershock  
18 is larger than the mainshock (Table S1, Figure S1); (2) Plots of aftershock ranked  
19 distance as a function of aftershock-mainshock magnitude difference, for alternative  
20 Catalogs and reference frames (Figures S2 – S4); and (3) Descriptions of the

21 procedures used to fit parameters (Text S1) and establish statistical significance  
22 (Text S2).

23

#### 24 **Text S1. Fitting the power-law kernel to the aftershock distributions**

25 We fit the parameters of the spatial kernel (Eqs. 4-6) using a grid search for  
26 maximum likelihood (Fig. S5). We use the population of aftershocks with  $\Delta M \leq 0$ ,  $r \leq$   
27 10 km, and  $t \leq 3.2$  days.

28

29 The maximum likelihood estimate (MLE) should weight the aftershocks of each  
30 magnitude bin approximately evenly, since the combined mainshocks of each  
31 magnitude bin tend to produce the same total number of aftershocks [*Helmstetter et*  
32 *al.*, 2005]. However, due to secondary triggering and other factors, the population of  
33 aftershocks may not be entirely independent and identically distributed. This  
34 makes it difficult to rigorously compare the likelihoods of the best-fit and the  
35 theoretically constrained model. Nevertheless, we use the Akaike Information  
36 Criterion (AIC) [*Akaike*, 1973] to adjust for the difference in free parameters and  
37 calculate rough confidence bounds on the best-fit parameters. The difference in AIC  
38 between the two models is defined as

$$39 \quad \Delta AIC = 2k_{th} - 2\log(L_{th}) - 2k_{mle} + 2\log(L_{mle}). \quad (S1)$$

40 Here  $k_{th}$  and  $k_{mle}$  are the number of free parameters, and  $L_{th}$  and  $L_{mle}$  are the  
41 maximum likelihoods computed for the theoretical and MLE parameter sets,  
42 respectively. The ‘plausibility’ of a candidate model relative to the best-fit model  
43 under the AIC framework can be expressed as  $P = \exp(-\Delta AIC/2)$  [*Burnham and*  
44 *Anderson*, 2002] (equivalent to a likelihood ratio corrected for number of  
45 parameters). Contours of  $\Delta AIC$  (Eq. S1) are given in Fig. S5, along with the  $P = 5\%$   
46 ( $\Delta AIC \approx 6$ ) boundary. The difference in AIC score between the MLE best-fit and the  
47 theoretically constrained parameter set is 7.4, putting the theoretical parameters  
48 near the boundary of what is considered to be plausibly supported by the data.

49

50 The AIC also justifies the use of the geometrical correction given by Equation 5. The  
 51 maximum likelihood score for a model with no geometrical correction, i.e.  $N(r) = 1$ ,  
 52 gives a power-law decay constant  $\hat{\gamma} = 3.21$ , and a log-likelihood 3714 units smaller  
 53 than with Eq. 5. Using a simpler geometrical correction of the form  $N(r) = r$  gives  $\hat{\gamma} =$   
 54 3.67 and a likelihood score 179 log-likelihood units smaller. The reduced number of  
 55 free parameters in these alternative models can only account for a difference of 2  
 56 log-likelihood units, meaning the AIC dramatically prefers the geometrical  
 57 correction given by Eq. 5.

58

59 For comparison with previous studies, we also directly fit the linear density using  
 60 least squares, following Felzer and Brodsky [2006]. This approach yields a  
 61 geometrical increase of  $r^{1.05 \pm 0.02}$  for distances less than  $\frac{1}{2}$  the mainshock rupture  
 62 length (from Eq. 6), and a power-law decay of  $r^{-1.77 \pm 0.007}$  for distances greater than  
 63 the mainshock rupture length. The uncertainty estimates given are 95% confidence  
 64 bounds.

### 65 **S1.1. Fitting the internal length scale for individual magnitude bins**

66 In the inset of Figure 4, we plot the internal length-scale  $d$  for individual magnitude  
 67 bins. To get these individual  $d$  values, we solve for the expected median  $r_{50}$  of the  
 68 power law kernel (Eq. 4), and use this formula to compute  $d$  from the observed  
 69 median of the aftershock population. The median of Eq. 4 is given by the formula

$$70 \quad r_{50} = d \left[ \left( 1 - 0.5D \right)^{1/(1-\gamma/2)} - 1 \right]^{1/2}, \quad (S2)$$

71 where

$$72 \quad D = 1 - \left[ \left( \frac{R}{d} \right)^2 + 1 \right]^{1-\gamma/2}$$

73

74 The constant  $R$  is the maximum bound on the aftershock distance. For fitting the  
 75 median we use a maximum bound  $R = 0.75w$ , where  $w = 7.8$  km, and use the  
 76 approximation  $N(r) \sim r$ , valid for  $r \ll w$  (Eq. 5).

77

78 **Text S2. Comparison with the ETAS model**

79 This study rests on the interpretation of combined aftershock sequences from a  
80 catalog in which the actual triggering relationships between foreshocks,  
81 mainshocks, and aftershocks are ambiguous. Whenever such a stacking method is  
82 employed, there is a possibility that unanticipated signals will show up due to  
83 unconsidered interactions between overlapping sequences. We therefore compare  
84 the results from the earthquake catalogs to results from simulated earthquake  
85 catalogs constructed with Epidemic-Type Aftershock sequence (ETAS) model.

