

Solid Earth Dynamics

Bill Menke, Instructor

Lecture 22

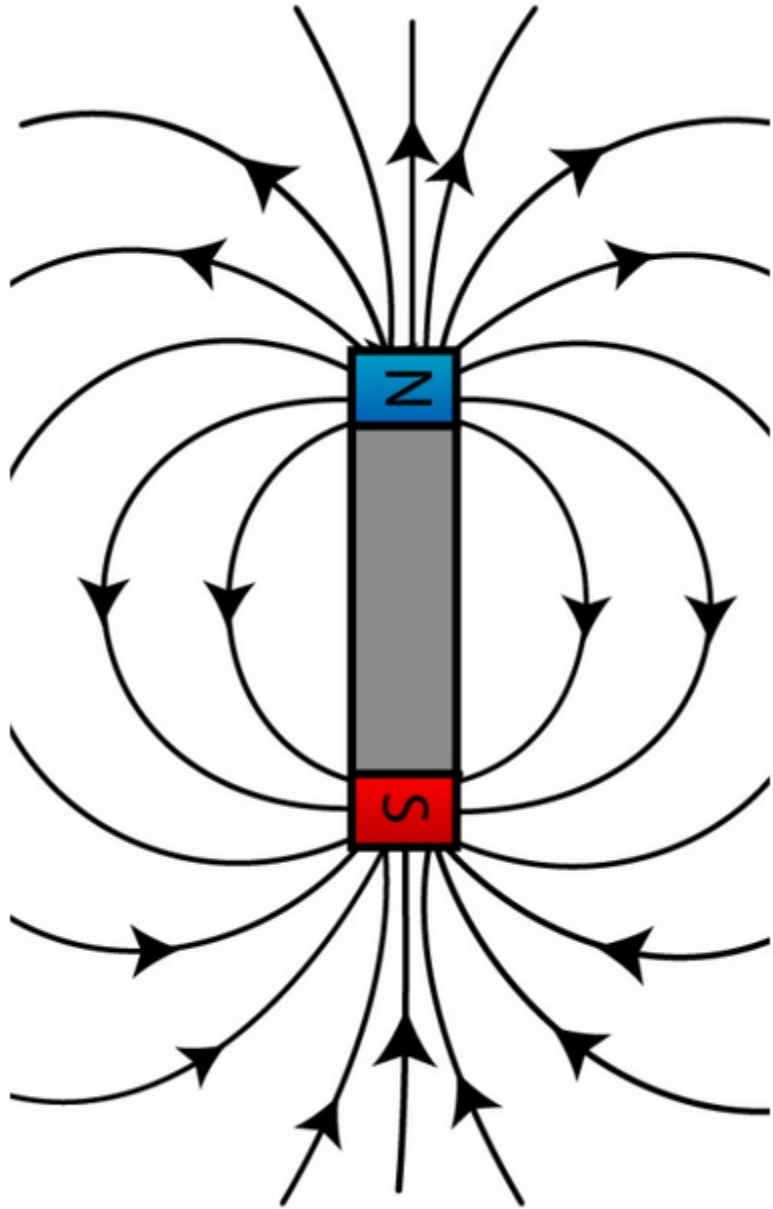
Today:

Geomagnetism:

