Monsoon drought over Java, Indonesia, during the past two centuries

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Monsoon droughts, which often coincide with El Niño warm events, can have profound impacts on the populations of Southeast Asia. Improved understanding and prediction of such events can be aided by high-resolution proxy climate records, but these are scarce for the tropics. Here we reconstruct the boreal autumn (October–November) Palmer Drought Severity Index (PDSI) for Java, Indonesia (1787–1988). This reconstruction is based on nine ring-width chronologies derived from living teak trees growing on the islands of Java and Sulawesi, and one coral δ18O series from Lombok. The PDSI reconstruction correlates significantly with El Niño-Southern Oscillation (ENSO)-related sea surface temperatures and other historical and instrumental records of tropical climate, reflecting the strong coupling between the climate of Indonesia and the large scale tropical Indo-Pacific climate system. 


1. Introduction

Droughts can have devastating effects on the peoples of largely agrarian areas of monsoon Asia, including Indonesia. Indonesian drought results from the failure of the east monsoon, and tends to be linked to ENSO warm events [Hastenrath, 1987; Harger, 1995]. During such times the Walker Circulation weakens as the Indonesian Low migrates eastward into the tropical Pacific, resulting in drought over much of the country [Allan, 2000]. Indonesia is, therefore, particularly relevant to the study of the Indo-Pacific climate system [Hackert and Hastenrath, 1986; Hastenrath, 1987; Nicholls, 1995]. The correspondence between Indonesian climate and ENSO may be the most intense of any land area worldwide, with correlations between ENSO indices and Indonesian rainfall ranging from 0.7–0.8 for some periods [Moore, 1995; van Oldenborgh, 2002].

Land-based rainfall proxies can considerably improve our understanding of past drought variability [e.g., Cook and Krusic, 2004], yet are scarce for Indonesia and the tropics as a whole. A network of ring-width chronologies has been derived from living teak (Tectona grandis L. F.) trees in Java and Sulawesi [D’Arrigo et al., 1994]. The development of these records follows pioneering efforts by Berlage [1931], who produced the first such record for Java. In Java and Sulawesi, annual growth rings can form in teak during the dry, east monsoon season which extends from ~June to November [Hastenrath, 1987]. Here we use the teak records, along with one coral oxygen isotope series, to develop a reconstruction of a drought index for Java, and describe its relation to past monsoon and ENSO variability.

2. Data and Methods

2.1. Climate Data

We use the updated global monthly Palmer Drought Severity Index (PDSI) data set (1870–2002) [Dai et al., 2004] (http://www.cdc.noaa.gov; v. 2). An index of meteorological drought, the PDSI is a proxy for soil moisture and streamflow [Dai et al., 2004]. The PDSI is based on temperature and precipitation data for global land areas, on a 2.5° grid. It was averaged over 5–10°S, 105–115°E, representing Java and vicinity (Figure 1). The Javanese PDSI series are compared to several tropical climate records, including El Niño sea surface temperature (SST) indices [Kaplan et al., 1998], the Southern Oscillation Index (SOI) [Ropelewski and Jones, 1987], and station records of precipitation for Jakarta, Indonesia (Global Historical Climate Network) [Peterson and Vose, 1997], and sea-level pressure (SLP) data for Darwin, Australia [Basnett and Parker, 1997].

Figure 1. Map of Java and vicinity showing PDSI grid (5-10°S-105-115°E) and tree-ring (blue, D’Arrigo et al. [1994]) and coral δ18O records (pink, Moore [1995]) used as predictors in PDSI regression model over 1870–1988 [Cook and Kairiukstis, 1990]. Values show significant (90% C.L.) correlations between each proxy (at lag T and/or T+1) and Oct–Nov PDSI.
2.2. Tree-Ring and Coral Data

Nine individual Indonesian tree-ring series were used (Table 1 and Figure 1). Due to low replication at some locations, some sites were composited to develop more robust time series (Table 1 and Figure 1). Some of these records were previously shown to correlate with ENSO indices [D’Arrigo et al., 1994]. A few of these chronologies were included in previous ENSO reconstructions [Stahle et al., 1998; Mann et al., 2000], and several of the teak series, along with coral records, were used to reconstruct Indonesian SSTs (R. D’Arrigo et al., Reconstructed Indonesian warm pool SSTs from tree rings and corals: Linkages to the Asian monsoon and ENSO, submitted to Paleoceanography, 2005, hereinafter referred to as D’Arrigo et al., submitted manuscript, 2005). The coral series is a monthly skeletal $\delta^{18}O$ record from Lombok Strait (composited to create an annual series), which also relates to local SSTs (Table 1 and Figure 1) [Moore, 1995; D’Arrigo et al., submitted manuscript, 2005]. After composing some of the chronologies, 5 tree-ring series and 1 coral series were used to develop the reconstruction.