86 **S2.1. ETAS ingredients**

87 The ETAS model combines well-established statistical laws of aftershocks to  
88 produce realistic simulated earthquake catalogs. We start with an initial catalog of  
89 spontaneous or “background” earthquakes, with random magnitudes assigned from  
90 the Gutenberg-Richter distribution. To each of these earthquakes, we assign a  
91 number of aftershocks (offspring) that depends exponentially on the parent  
92 magnitude, and distribute these offspring in time according to Omori’s law.  
93 Magnitudes are then assigned to the offspring, and the process repeats until a  
94 generation arises with no new earthquakes. The ETAS simulation procedure is well  
95 described elsewhere [*Helmstetter and Sornette, 2003; Ogata, 1998; 2011; Zhuang et*  
96 *al., 2002*]. We generate an initial catalog of background earthquakes proportional to  
97 a gridded estimate of background rate for the California catalog, and set the ETAS  
98 parameters to produce catalogs with on average the same total number of events as  
99 in the observed catalog. For reproducibility, we report the parameters in Table S2.  
100 See the aforementioned references for a more elaborate description of the function  
101 of each parameter in the simulation.

102  
103 We assign parent-offspring distances using a kernel of the form

104 
$$p(r) \propto r(r^2 + d_{MS}^2)^{-\gamma/2} \quad (S3)$$

105 where  $d_{MS}$  is the length scale of the parent earthquake from Eq. (6), and  $\gamma = 3.0$ .

106

107 Depths for the initial seed catalog of ‘background’ earthquakes are drawn from the  
108 observed depth distribution of the catalog mainshocks, but aftershock depths are  
109 allowed to range from 0 to 30 km. In the real catalog, the aftershock depths are  
110 more restricted, leading to the geometrical correction  $N(r)$  given as Eq. 5 in the main  
111 text. In the simulations we use the approximation  $N(r) \sim r$  (Eq. S3) to permit the  
112 efficient numerical generation of random aftershock distances. The ETAS  
113 simulations generated with Eq. S3 produce median aftershock distances that are  
114 slightly larger than those in the catalog ( $\sim 2.1$  km for the simulations vs.  $\sim 1.4$  km for  
115 the real catalog). Since the goal of the simulations is to look for differences in the  
116 distributions of large and small aftershocks, and not to perfectly reproduce the  
117 observed catalog, these small discrepancies should not be a problem. To be safe we  
118 compare the results as a function of ranked distance, which normalizes each  
119 aftershock zone by the distribution of smaller events and corrects for any small  
120 systematic differences between the observation and simulations.

## 121 **S2.2. ETAS results**

122 The results are summarized in Fig. S7. While the scatter in the distribution of ranked  
123 distance increases with  $\Delta M$  (as the sample size diminishes), the median ranked  
124 distance stays centered on 0.5, and the ETAS simulations do not reproduce the  
125 observed tendency for the larger aftershocks to occur farther away (Fig S7). The  
126 difference between observations and simulations is significant above the 95% level  
127 for  $\Delta M \geq -0.5$ , as measured by the fraction of simulations that produce a mean or  
128 median ranked distance as large as that observed. While the mean ranked distance  
129 in the observational data continues to grow with  $\Delta M$ , the sample size becomes too  
130 small for a meaningful comparison above  $\Delta M = 1$ .