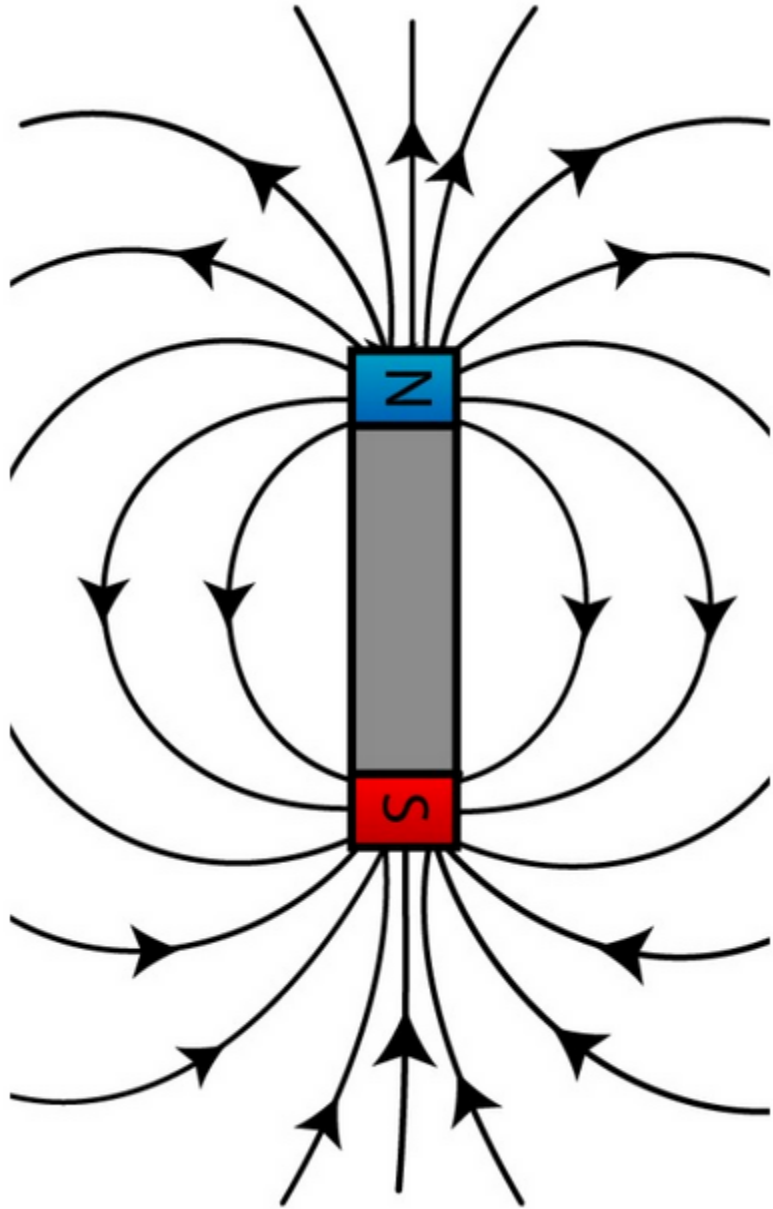
Paleomagnetism

Return to the Dynamo

Paleomagnetism



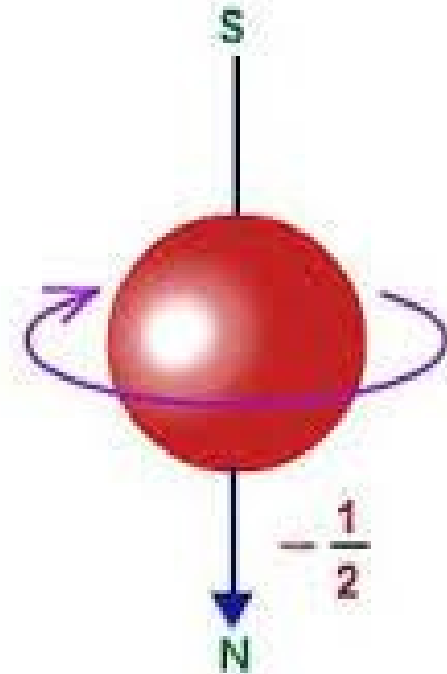
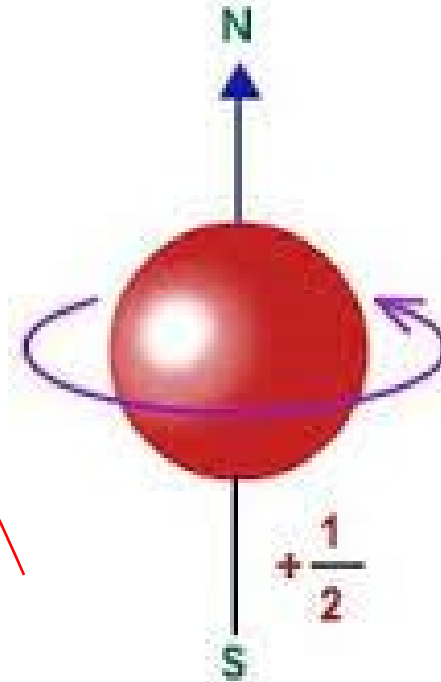
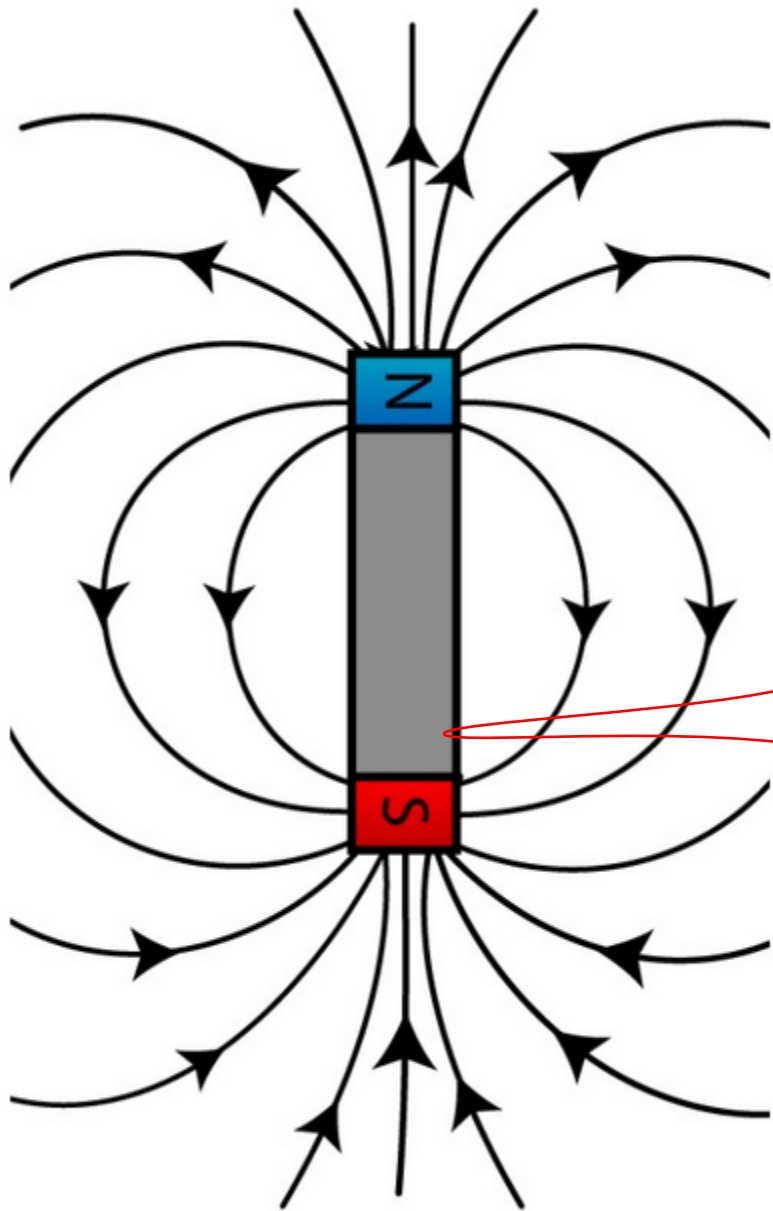
dipole field around a bar magnet

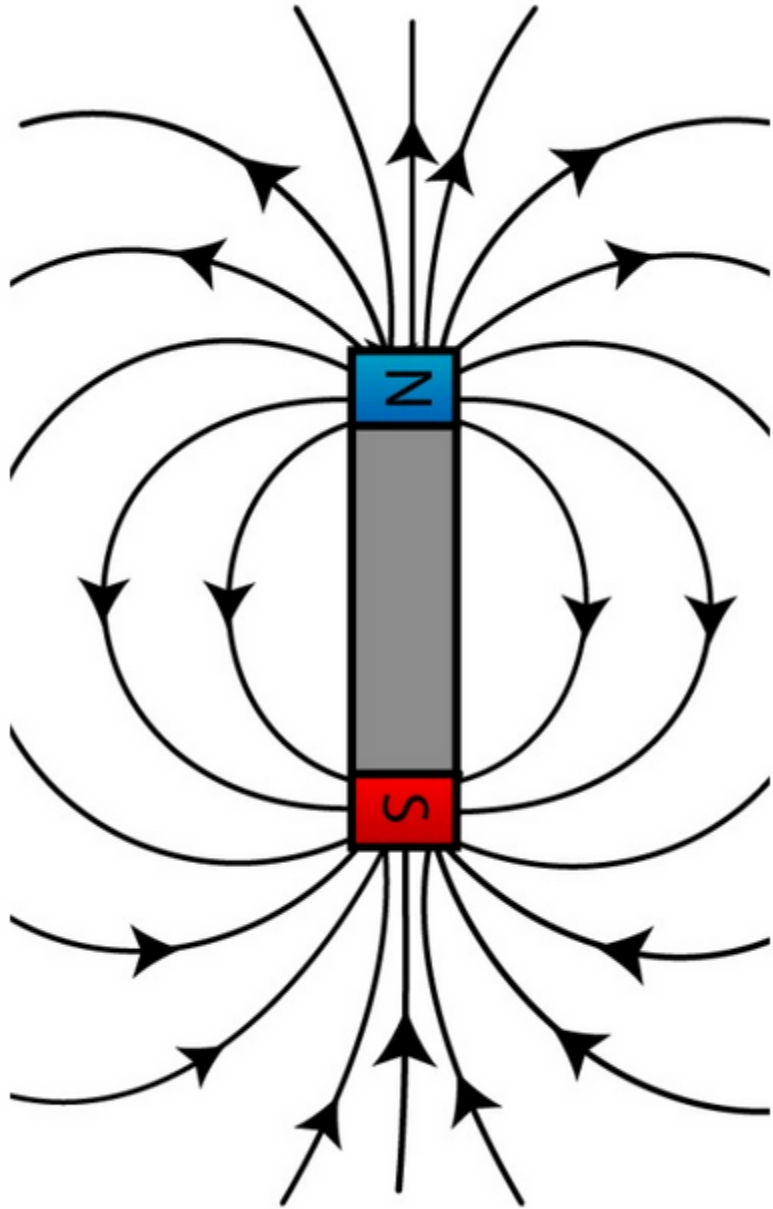


where does this magnetic field come from?

