

REPORT OF UNDERWAY pCO<sub>2</sub> MEASUREMENTS IN SURFACE WATERS  
AND THE ATMOSPHERE DURING November-December 2001

(RVIB Nathaniel B. Palmer Cruise 01/6)

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**1. General**

**1-a) The Cruise:**

This cruise participated in the Cape Shirreff opening, and research led by the Principal Investigator Martin Visbeck (LDEO). It departed Punta Arenas, Chile on 9 November 2001, transitted south to the west side of the Antarctic Penninsula. The cruise track then proceeded south along the Penninsula, then retracing the track and along 62°S to near 40°W. The track then turned back toward Punta Arenas . The ship arrived in Punta Arenas on 1 December 2001. Our system was operated by Bruce Huber of LDEO.

This cruise was also the first use of the modified system containing the mass flow controller. Most air lines were also replaced on a major system overhaul by Colm Sweeney.

We have applied the revised algorithm for correcting for the time lag between water passing the remote temperature probe and arriving at our equilibrator. We move the pCO<sub>2</sub> data one scan earlier.

Note that the uncorrected sea surface temperature, that is without adjusting by one scan, showed a slightly smaller standard deviation of  $T_{eq} - T_{SG}$  Temp than the corrected file. Specifically these were 0.0473 uncorrected, and 0.0494 corrected. We chose the latter to maintain consistency with other data files.

**1-b) The pCO<sub>2</sub> data:**

The program which combines our raw pCO<sub>2</sub> data files with other shipboard data was not run on this cruise. The program which creates "JGOFS" type files was run, however, so we used a combination of those files and our raw data files to create our final file. The raw pCO<sub>2</sub> data file contained entries for Latitude and Longitude (all -99.999) in each of the unknown data lines, but NOT in the standard data lines. We used a text editor to eliminate these lat/lon data and modified both **procpc2.prg** and the format of the database file **pc2proc.dbf** to accomodate these changes.

The program **makepco2.prg** combines our pCO<sub>2</sub> data with the ship's data from the JGOFS-type file. The program **editpco2.prg** then performs the editing.

Thirty-two seawater and one air observation are made between standard sets. Each observation involves flowing the equilibration gas for 180 seconds. We also flow the air for 180 seconds before stopping for an observation. Details on these and other aspects of the analysis system are explained later.

**Standards:**

Cylinder	Concentration (ppm)
Nitrogen (UHP)	0.0
CC02231	236.29
CC02235	105.20
CC15551	362.8
CA02205	495.18

A 4<sup>th</sup> order calibration curve is used with all 5 standards, as described in Section 2.

Standards flow for about 70 seconds before an observation is made.

**1-c) AIR data:**

There are 346 air observations. We apply a filter based on a running mean of 9 observations, and reject values outside one standard deviation of that mean. We recalculate a running mean of 9 values, this time using only those which have passed the first filter, and reject those more than one standard deviation away from the mean. There were many data points rejected here, and only 164 remained. The result was **371.62 ± 0.20 ppm (N=164)**.

To calculate  $\Delta p\text{CO}_2$  we use the Global View CO<sub>2</sub> data, observed through the end of 2001. This value goes into the "vco2\_air" and "gvvco2\_air" variables in the **0106sfc** database. Using the values coincident in time with our observations yields **369.54±0.05(N=4)**.

For this cruise we also calculate a sea-air pCO<sub>2</sub> difference using the LDEO measured and filtered air concentration and store it as a separate variable in the database. The comparison between the two values is as follows:

LDEO vCO<sub>2</sub> air - Globalview vCO<sub>2</sub> Air = **+2.08**. So our value is larger by 2.08 ppm

**1-d) Editing:**

In this data set editing was far more challenging than it normally is. Fortunately being a relatively short cruise it was not overly difficult. Because there are no water flow data in the JGxxx.dat type file, that primary editing tool was not available. However, because our system was not started until the ship had passed south of the Antarctic Front,

and was shut down before passing north of it on the return leg, we were able to apply a simple threshold of 2°C as a rejection criteria. Each time this value was reached, we went to 4 scans before and 6 after the temperature returned under 2°C to reject values. This was surprisingly effective in removing obvious outliers. "Hand edits" were still required, mostly to extend the rejections after a positive temperature spike, and to locate those spikes which did not reach all the way to 2°C. Of the **11,048** observations of seawater pCO<sub>2</sub> we rejected **1400**, retaining **9,648** records. The program **editpco2.prg** performs this editing operation.