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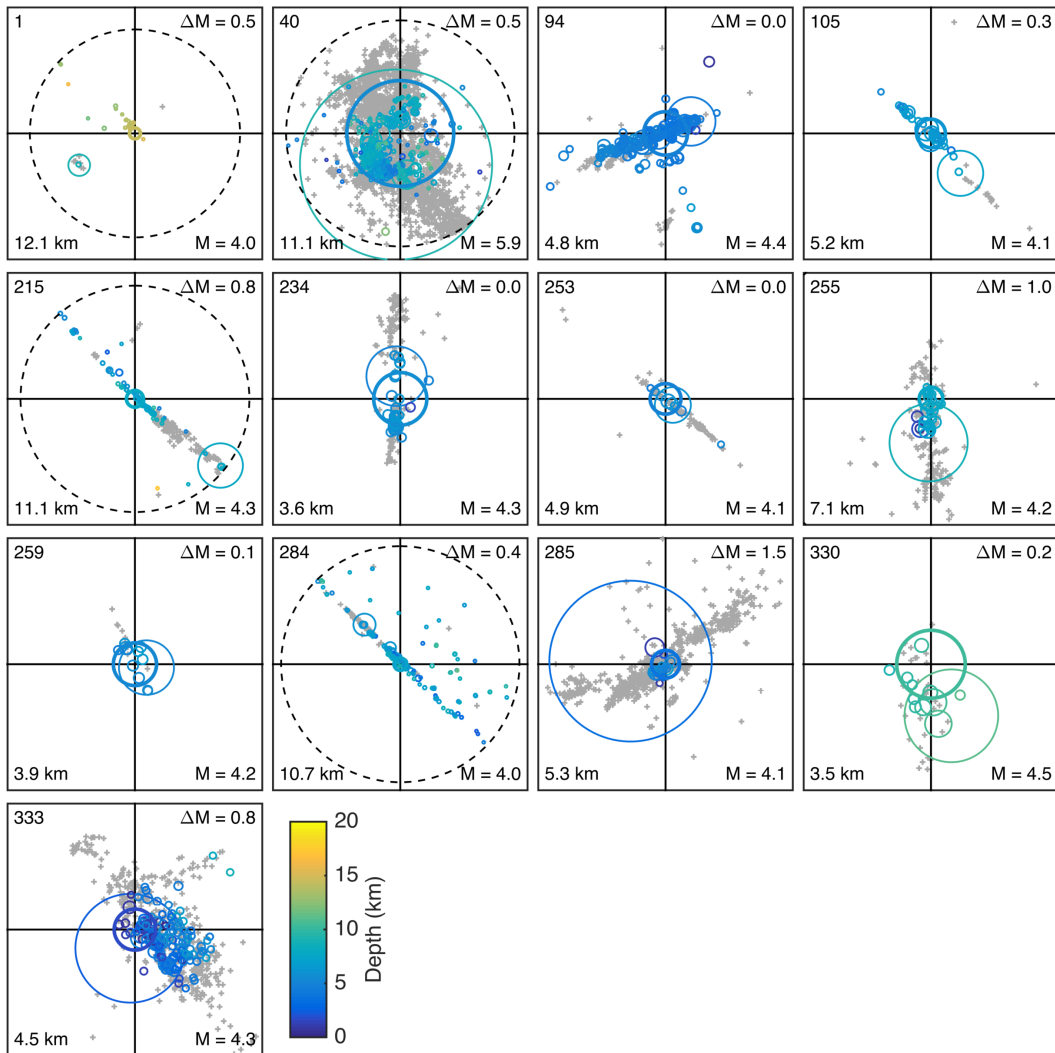
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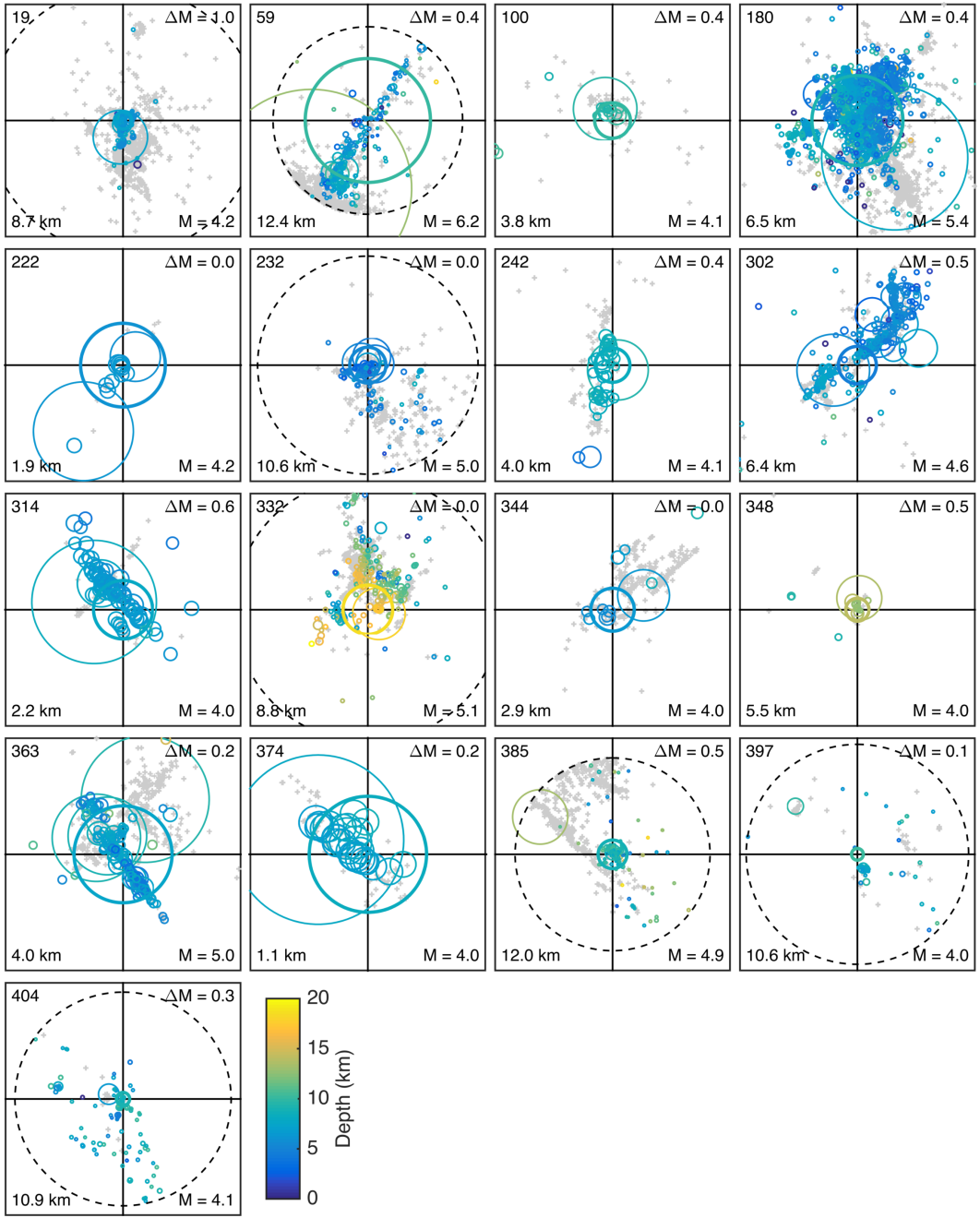
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140 **Fig S1a.** Northern California mainshock-aftershock sequences for which the largest aftershock is as  
 141 large or larger than the mainshock in map view. Mainshock is the thick circle centered on the  
 142 crosshairs. Other colored circles are aftershocks up to and including the first aftershock larger than  
 143 the mainshock. Circles have radius equal to the rupture radius (Eq. 6). Grey crosses are aftershocks in  
 144 the 10 days after the first aftershock larger than the mainshock. Numbers in corners are: NW -  
 145 sequence number (Fig. 1); NE - magnitude difference between largest aftershock and mainshock; SE -  
 146 mainshock magnitude; SW - edge length of one quadrant of each plot for scale. The dashed circle  
 147 shows the 10 km limit for collecting aftershocks.