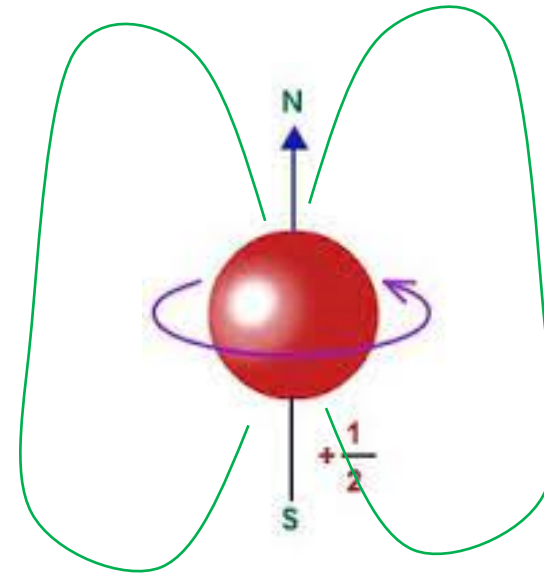
No (obvious) currents to make a solenoid

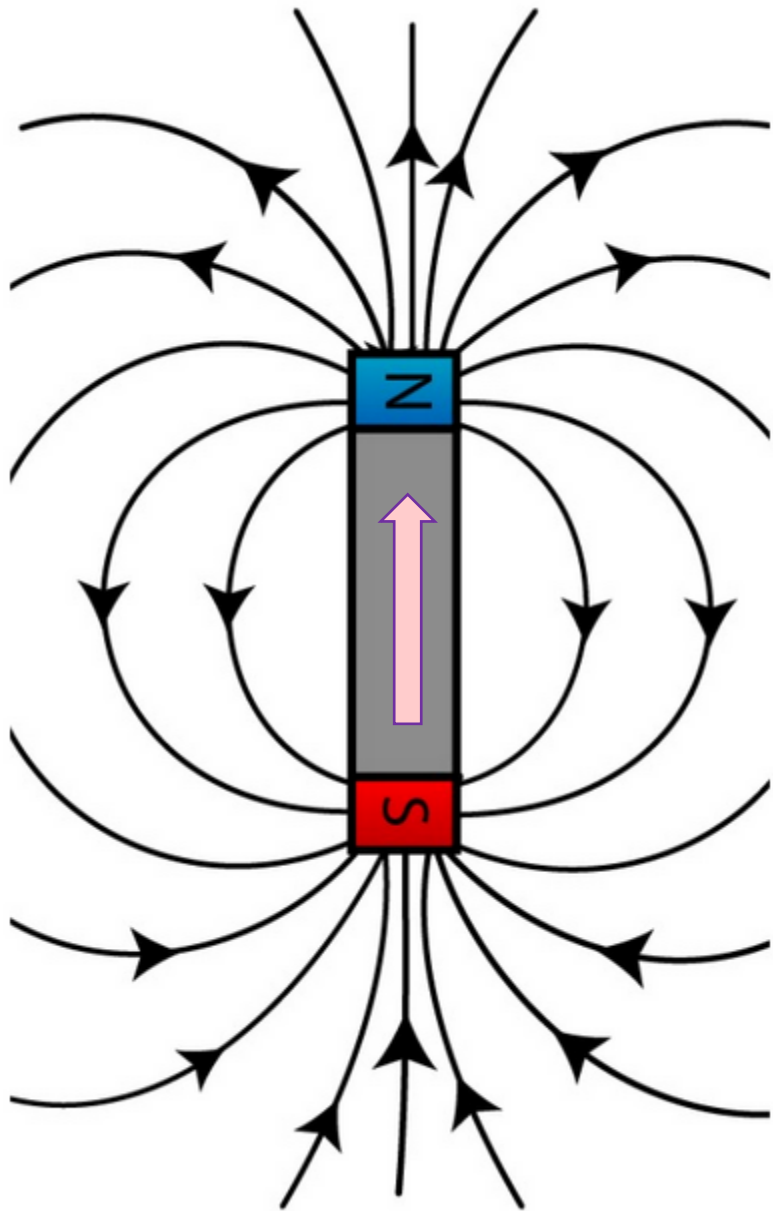
electrons have “spin”
(intrinsic angular momentum)





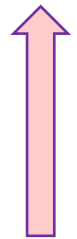
everything with spin
generates a dipolar
magnetic field