To create a **0106sfc.dbf** surface format file, we use only ACCEPTED pCO<sub>2</sub> values, but ALL the data are retained in the pco2data.dbf database in the subdirectory: \NBPALMER\NBP01\_6\SURFACE\PCO2DATA. We estimate the air value from the Globalview CO2database (as explained under 1-d) AIR data . The program **make0106.prg** performs this process. The program also adds salinity from the edited salinity file saldata.dbf in the salinity subdirectory (see other data below).

#### **1-f) Estimating Temperatures:**

No temperature data were missing.

#### **1-g) Other Data:**

Salinity: There is a program, **procsal.prg** in the subdirectory salinity which performs editing of the salinity data. The system uses a mean and standard deviation of 15 data points. There are five passes through the data. Three calculate mean and standard deviation, two perform edits. After each of the first 2 calculation passes, data points are rejected if outside 1 standard deviation. A minimum standard deviation of 0.02 is assigned. After the second editing pass (4th through the data), mean and standard deviation of the remaining data are calculated using only accepted values. Between accepted values the mean is unchanged. This file **saldata.dbf** is then used as the source for salinity in the master file **0106sfc.dbf**.

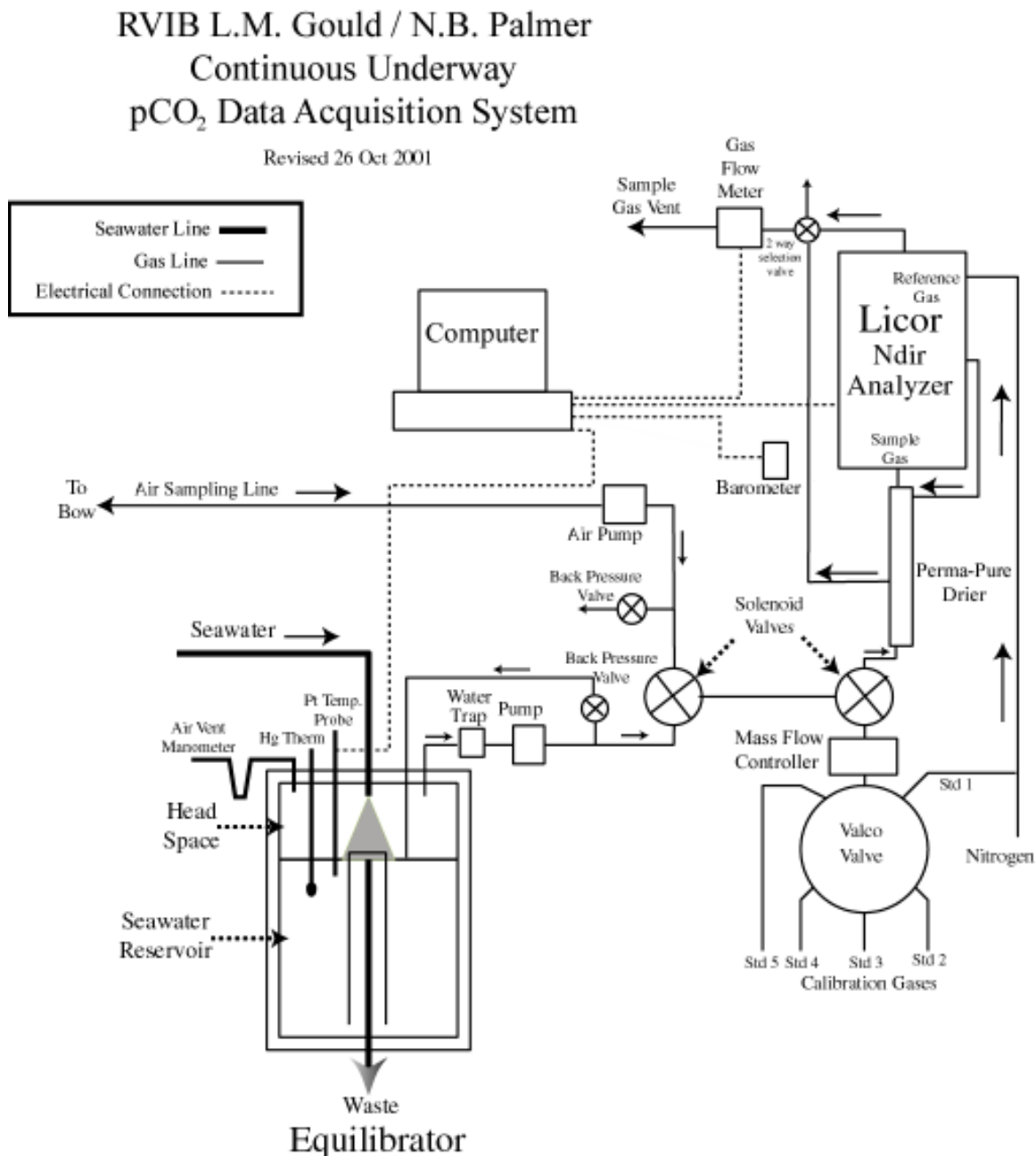
Wind speed units are meters seconds<sup>-1</sup>.