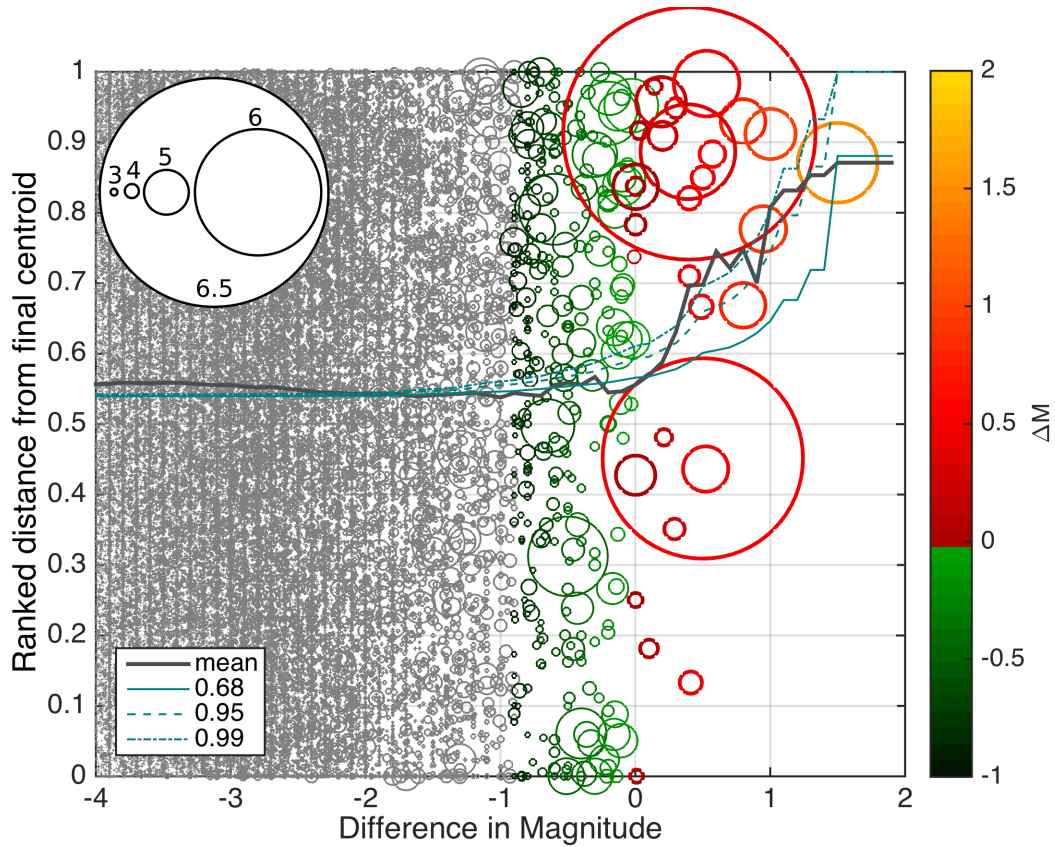
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**Fig S1b.** Sequence plots for Southern California. See Fig. S1a caption for additional details.



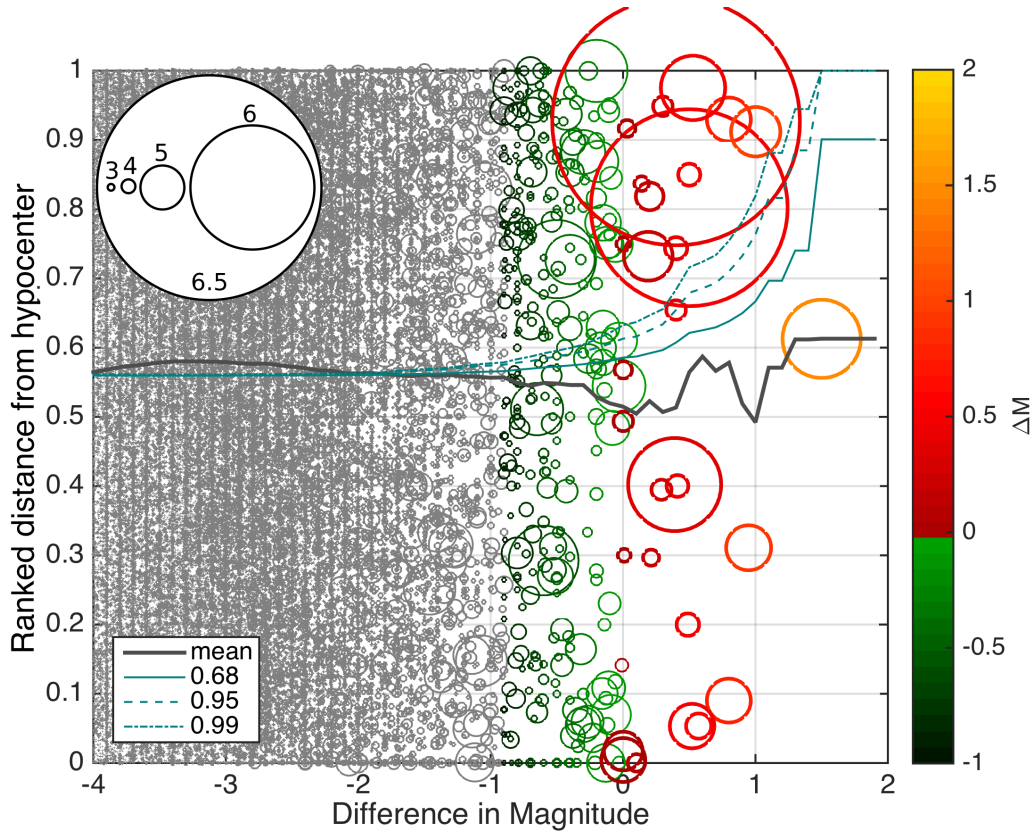
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152 **Fig S2.** Ranked distance of aftershocks with respect to the *final centroid* location of all aftershocks, up  
 153 to and including the first aftershock with magnitude as large or larger than the mainshock. Compare  
 154 to Fig. 3 in the main text. Circles are scaled to aftershock magnitude. Solid black line is the running  
 155 average over one magnitude unit. Green lines are 68, 95, and 99% confidence bounds on the mean  
 156 from repeated random sampling of the smaller aftershocks ( $-2 \leq \Delta M < -1$ ).