3. The Java PDSI Reconstruction

The October-November season optimized the PDSI signal in the combined proxies. These months fall within the period when rainfall over Java is best linked to ENSO [Haylock and McBride, 2001; van Oldenborgh, 2002; Aldrian and Susanto, 2003]. Variability in precipitation over Java is due in part to orographic effects, and there is a moisture gradient with drier conditions in eastern Java. This variability can complicate efforts to reconstruct drought and pinpoints the need for more diverse proxy data coverage. Our reconstruction is thus best representative of conditions in central-east Java and vicinity. Details on reconstruction methodology are provided as auxiliary material.

Table 1. Description of Teak Chronologies Used to Reconstruct Java PDSI

<table>
<thead>
<tr>
<th>Site Name</th>
<th>Site Code</th>
<th>Latitude</th>
<th>Longitude</th>
<th>Full Period</th>
<th>No of Series</th>
<th>Mean r</th>
<th>Period &gt; 10 Series</th>
</tr>
</thead>
<tbody>
<tr>
<td>Randublatung</td>
<td>RAN</td>
<td>7.06 S</td>
<td>111.22 E</td>
<td>1925–2004</td>
<td>13</td>
<td>0.47</td>
<td>1938–2004</td>
</tr>
<tr>
<td>Bekutuk</td>
<td>BEK</td>
<td>7.07 S</td>
<td>111.22 E</td>
<td>1668–2004</td>
<td>20</td>
<td>0.40</td>
<td>1834–2004</td>
</tr>
<tr>
<td>Gubug Payung</td>
<td>GUB</td>
<td>7.05 S</td>
<td>111.29 E</td>
<td>1864–2004</td>
<td>18</td>
<td>0.34</td>
<td>1879–2004</td>
</tr>
<tr>
<td>Donoloyo Cagar Alam</td>
<td>DNLY</td>
<td>7.52 S</td>
<td>111.12 E</td>
<td>1714–2004</td>
<td>13</td>
<td>0.39</td>
<td>1746–2004</td>
</tr>
<tr>
<td>Pagerwunung Darupono</td>
<td>DNRP</td>
<td>7.02 S</td>
<td>110.16 E</td>
<td>1776–2004</td>
<td>19</td>
<td>0.39</td>
<td>1820–2004</td>
</tr>
<tr>
<td>Klangon Natural Forest</td>
<td>KLA</td>
<td>7.30 S</td>
<td>111.47 E</td>
<td>1707–2004</td>
<td>15</td>
<td>0.49</td>
<td>1812–2004</td>
</tr>
<tr>
<td>Saradan</td>
<td>SAR</td>
<td>7.29 S</td>
<td>111.42 E</td>
<td>1689–2000</td>
<td>30</td>
<td>0.46</td>
<td>1812–2000</td>
</tr>
<tr>
<td>Bigin</td>
<td>BIG</td>
<td>7.10 S</td>
<td>113.14 E</td>
<td>1839–1995</td>
<td>20</td>
<td>0.59</td>
<td>1853–1995</td>
</tr>
<tr>
<td>Muna</td>
<td>MUN</td>
<td>5.30 S</td>
<td>123.00 E</td>
<td>1564–1995</td>
<td>39</td>
<td>0.42</td>
<td>1673–1994</td>
</tr>
<tr>
<td><strong>Composite Series</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>SAKLA</td>
<td>-</td>
<td>-</td>
<td>-</td>
<td>1689–2004</td>
<td>45</td>
<td>0.46</td>
<td>1759–2004</td>
</tr>
<tr>
<td>RBGB</td>
<td>-</td>
<td>-</td>
<td>-</td>
<td>1668–2004</td>
<td>71</td>
<td>0.36</td>
<td>1834–2004</td>
</tr>
<tr>
<td><strong>Coral Data</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Lombok Strait</td>
<td>LOM</td>
<td>8.15 S</td>
<td>115.30 E</td>
<td>1782–1989</td>
<td>-</td>
<td>-</td>
<td>-</td>
</tr>
</tbody>
</table>