## 2. MEASUREMENTS OF pCO<sub>2</sub> IN SURFACE WATERS

### 2-a) The LDEO Underway System for Surface Water pCO<sub>2</sub> Measurements:

The system for underway measurements of pCO<sub>2</sub> in surface waters consists of a) a water-air equilibrator, b) a non-dispersive infra-red CO<sub>2</sub> gas analyzer and c) a data logging system. The measurement system is schematically shown in Fig. 1, and is similar with the one described in Bates et al. (1998). Each of these units and the data reduction procedures used will be described below. This shows small revisions performed by Colm Sweeney prior to this cruise.

Figure 1 - The underway pCO<sub>2</sub> system used for the measurements of pCO<sub>2</sub> in surface waters during the Southern Ocean JGOFS (AESOP) Program.



## **2-b) Water-air Equilibrator:**

The equilibrator has a total volume of about 30 liters and is equipped with a specially designed drain which maintains automatically the level of water in the equilibrator at a constant level at about half the height of the equilibrator leaving about 15 liters of headspace. Seawater from the ship's uncontaminated water line is continuously pumped into the equilibrator at a rate of about 10 liters/min, giving a mean residence time of water in the equilibrator of about 1.5 minutes. The headspace above the water serves as an equilibration chamber. A carrier gas (commonly marine air) is drawn into the chamber by a diaphragm pump, and exchanges CO<sub>2</sub> with a continuous flow of seawater sprayed into the chamber through a shower head. Because of large gas-water contact areas created by fine water droplets as well as gas bubbles in the pool of water, CO<sub>2</sub> equilibration between the carrier gas and seawater is achieved rapidly with a e-folding time of 2 to 3 minutes. Under normal operating conditions, the carrier gas in the equilibration chamber is pumped into the infra-red gas analyzer at a rate of about 50 ml/min. At this rate, the residence time of the carrier gas in the equilibration chamber is about 300 minutes, that is about 100 times as long as the equilibration time. Therefore, the carrier gas in the head space is always in equilibrium with water. The over all response time of the equilibrator system has been estimated to be of an order of several minutes. The large volume of water in the equilibrator is chosen in order to have a large thermal inertia of the equilibrator, so that the effects of room temperature changes on the equilibration temperature may be minimized. The temperature of water in the equilibrator is monitored continuously using a Guildline platinum resistance thermometer (readable to 0.05 °C) and recorded on the data logging computer. A calibrated mercury thermometer is also inserted in the equilibrator for testing the performance of the platinum thermometer.

At the gas intake end of the equilibrator, a flow indicator based on U-tube manometer is attached. This gives a visual confirmation for the fact that marine air is taken into the equilibration chamber at a desired flow rate. Since we operate the system with the equilibration chamber at the same pressure as the ambient room pressure, the total pressure, at which the gas was equilibrated, is measured using a precision electronic barometer (Setra Model 270, Action, MA) outside the equilibrator. This equilibration pressure is also logged on the computer.

The temperature and salinity of seawater at the in situ conditions were measured using a SeaBird Model SBE-21 thermosalinograph. The precision of the report temperature data has been estimated to be about 0.005 °C.

## **2-c) Infra-red CO<sub>2</sub> Gas Analyzer:**

The equilibrated gas was passed through a water trap (to collect aerosols and condensates), mass flow controller and a reverse flow naphion dryer (PermaPure flushed with pure nitrogen gas) to remove water vapor (to a level of -20°C), and was introduced into the IR sample cell at a rate of about 50 ml/min for CO<sub>2</sub> determinations. A LI-COR infra-red gas analyzer (Model 6251, Lincoln, NB) was used. After about 3 minutes of purging period, the gas flow was stopped and readings were recorded on the computer.