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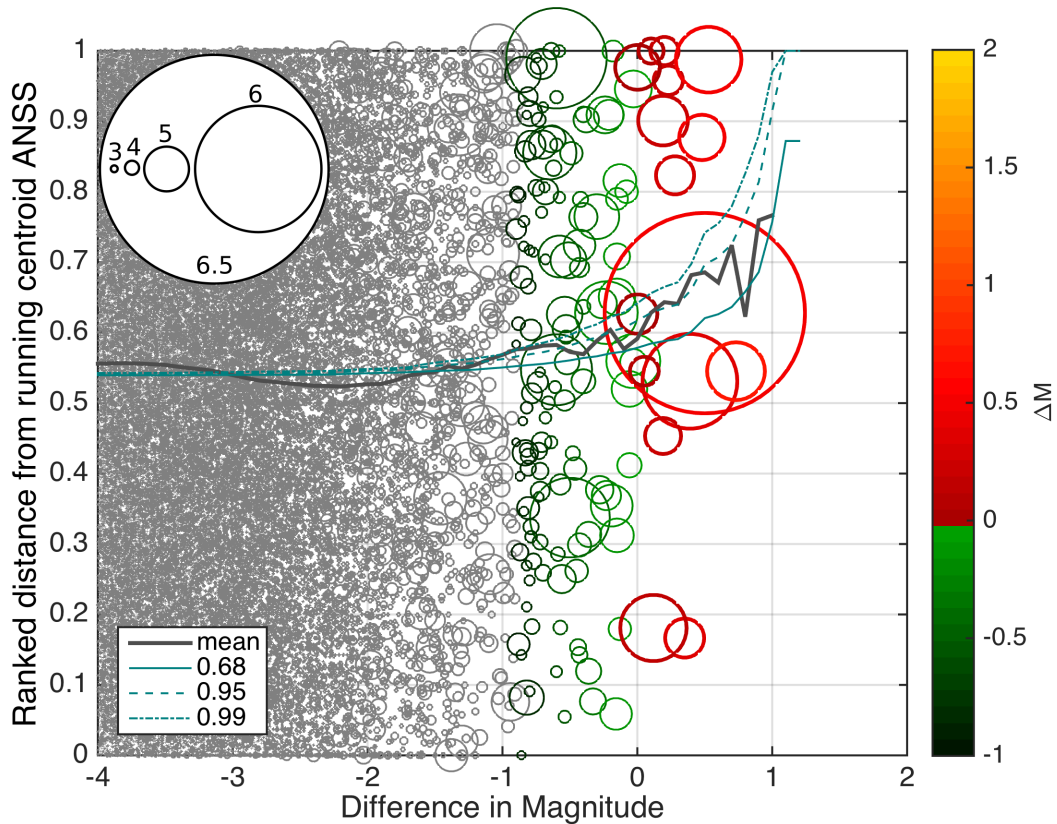


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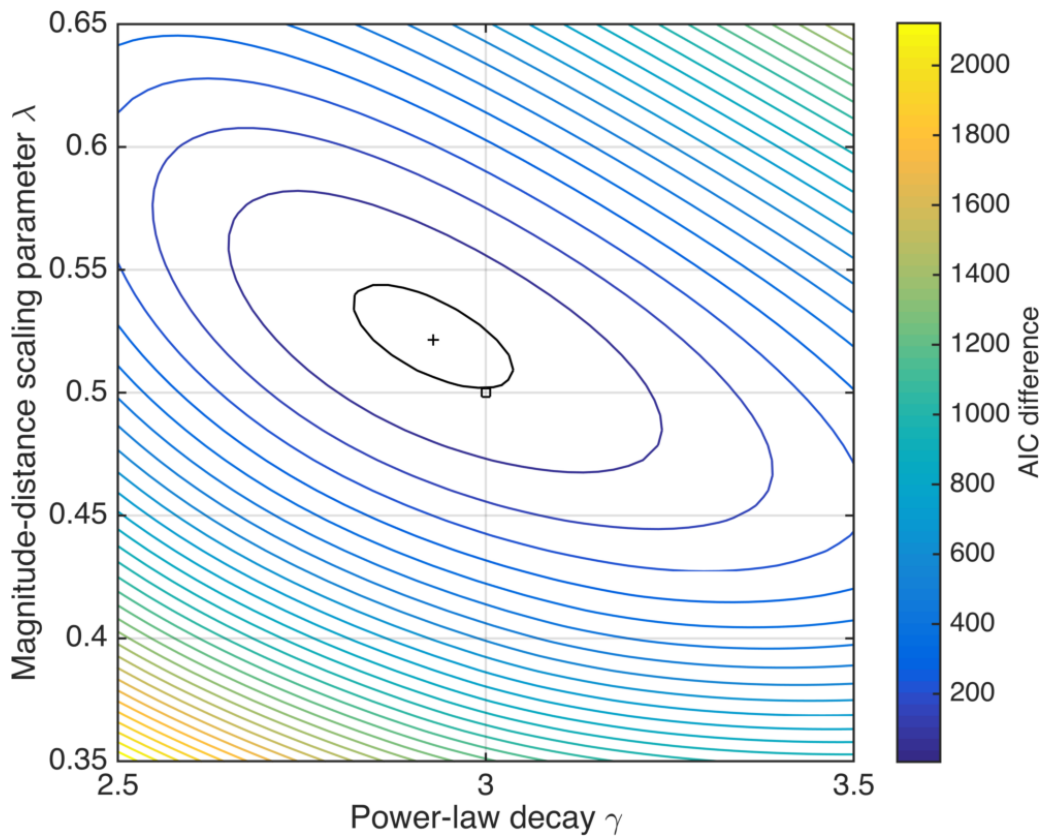
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**Fig S3.** Ranked distance to aftershocks with respect to the mainshock *hypocenter*, as a function of aftershock-mainshock magnitude difference. Compare to Fig. 3 in the main text. Circles are scaled to aftershock magnitude. Solid black line is the running average over one magnitude unit. Green lines are 68, 95, and 99% confidence bounds on the mean from repeated random sampling of the smaller aftershocks ( $-2 \leq \Delta M < -1$ ).