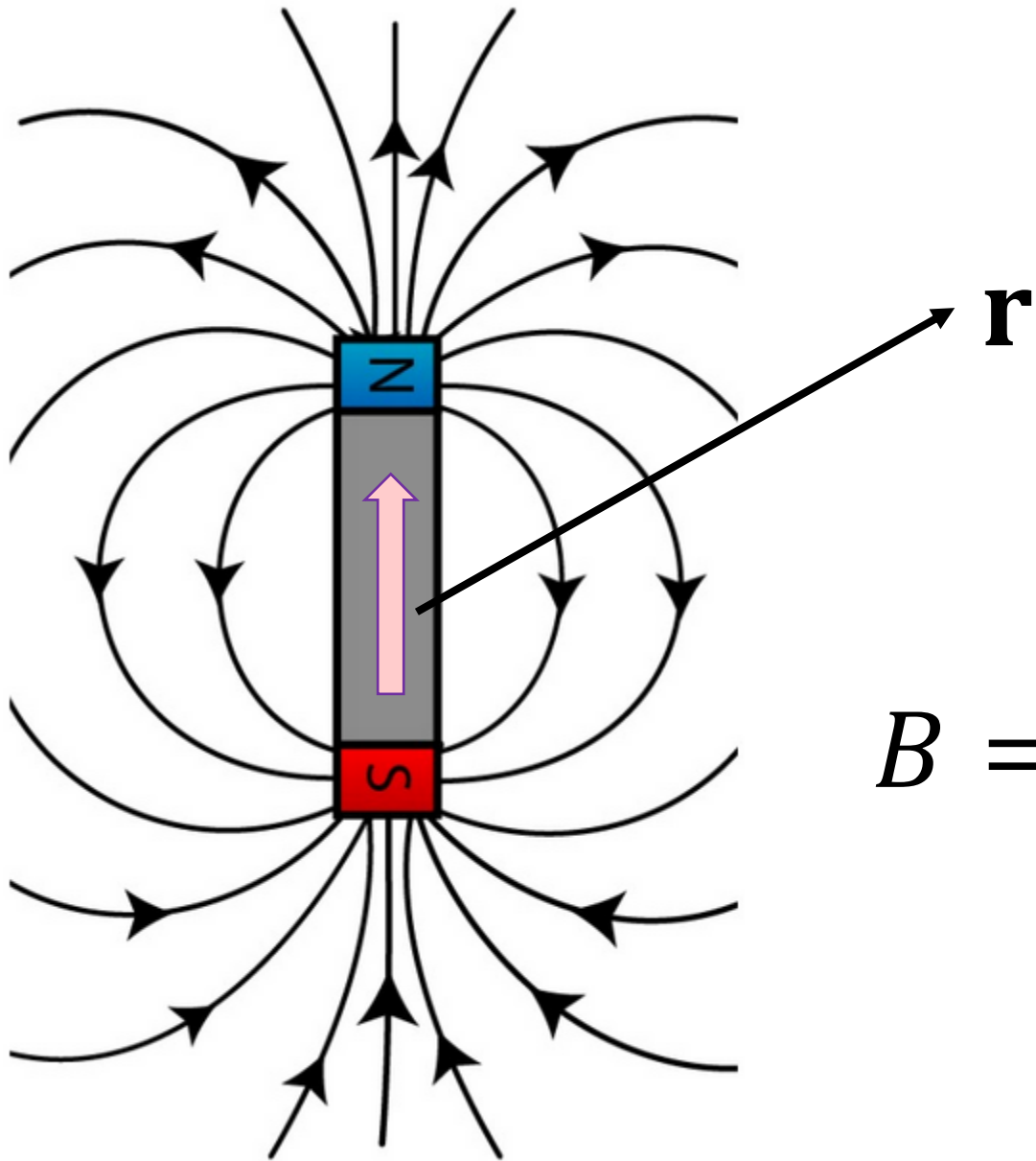




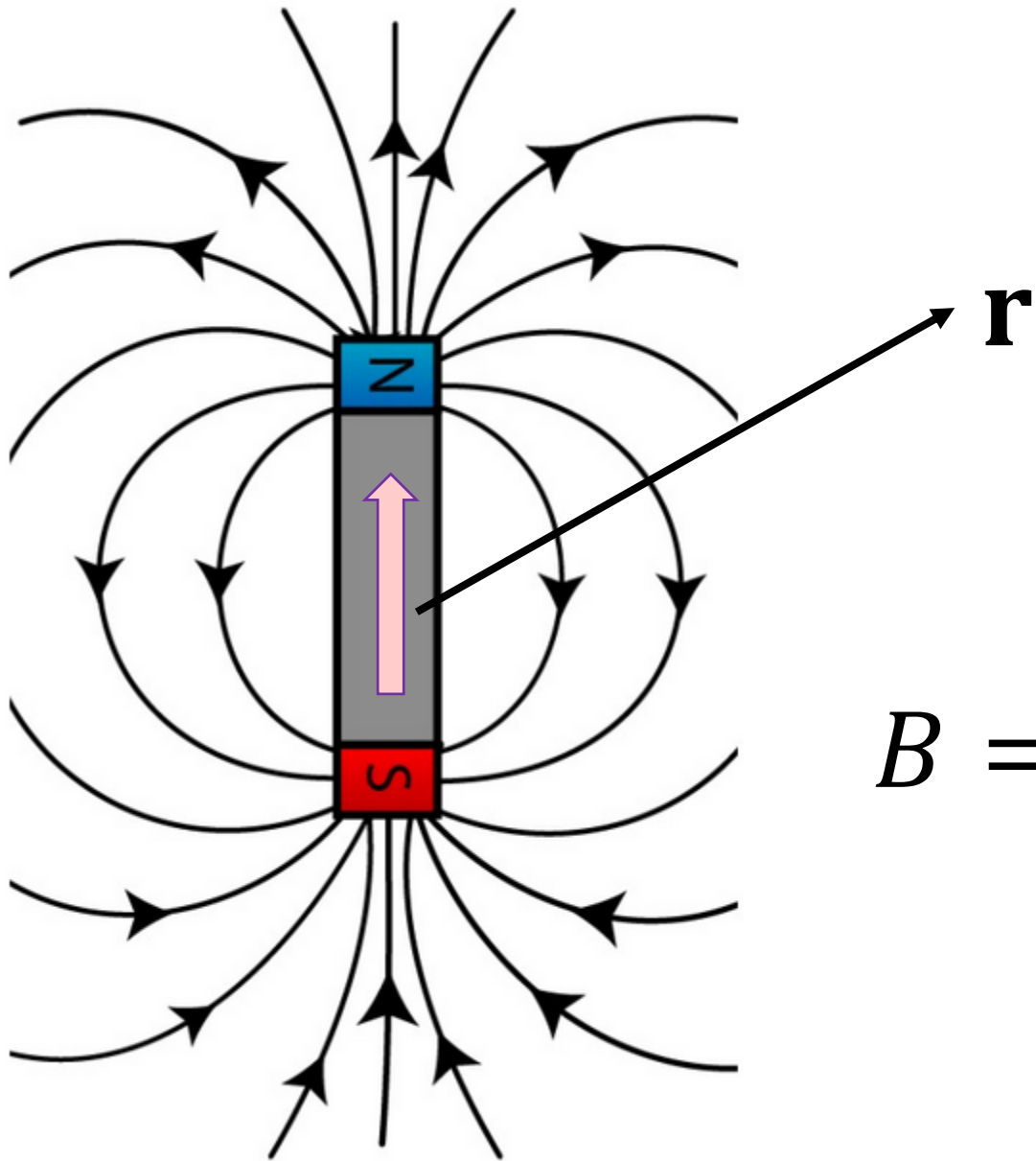
described by

magnetization vector \mathbf{M}



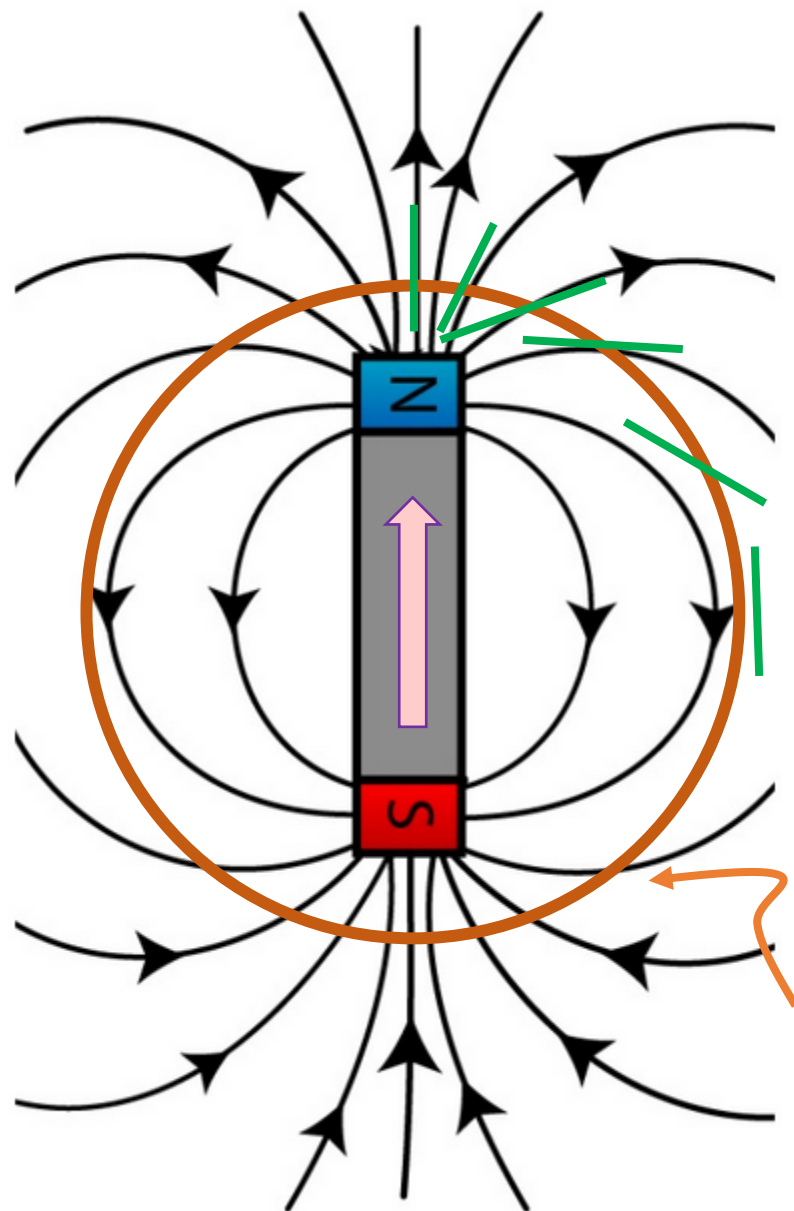


$$\mathbf{B} = \frac{\mu_0}{4\pi} \left(\frac{3\mathbf{r}(\mathbf{r} \cdot \mathbf{M}) - \mathbf{M}}{r^3} \right)$$



$$\mathbf{B} = \frac{\mu_0}{4\pi} \left(\frac{3\mathbf{r}(\mathbf{r} \cdot \mathbf{M}) - \mathbf{M}}{r^3} \right)$$