Although an electronic circuit was provided by the manufacturer in order to linearize the CO<sub>2</sub> response, it exhibited a few inflexions that deviated from linearity by a few ppm. Therefore, we chose not to use the outputs from the linearization circuit supplied by the manufacturer. Instead, we used five standard gas mixtures (one pure nitrogen and four CO<sub>2</sub>-air mixtures) during the expeditions, and established response curves using the raw output from the analyzer. The CO<sub>2</sub> concentrations in the gas mixtures were calibrated using the SIO standards determined by C. D. Keeling's group using the manometric method. The concentrations of CO<sub>2</sub> in the standard gas mixtures were summarized above.

During normal operations, each of the standard gas mixtures was passed through the analyzer for 70 to 90 seconds at a rate of about 60 ml/min. This replaced the IR analyzer cell completely with the new gas. The flow was stopped for 5 seconds and then a millivolt reading from the analyzer was taken and recorded. Samples of equilibrated air and marine air were pumped through the analyzer for 180 seconds (3 minutes) at a rate of about 50 ml/min to purge the previous sample in the IR cell. The flow was stopped for 5 seconds and a reading for the analyzer output was recorded. This procedure was intended to eliminate errors due to fluctuations of the dynamic pressure within the IR cell by irregular gas flow rates. The slow flow rates used for samples were required for the removal of water vapor using the PermaPure membrane dryer. Between two sets of calibration runs using the five standard gases, 6 to 20 samples were analyzed depending upon the stability of the IR analyzer.

#### **2-d) Data Logging System:**

The following values were recorded on a laptop computer. The sample locations were derived from a GPS positioning unit that is a part of our surface water pCO<sub>2</sub> system. The CO<sub>2</sub> readings for samples were recorded once every 3 minutes (180 seconds), and those for the standard gas mixtures once every 1.5 minutes.

Date,  
Time (GMT),  
Sample ID (standard gas cylinder numbers, seawater CO<sub>2</sub>, atmospheric CO<sub>2</sub>)  
Barometric pressure in the laboratory (to 0.1 mb)  
IR cell temperature,  
Gas flow rate in the IR cell (to 0.1 ml/min),  
Temperature of equilibration (to 0.01 °C),  
Analyzer output (millivolts to 0.1 mv)  
CO<sub>2</sub> concentration in dry gas sample (preliminary based on the last response curve), and  
pCO<sub>2</sub> (preliminary value based on the last response curve).

This cruise created only the JGOFS-type file, named JGxxx.DAT where xxx = Julian Date. There are two fields added to the file which are not identified in my documentation. The date list is as follows:

GMT Date (dd/mm/yy)  
 GMT Time (hh:mm:ss)  
 Latitude (dd.dddd)  
 Longitude (ddd.dddd)  
 Speed over Ground (knots)  
 HDOP (GPS accuracy indicator)  
 Heading (ddd.d°T)  
 Course Made Good (ddd.d°T)  
 PAR (microeinsteins m<sup>-2</sup> sec<sup>-1</sup>)  
 Conductivity (Siemens m<sup>-1</sup>)  
 Salinity (Practical Salinity Scale)  
 Water Depth (m)  
 Wind Speed (m sec<sup>-1</sup>)  
 Wind Direction (ddd°T)  
 Air Temperature (°C)  
 Relative Humidity (%)  
 Barometric Pressure (mbar)  
 Fluorometer #1 Voltage  
 Fluorometer #2 Voltage  
 unidentified #1  
 unidentified #2

**2-e) Data Reduction Procedures:**

The concentration of CO<sub>2</sub> in the sample was computed by the following way based on the millivolt reading and time of the reading. The millivolt reading taken for each of the five standard gases at the time of sample measurement was computed by linearly interpolating as a function of time using the readings taken before and after the respective standard gases were analyzed. This yields millivolt reading for each of the five standard gases at the time when the sample was analyzed. These five values were fit to a fourth-order polynomial equation (with five constants to be determined). This serves as the response curve. The CO<sub>2</sub> concentration in the sample was computed using the response curve that was established at the time of each sample analysis. This method has been demonstrated to yield more reliable CO<sub>2</sub> values compared with those computed, for example, using a least-squares fit of a quadratic or cubic functions to the five calibration points. The method described above yields atmospheric CO<sub>2</sub> values that are consistent with those reported for the South Pole and the Cape Grim by the Climate Monitoring and Diagnostics Laboratory/NOAA in Boulder, CO.