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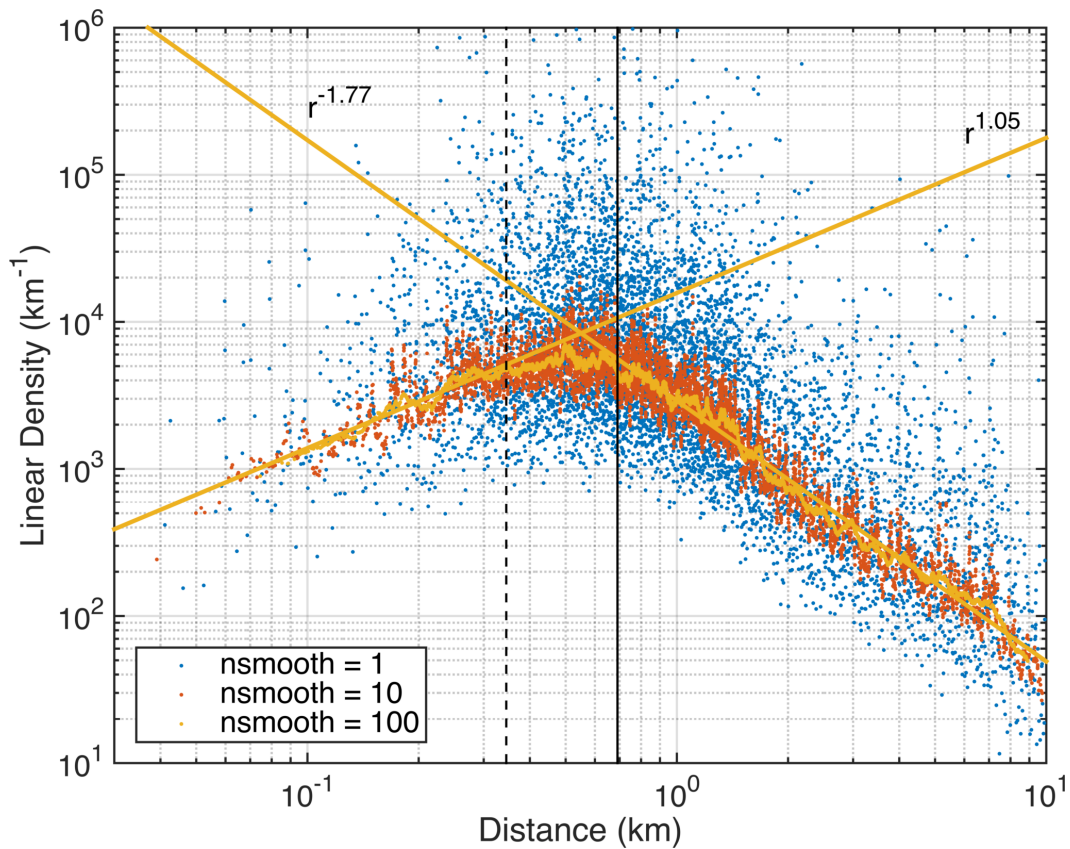
**Fig S4.** Ranked aftershock distance in the ANSS catalog, with respect to the running centroid location of all aftershocks, up to and including the first aftershock with magnitude as large or larger than the mainshock. Compare to Fig. 3 in the main text. Mainshock magnitudes in the ANSS catalog are limited to M4.5 and above ( $\sim 1$  km source radius), to account for the greater location uncertainty. Circles are scaled to aftershock magnitude. Solid black line is the running average over one magnitude unit. Green lines are 68, 95, and 99% confidence bounds on the mean from repeated random sampling of the smaller aftershocks ( $-2 \leq \Delta M < -1$ ).