inverse cubed law



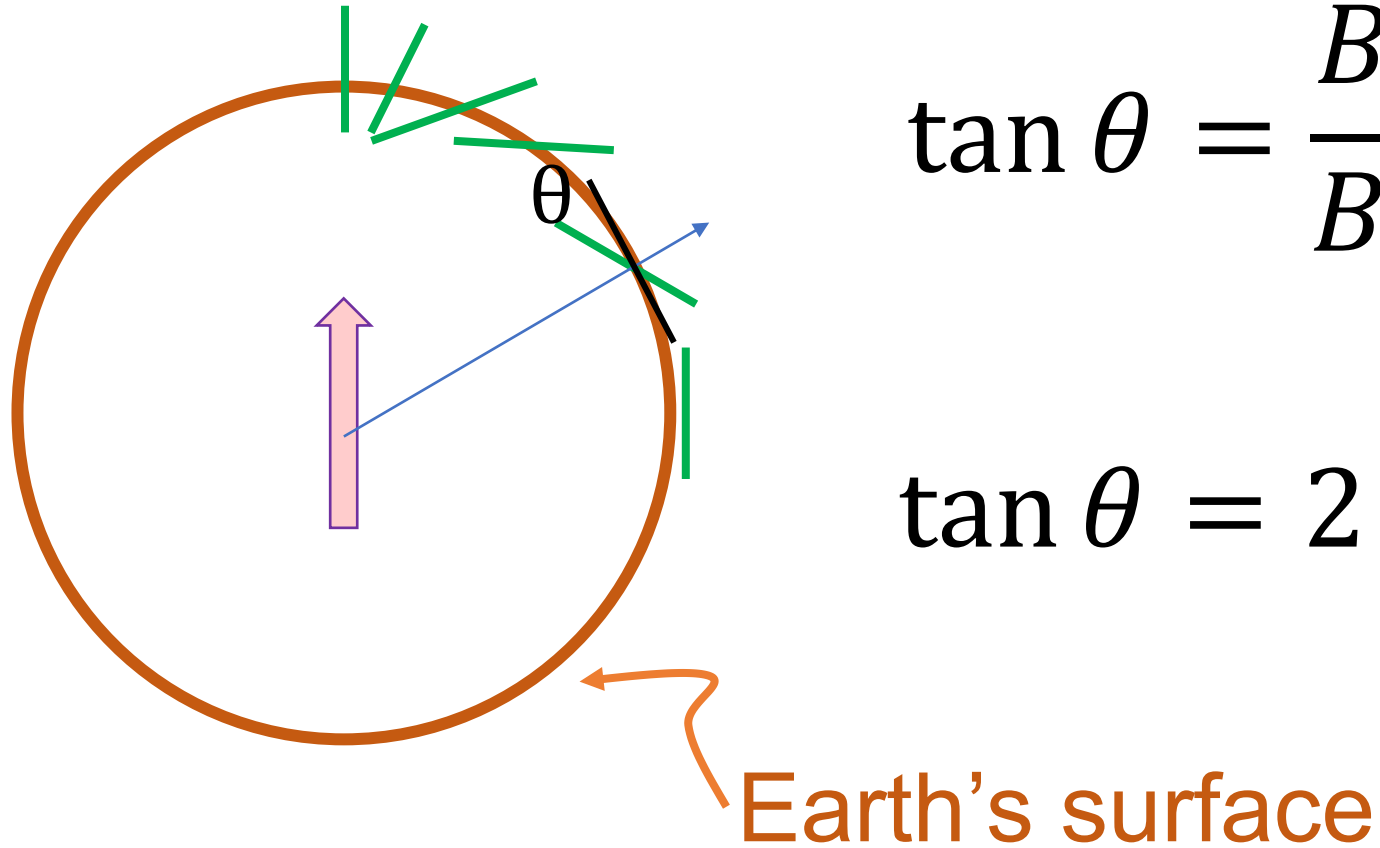
magnetic field
not horizontal

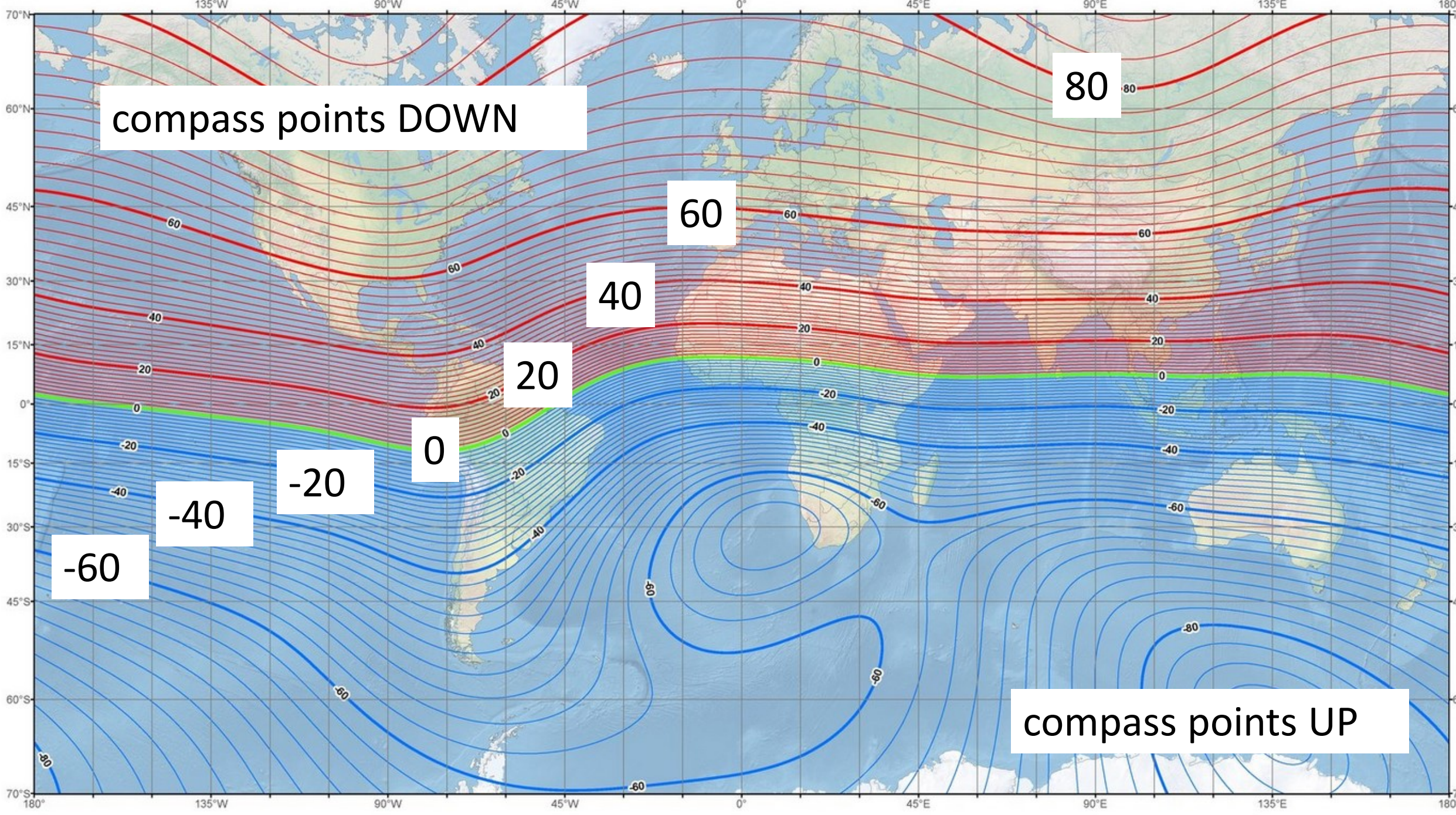
Earth's surface

formula for inclination

$$\tan \theta = \frac{B_V}{B_H}$$

$$\tan \theta = 2 \tan(\text{latitude})$$





