# Solid Earth Dynamics 

## Bill Menke, Instructor

## Lecture 12

## Today:

# more on glacial isostatic rebound 

## starting discussion of deformation

newton's law in a material atmospheric pressure
volumetric and shear strain

## Glacial Isostatic Rebound



$$
\begin{aligned}
& \rho_{C} L+\rho_{M} d=\rho_{I} H+\rho_{C} L \\
& \rho_{M} d=\rho_{I} H \quad d=\frac{1000}{3000} H=\frac{1}{3} H=1 \mathrm{~km}
\end{aligned}
$$

$n=1$
patterns with increasing number of wiggles

$\mathrm{n}=$ number of half-wavelengths

let's call them $P_{n}(\cos \theta)$<br>Legendre Polynomials

each decays with
characteristic decay time $\tau$
$n=1$

$n=3$
$\mathrm{t}=0$
$c_{n} P_{n}(\cos \theta)$
at a later time, t
$c_{n} \exp (-t / \tau) P_{n}(\cos \theta)$
formula for $\tau$ known, depends on $n$

$$
\begin{aligned}
& F=\frac{\rho R \sqrt{g R}}{\mu} \quad \text { dimensionless constant } \\
& f=\sqrt{\frac{R}{g}} \quad \text { time scale, units of } s \\
& \tau=\frac{f}{F} \frac{\left(2 n^{2}+4 n+3\right)}{n} \quad \text { decay time, units of } s
\end{aligned}
$$



$$
n=4 \text { gives } \tau=\frac{f}{F} \frac{\left(2 n^{2}+4 n+3\right)}{n} \approx 2200 \text { years }
$$

## because both viscous flow and gravity are linear processes

## if load $L_{A}$ leads to uplift $U_{A}$ and

if load $L_{B}$ leads to uplift $U_{B}$ then
if load $L_{A}+L_{B}$ leads to uplift $U_{A}+U_{B}$
strategy for dealing with complicated load, $\mathrm{L}(\theta)$
(A) figure out the proportionality between Load and uplift at time $\mathrm{t}=0$, for loads of shape $P_{n}(\cos \theta)$
at time $\mathrm{t}=0$
Load $L=P_{n}(\cos \theta)$
leads to uplift $U=c_{n} P_{n}(\cos \theta)$
only work for some "the right" $P_{n}(\cos \theta)$ s

## $c_{n}$ is known for this problem

$$
c_{n}=\frac{2 n+1}{2(n-1)}
$$

when $P_{n}(\cos \theta)$ are the ones shown previously
(B) approximate Load $\mathrm{L}(\theta, t=0)$ as as sum of the functions $P_{n}(\theta)$

$$
f(\theta) \approx \sum_{n=1}^{N} b_{n} \underbrace{P_{n}(\cos \theta)}_{\text {need to find } b_{n}}
$$

(C) invoke superposition:
each $c_{n}$ decays independently with time
$c_{n}$ becomes $c_{n} \exp (-t / \tau)$
(D) perform the summation

$$
R(\theta, t=0) \approx \sum_{n=1}^{N} b_{n} c_{n} \exp (-t / \tau) P_{n}(\cos \theta)
$$

I cooked up the b's that give this load



## uplift

time $=0$

time $=3000 \mathrm{yrs}$



## during the Ice Age


today
uplift



## linear elasticity

and

Newton's Law in fluids

## elasticity: reversible deformation

during

before


## linear elasticity: deformation proportional to force


double force double deformation


## force per unit area

 is called traction, Tpositive when outward pointing
traction is always normal to surface


## fluid

## all (nearby) surfaces have the same traction


tractoin in fluid: minus the pressure, p

Newton's Law (horizontal motion only)

pressure force


$$
\begin{gathered}
F_{F}=A p_{L} \quad F_{R}=-A p_{L} \\
F=F_{L}-F_{R}=A\left(p_{L}-p_{R}\right)
\end{gathered}
$$

volumetric force, like gravity

$$
\begin{aligned}
& F_{F}=A p_{L} \\
& F_{R}=-A p_{L} \\
& F=f A \Delta x
\end{aligned}
$$