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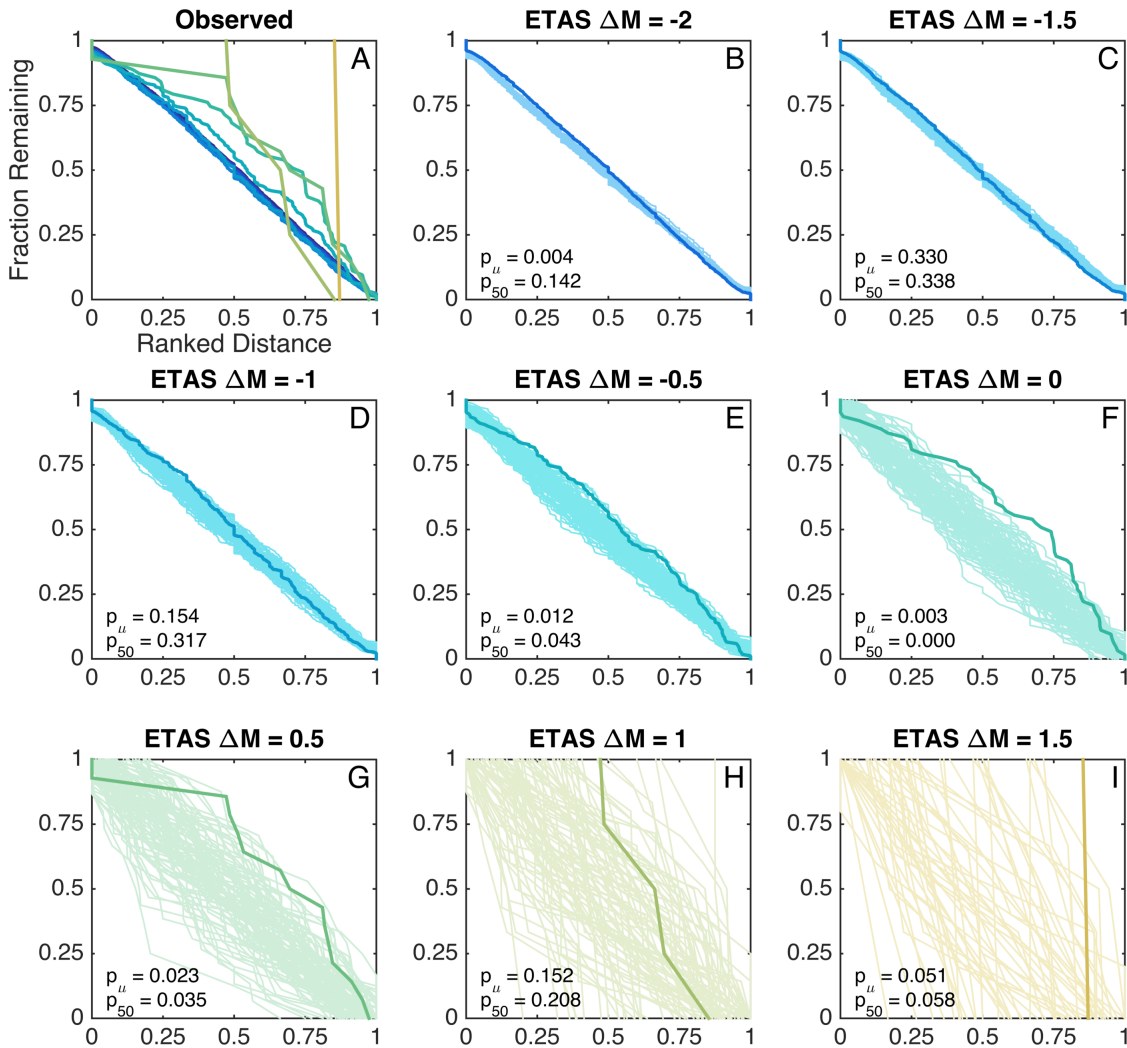
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179 **Fig S5.** Maximum likelihood estimate (MLE) of the spatial kernel parameters (+). Contours give the  
 180 difference in Akaike Information Criterion (AIC), relative to the MLE (Eq. S1) The black contour  
 181 encloses the parameter space with greater than 5% relative likelihood compared to the MLE.  
 182 The theoretically-motivated parameters are marked by a square.



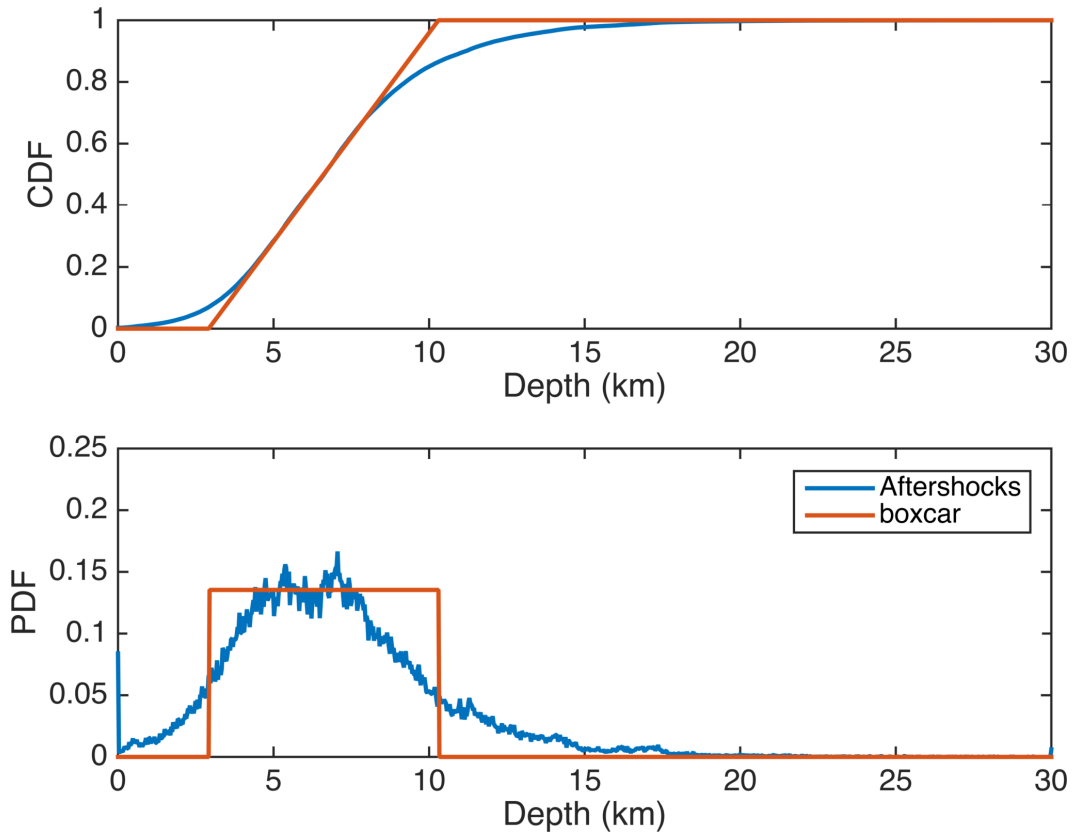
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184 **Fig S6.** Least-squares fit to linear density for M4 – 4.5 mainshocks, using aftershocks with  $\Delta M < 0$ , up  
 185 to 3.2 days after the mainshock. The decay is fit to points outside the estimated rupture length (solid  
 186 black line), and the geometrical increase is fit to points inside  $\frac{1}{2}$  the rupture length (dashed black  
 187 line). The legend gives the number of points over which the data are smoothed. The three smoothing  
 188 windows yield spatial decays that agree within 0.01. The 95% confidence bounds for the two  
 189 exponents are  $1.05 \pm 0.02$  and  $-1.77 \pm 0.007$ .