compass points DOWN

80

60

40

20

0

-20

-40

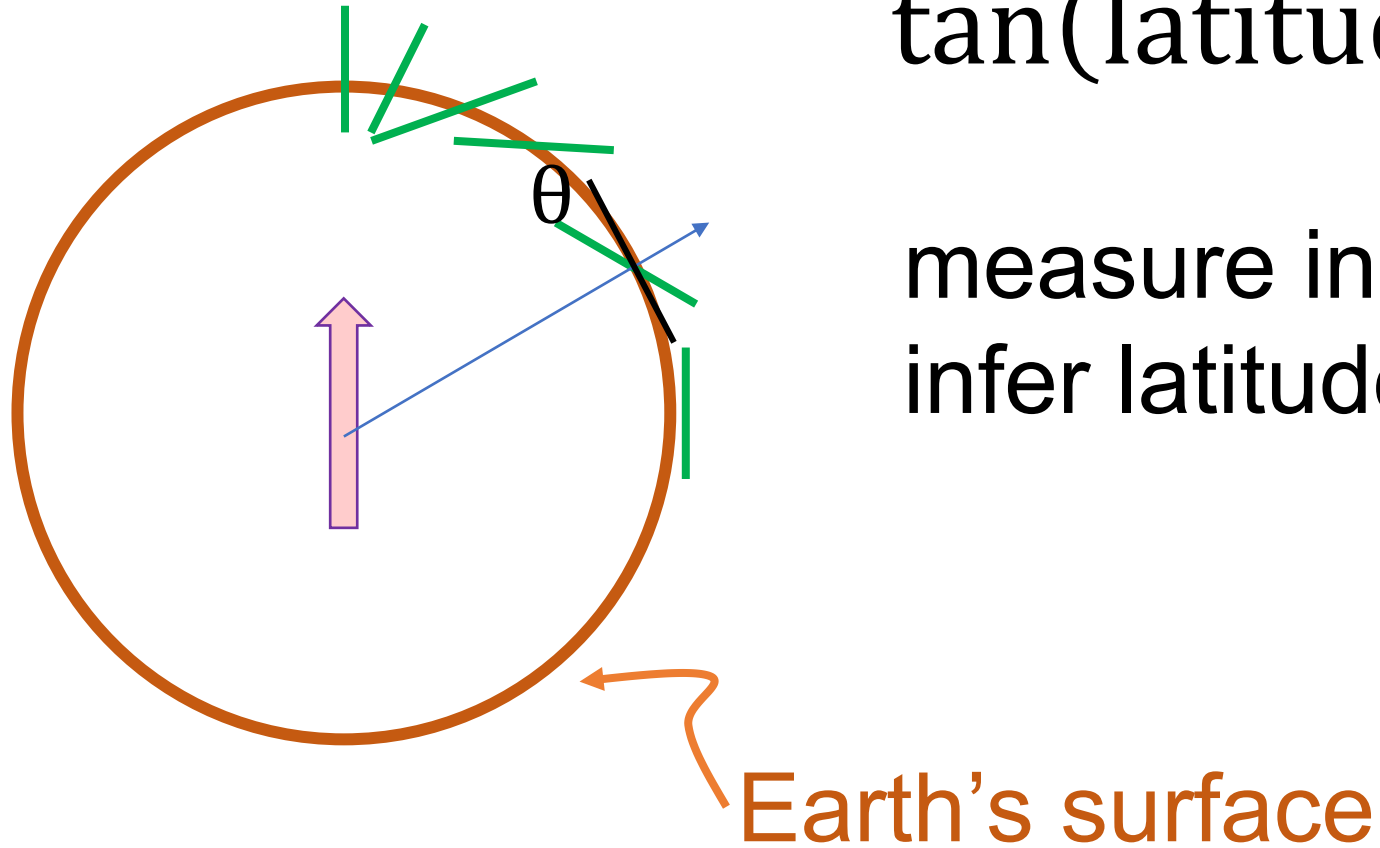
-60

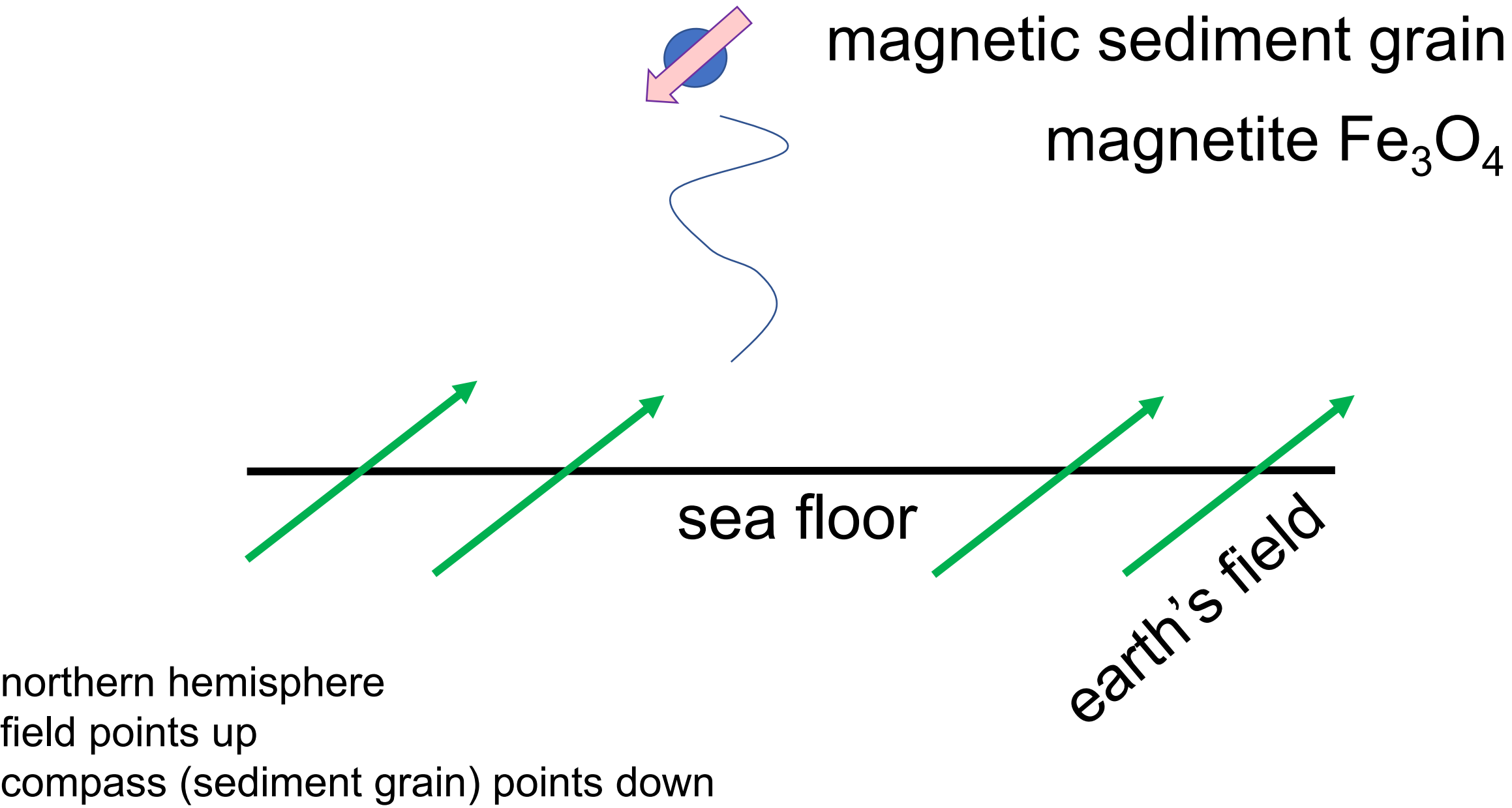
compass points UP

turn formula around

$$\tan(\text{latitude}) = \frac{1}{2} \tan \theta$$