Mass


$$
M=\rho A \Delta x
$$

## Acceleration



$$
a=\frac{d^{2} u}{d t^{2}}
$$

Newton's law $\quad F=M a$

$$
\begin{array}{ll}
F=A\left(p_{L}-p_{R}\right) & A\left(p_{L}-p_{R}\right)+f A \Delta x=\rho A \Delta x \frac{d^{2} u}{d t^{2}} \\
F=f A \Delta x & \frac{\left(p_{L}-p_{R}\right)}{\Delta x}+f=\rho \frac{d^{2} u}{d t^{2}} \\
M=\rho A \Delta x & f-\frac{d p}{d x}=\rho \frac{d^{2} u}{d t^{2}}
\end{array}
$$

## Linear Elasticity in a fluid

$$
\frac{\Delta V}{V}=-c \Delta p
$$

fractional change in volume is proportional to pressure

Linear Elasticity in a fluid

$$
\begin{array}{ll}
\text { or if you prefer } & \frac{\Delta \frac{M}{\rho}}{\frac{M}{\rho}}=-c \Delta p \\
\text { with } \mathrm{V}=\frac{M}{\rho} & \\
\rho \Delta \rho^{-1}=-c \Delta p & \\
\rho \frac{d \rho^{-1}}{d p}=-c & -\rho \rho^{-2} \frac{d \rho}{d p}=-c
\end{array} \rho^{-1} \frac{d \rho}{d p}=c
$$

## what's atmospheric pressure do with height?

## 


pressure




Linear Elasticity in an isothermal ideal gas

$$
\begin{array}{ll}
P V=n R T & \begin{array}{l}
n=\text { number of moles } \\
m=\frac{n}{V} \frac{R T}{1}
\end{array} \begin{array}{l}
\text { mass }=n m
\end{array} \\
P=\frac{n m}{V} \frac{R T}{m} & \rho=\frac{\text { mass }}{\text { volume }}=\frac{n m}{V}
\end{array}
$$

Linear Elasticity in an isothermal ideal gas

$$
\begin{array}{rlrl}
\mathrm{p} V & =n R T & \mathrm{p}=\frac{R T}{m} \rho \\
p & =\frac{n}{V} \frac{R T}{1} & \rho=\frac{m}{R T} p \\
p & =\frac{n m}{V} \frac{R T}{m} &
\end{array}
$$




## Shear forces in solids

Shear forces in solids


Shear forces in solids
undeformed object

deformed object

displacement

a point initially at $(x, y)$
moves to $(x+u, y+v)$
volumetric strain: change in volume

volumetric strain: change in volume

$$
(\Delta x+\Delta u, \Delta y+\Delta v)
$$


change in area $\Delta A=\Delta x \Delta v+\Delta y \Delta u$
volumetric strain: change in volume

$$
(\Delta x+\Delta u, \Delta y+\Delta v)
$$


$(0,0)$
change in area

$$
\Delta A=\Delta x \Delta v+\Delta y \Delta u
$$

change in area

$$
\frac{\Delta A}{A}=\frac{\Delta v}{\Delta y}+\frac{\Delta u}{\Delta x}
$$

no change in area

$$
0=\frac{d v}{d y}+\frac{d u}{d x}
$$

no change in volume

> no change in area

$$
\begin{gathered}
0=\frac{d v}{d y}+\frac{d u}{d x} \\
\text { called }
\end{gathered}
$$ shear strain

$(0,0)$
volumetric strain: change in volume

$$
(\Delta x+\Delta u, \Delta y+\Delta v)
$$


easy way to satisfy no change in area

$$
0=\frac{d v}{d y}+\frac{d u}{d x}
$$

$u$ only varies with $y$
$v$ only varies with $x$

## undeformed object


undeformed object
shear strain only