190

191 **Fig S7. (A)** Ranked distance as a function of the difference between aftershock and mainshock  
 192 magnitude  $\Delta M$  (Fig 5C, main text). **(B-I)** Distribution of ranked distance curves from 100 simulated  
 193 ETAS catalogs in which distance depends only on mainshock magnitude, for varying  $\Delta M$ . Thick lines  
 194 are observed distributions from (A). Reported  $p$ -values are the fraction of 1000 ETAS simulations  
 195 that produce mean ( $p_{\mu}$ ) or median ( $p_{50}$ ) ranked distances larger than observed.



196

197

198 **Fig. S8.** Distribution of aftershock depths compared to the boxcar used to derive the geometrical  
 199 correction  $N(r)$  (Eq. 5). The width of the boxcar (7.39 km) was independently constrained by a  
 200 maximum likelihood fit to the distance decay of aftershocks through Eq. 5, and is centered on the  
 201 median of the data for comparison.

201

202 **Table S1.** Sequences with aftershocks larger than the mainshock

<b>Date</b>	<b>Time</b>	<b>Mag</b>	<b>Lat.</b>	<b>Lon.</b>	<b>Depth</b>	<b>Max</b>	<b>Seq.</b>	<b>Cat.</b>
	<b>(UTC)</b>				<b>(km)</b>	<b><math>\Delta M</math></b>	<b>#</b>	
<b>27-Sep-1982</b>	18:21	4.2	35.75	-117.75	7.2	1.0	19	SC
<b>12-Jan-1984</b>	3:11	4.0	37.41	-118.53	14.5	0.5	1	NC
<b>18-Jul-1986</b>	14:29	5.9	37.56	-118.44	5.5	0.5	40	NC
<b>22-Nov-1987</b>	1:53	6.2	33.08	-115.78	10.1	0.4	59	SC
<b>5-Apr-1990</b>	2:39	4.4	37.86	-121.99	4.8	0.0	94	NC
<b>29-Aug-1990</b>	1:06	4.1	33.24	-116.05	9.9	0.4	100	SC
<b>6-Sep-1990</b>	12:48	4.1	36.68	-121.31	6.1	0.3	105	NC
<b>15-Aug-1995</b>	22:39	5.4	35.77	-117.66	9.1	0.4	180	SC
<b>26-May-1998</b>	20:31	4.3	36.81	-121.54	7.8	0.8	215	NC
<b>17-Feb-1999</b>	3:08	4.2	32.60	-116.16	6.0	0.0	222	SC
<b>19-Oct-1999</b>	1:53	5.0	34.86	-116.40	5.3	0.0	232	SC
<b>8-Jan-2000</b>	21:41	4.3	38.76	-122.92	5.6	0.0	234	NC
<b>12-Jun-2000</b>	19:00	4.1	32.89	-115.50	8.2	0.4	242	SC
<b>30-Jun-2001</b>	17:34	4.1	36.70	-121.33	4.8	0.0	253	NC
<b>12-Jul-2001</b>	17:31	4.2	36.03	-117.87	7.5	1.0	255	NC
<b>5-Dec-2001</b>	14:29	4.2	39.05	-123.12	5.2	0.1	259	NC
<b>28-Aug-2004</b>	4:30	4.0	36.58	-121.18	7.4	0.4	284	NC
<b>16-Sep-2004</b>	7:07	4.1	38.01	-118.67	3.8	1.5	285	NC
<b>29-Aug-2005</b>	22:48	4.6	33.16	-115.62	4.6	0.5	302	SC
<b>12-Sep-2006</b>	0:11	4.1	32.71	-116.04	7.8	0.6	314	SC
<b>17-Jan-2008</b>	17:18	4.5	40.16	-122.76	10.3	0.2	330	NC
<b>7-Feb-2008</b>	7:12	5.1	32.41	-115.31	18.7	0.0	332	SC
<b>22-Apr-2008</b>	22:55	4.3	39.53	-119.93	1.6	0.8	333	NC

<b>27-Nov-2008</b>	21:14	4.0	35.97	-117.31	6.4	0.0	344	SC
<b>12-Dec-2008</b>	8:41	4.0	32.53	-115.52	14.1	0.5	348	SC
<b>29-Sep-2009</b>	10:01	5.0	36.39	-117.86	7.4	0.2	363	SC
<b>7-Mar-2010</b>	4:17	4.0	33.00	-116.34	8.2	0.2	374	SC
<b>11-Jun-2010</b>	3:08	4.9	33.39	-116.40	8.8	0.5	385	SC
<b>7-Dec-2010</b>	9:02	4.0	32.63	-115.77	10.2	0.1	397	SC
<b>7-Apr-2011</b>	8:58	4.1	32.64	-115.73	9.1	0.3	404	SC

203

204

205 **Table S2.** ETAS parameters

206

Parameter	Value	Description
$\mu$	13.2508 day <sup>-1</sup>	Background rate
$c$	0.0055 days	Omori time constant
$p$	1.1	Omori decay exponent
$k$	0.0085	Aftershock productivity
$\alpha$	0.9	Productivity-magnitude scaling
$M_c$	1.0	Minimum simulated magnitude
$b$	1.0	Gutenberg-Richter parameter

207

208 **References**

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