measure inclination
infer latitude





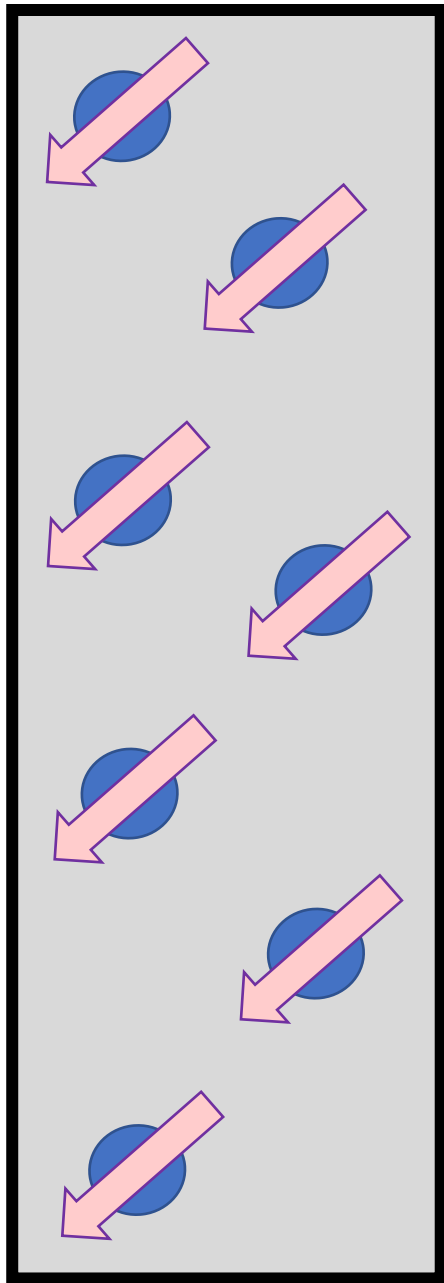
magnetic sediment grain

magnetite Fe₃O₄

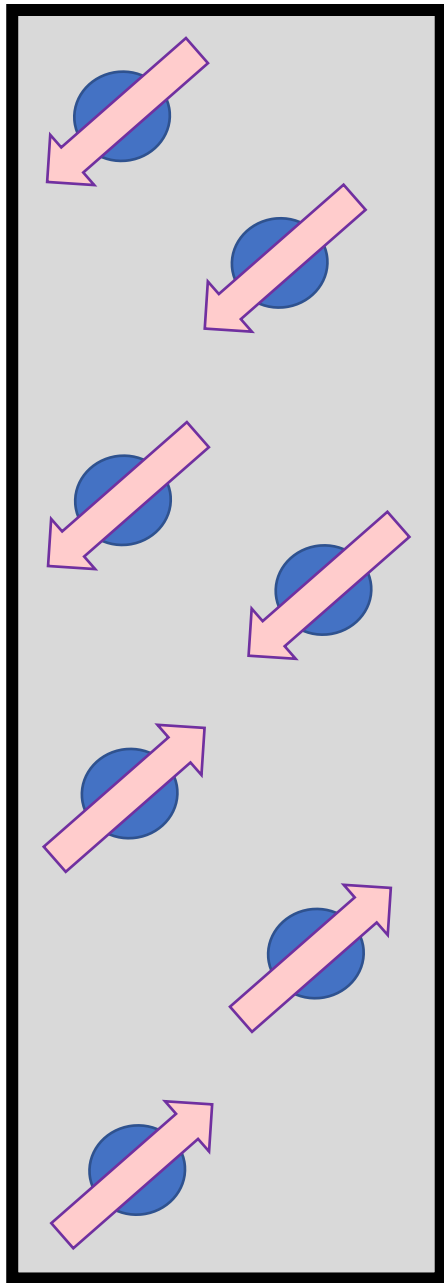
sea floor

earth's field

northern hemisphere
field points up