The partial pressure of CO<sub>2</sub> in seawater, (pCO<sub>2</sub>)<sub>sw</sub>, at the temperature of equilibration, T<sub>eq</sub>, in the unit of microatmospheres (µatm) was computed using the expression:

$$(pCO_2)_{sw} @ T_{eq} = (V_{CO_2})_{eq} \times (P_b - P_w), \dots \dots \dots [1]$$

- (V<sub>CO<sub>2</sub></sub>)<sub>eq</sub> = the mole fraction concentration (ppm) of CO<sub>2</sub> in the dried equilibrated carrier gas;  
 P<sub>b</sub> = the barometric pressure (that is equal to the total pressure of equilibration) in atmospheres; and  
 P<sub>w</sub> = the equilibrium water vapor pressure at T<sub>eq</sub> (°C) and salinity.

The water vapor pressure was computed using the following formulation;

$$P_w \text{ (atm)} = (1/760) \times (1 - 5.368 \times 10^{-4} \times \text{Sal}) \times \text{EXP}\{[0.0039476 - (1/\text{TK})]/1.8752 \times 10^{-4}\}, \dots\dots [2]$$

where Sal is salinity in PSU measured using the ship's thermosalinograph, and TK is the temperature of equilibration in °K.

The (pCO<sub>2</sub>)<sub>sw</sub> at the in situ temperature, T in situ, was computed using a constant value of 0.0423 % per °C for the effect of temperature (Takahashi et al., 1993):

$$(p\text{CO}_2)_{\text{sw}} @ T_{\text{in situ}} = (p\text{CO}_2)_{\text{sw}} @ T_{\text{eq}} \times \text{EXP}[0.0423 \times (T_{\text{in situ}} - T_{\text{eq}})].$$

The value for T<sub>in situ</sub> is taken to be the seawater temperature measured by the ship's thermosalinograph at the time of pCO<sub>2</sub> measurements. T<sub>eq</sub> is generally warmer than T<sub>in situ</sub> by 0.5 ~ 0.8 °C. Hence the temperature correction is normally less than 3% of pCO<sub>2</sub> values.

The over all precision of the reported pCO<sub>2</sub>)<sub>sw</sub> values has been estimated to be about ±1.5 uatm.

### **3. MEASUREMENTS OF pCO<sub>2</sub> IN THE ATMOSPHERE**

#### **3-a) Measurements:**

The air measurement system is shown schematically in Fig. 1. Uncontaminated marine air samples were collected about 10 m above the sea surface using a DEKORON tubing (1/4" i.d., Calco Inc., PA), a thin-wall aluminum tubing protected by plastic casing. The intake was located at the middle of the foremast about 10 m above the sea surface. A KNF Neuberger air pump that was located near the IR analyzer was used to pump air through the tubing and into the IR analyzer. Even when air samples were not analyzed, the pump was on all the time to keep the air flowing through the sampling line. For the analysis, the air sample was passed through a water trap and a drying column to remove water vapor (the same PermaPure column as used for the equilibrated gas) and introduced into the IR cell for CO<sub>2</sub> analysis at a rate of about 50 ml/min. After 3 minutes of purging the cell, the flow was stopped for 5 seconds and the IR millivolt output reading was recorded.

#### **3-b) Data Processing:**



The partial pressure of CO<sub>2</sub> in the air, (pCO<sub>2</sub>)<sub>air</sub>, was computed in the unit of microatmospheres (µatm) in the same way as that for seawater using Eq. [3] below:

$$(pCO_2)_{air} = (V_{CO_2})_{air} \times (P_b - P_w), \dots \dots \dots [3]$$

- (V<sub>CO<sub>2</sub></sub>)<sub>air</sub> = the mole fraction concentration (ppm) of CO<sub>2</sub> in the dried air sample;
- P<sub>b</sub> = the barometric pressure at sea surface in atmospheres; and
- P<sub>w</sub> = the equilibrium water vapor pressure at T<sub>in situ</sub> (°C) and salinity given by Eq. [2].

The precision of the atmospheric pCO<sub>2</sub> values have been estimated to be about ± 1 µatm.

#### 4. REFERENCES CITED

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