*Mean r = mean inter-series correlation derived from program COFECHA, and is an indication of common signal among tree-ring series within a chronology [Holmes, 1983]. Period > 10 series is an arbitrary cutoff for acceptable replication. SAKLA = composite of KLA and SAR; RBGB = composite of RAN, BEK, GUB and BIG. Series used in final reconstruction are: (tree rings): DNLY, DNRP, MUNA, SAKLA, and RBGB; (coral): Lombok $\delta^{18}O$ [Moore, 1995].

[7] The PDSI reconstruction explains 38% of the variance for the most replicated (1840–1988) period, dropping to 26% for the least replicated (1787–1819) period (Figure 2). Although these values are not very high, the statistical results (Figure 2) demonstrate that the model is robust and well-calibrated.

Figure 2. (a) Reconstructed (black) and instrumental (red) Oct–Nov Java PDSI over 1787–1888 and 1870–2004, respectively. Reconstruction scaled to instrumental data over the calibration period. Historical droughts identified from sea salt records (vertical bars, Quinn et al. [1978], 1844–1976) and planter’s records and rain measurements (triangles, van Bemmelen [1916], 1833–1914) are highlighted. White = reconstructed PDSI > 0; red = PDSI between 0 and –3; blue = PDSI values < –3.0. Horizontal green dashed line highlights –3.0 PDSI level. (b) and (c) Calibration and verification statistics for nested reconstruction models. $r^2$ = square of multiple correlation coefficient adjusted for loss of degrees of freedom; CE = Coefficient of Efficiency. CE’s greater than zero indicate model skill [Cook and Kairiukstis, 1990]. ST = Sign test [Fritts, 1976] shows ratio of agreement/disagreement; all values are significant above 99% level.
reasonably robust over its length. The reconstruction reveals PDSI variations during the past two centuries, with good agreement between instrumental and reconstructed values (Figure 2). The reconstructed PDSI ranges from −4.7 to 6.0. Analysis of the frequency of extreme dry/wet events per decade (exceeding ±1 std. dev. or SD) did not indicate any significant trends over the past two centuries. The PDSI series correlate significantly with several tropical climate indices (Table 2). The strongest correlations with the reconstruction are found with Niño-3 and 3.4 SSTs (for both, $r = -0.56$, 99.9% level). Consistent with these observations, the strongest coherence between the reconstructed and actual PDSI is within the 5–6 year ENSO bandwidth (Figure 3a). Evidence for a large-scale ENSO signal is further supported by multi-taper method spectral analysis [Mann and Lees, 1996], which also shows peaks consistent with ENSO (e.g., at $\pm 5–6$ and $2–3$ years; Figures 3b and 3c). Peaks in the coherency spectrum and full reconstruction at $\approx 10$ years (Figures 3a and 3c) may correspond to ENSO-like quasi-decadal variability [Allan, 2000].

### 4. Comparison With Javanese Historical Records

The PDSI reconstruction was compared to a historical record of Jakarta rain days [Konnen et al., 1998]. This record begins in 1829 (with a gap from 1851–1863) and has been extended using modern data. Correlation is strongest for June to October ($r = 0.56$, 99.9% level). This coherence is good considering that rainfall in Jakarta, northwestern Java, can display differences from conditions at the tree sites (mainly in central and east Java). We also compare our reconstruction to two independent listings of drought for sites (mainly in central and east Java). We also compare our reconstruction to two independent listings of drought for sites (mainly in central and east Java). The most negative departures following the 1809 unknown volcanic event [Dai et al., 1991], [1810 ($-4.0$), 1811 ($-4.72$) and 1812 ($-3.54$)] which has been linked with unusual tropical cooling [Chenoweth, 2001; D’Arrigo et al., submitted manuscript, 2005; R. Wilson et al., 250 years of reconstructed and modeled tropical temperatures, submitted to Journal of Geophysical Research, 2005]. There is a negative response following the Tambora, Indonesia 1815 eruption [1816 ($-1.48$), 1817 ($-2.10$), 1818 ($-3.12$)]. Reconstructed values following such events may relate to the proxies’ response to cooling as well as drought. However, ENSO warm events may have also occurred during

**Table 2. Comparison of Reconstructed and Instrumental Java PDSI With Tropical Indices for Seasons of Optimum Correlation**

<table>
<thead>
<tr>
<th>Index</th>
<th>Reconstruction</th>
<th>Years</th>
<th>Instrumental</th>
<th>Years</th>
<th>Reference</th>
</tr>
</thead>
<tbody>
<tr>
<td>NINO4</td>
<td>−0.50 (Dec–Jan)</td>
<td>132 yr</td>
<td>−0.56 (Oct–Nov)</td>
<td>130 yr</td>
<td>5S–5N, 160E–150W</td>
</tr>
<tr>
<td>NINO3</td>
<td>−0.56 (Dec–Jan)</td>
<td>132 yr</td>
<td>−0.55 (Nov–Dec)</td>
<td>130 yr</td>
<td>5S–5N, 120W–170W</td>
</tr>
<tr>
<td>NINO12</td>
<td>−0.43 (Dec–Jan)</td>
<td>132 yr</td>
<td>−0.44 (Oct–Nov)</td>
<td>130 yr</td>
<td>0–105, 80W–90W</td>
</tr>
<tr>
<td>SOI</td>
<td>0.40 (Dec–Jan)</td>
<td>122 yr</td>
<td>0.61 (Jul–Aug)</td>
<td>130 yr</td>
<td>Ropelewski and Jones [1987]</td>
</tr>
<tr>
<td>Darwin SLP</td>
<td>−0.42 (Aug–Sep)</td>
<td>117 yr</td>
<td>−0.64 (Aug–Sep)</td>
<td>122 yr</td>
<td>Basnett and Parker [1997]</td>
</tr>
<tr>
<td>Jakarta PPT</td>
<td>0.32 (Aug–Sep)</td>
<td>125 yr</td>
<td>0.59 (Sep–Oct)</td>
<td>121 yr</td>
<td>GHCN</td>
</tr>
</tbody>
</table>

*Niño-SSTs from Kaplan et al. [1998]. All correlations significant above 99% confidence limit.

![Figure 3](image.png)
these periods [Ortlieb, 2000]. Only weakly negative departures are reconstructed during the severe 1790s ENSO and drought, which impacted much of Asia [Grove, 1998]. These values are prolonged, however spanning five years (1787–1791, mean PDSI = –0.95; SD 0.46; this event falls in the early, weakest part of the reconstruction). Drought is reconstructed during moderate-strong warm events of the 19th century: in ~1855 (–4.60), 1868 (–4.14), and 1877–78 (–3.31, –3.62) [Ortlieb, 2000] (Figure 2). These values are comparable to, or exceed, those during intense warm events of recent decades: in 1882 (–4.54), 1994 (–2.95), 1997 (–3.17) and 2002 (–2.39). Two of these (1855 and 1877–78) are also identified by Quinn et al. [1978] and van Bemmelen [1916] (Figure 2a). Except for departures during the unknown volcanic event (1810–12), the 1877–78 event is the only one in which reconstructed values fall below –3.0 for two consecutive years.

The greatest positive anomaly occurs in 1861, when the reconstructed value is 5.99; it falls within one of the wettest decades in the reconstruction (1856–1865). This decade coincides with the persistent “Civil War Drought” observed in tree-ring data over much of the USA, attributed to the effects of “La Niña”-like cold tropical Pacific SSTs [Cook and Krusic, 2004; Herweijer et al., 2006] (Figure 2).

5. Summary

Few high-resolution proxies exist for monsoon Asia. We have developed a reconstruction of PDSI for Java, Indonesia, that reflects past monsoon drought variability and the influence of ENSO in this region. Although short, spanning only two centuries, the reconstruction extends the available instrumental PDSI record, which is more uncertain in its early decades. Our results suggest strong coupling between Javanese drought and ENSO-related variability. This reconstruction should prove useful in future efforts to model Indonesian drought relationships with the Asian monsoon and ENSO over the past several centuries.

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