compass (sediment grain) points down

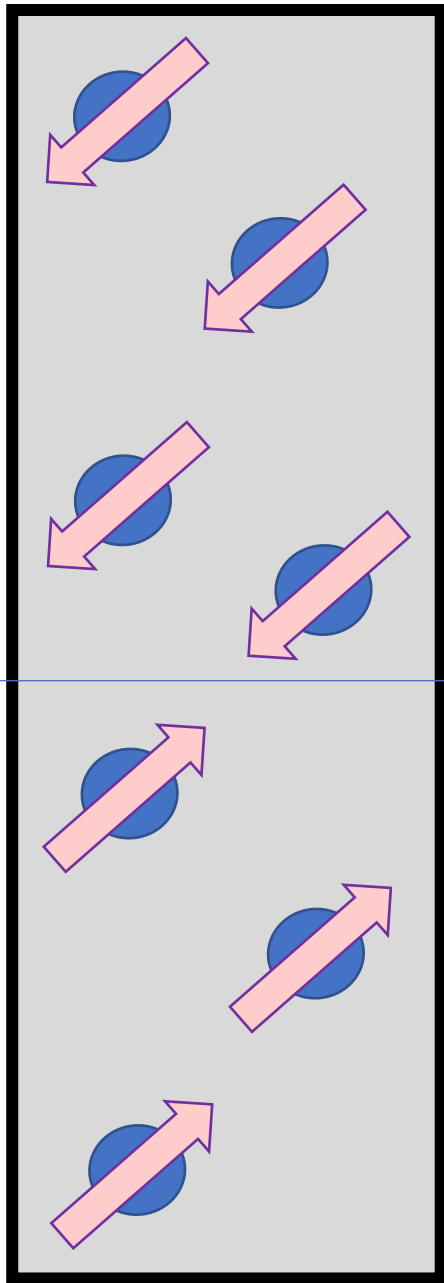


sediment core



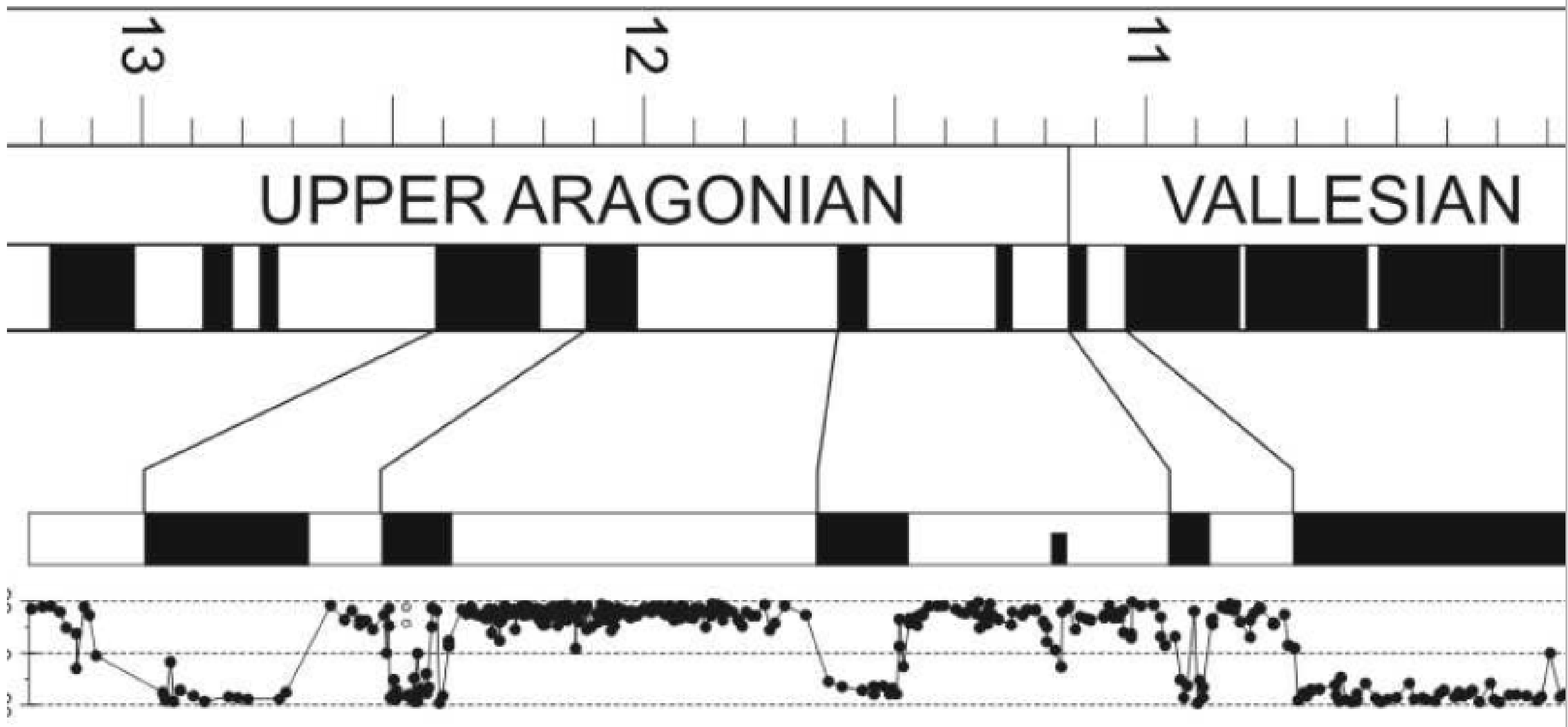
what's happened here?

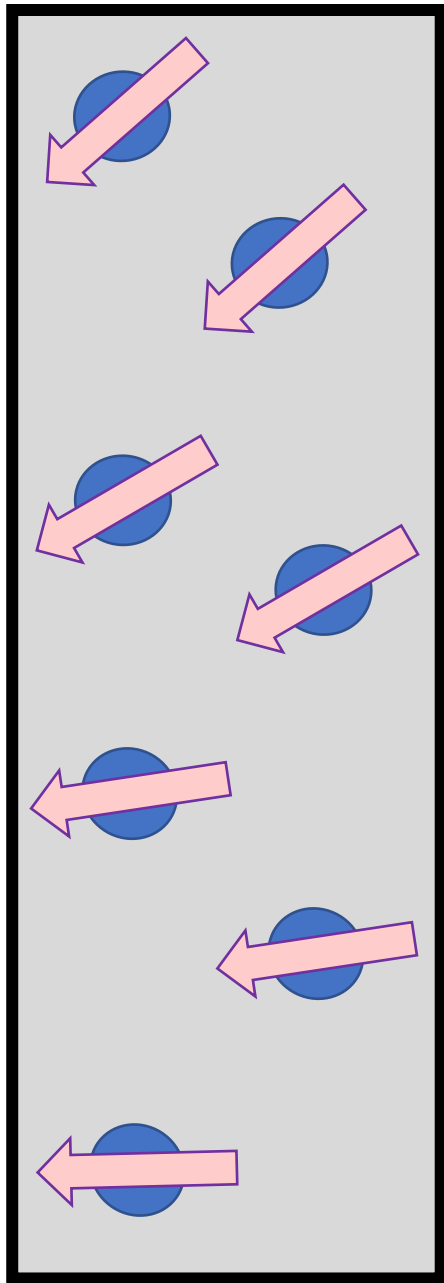
sediment core



reversal

sediment core

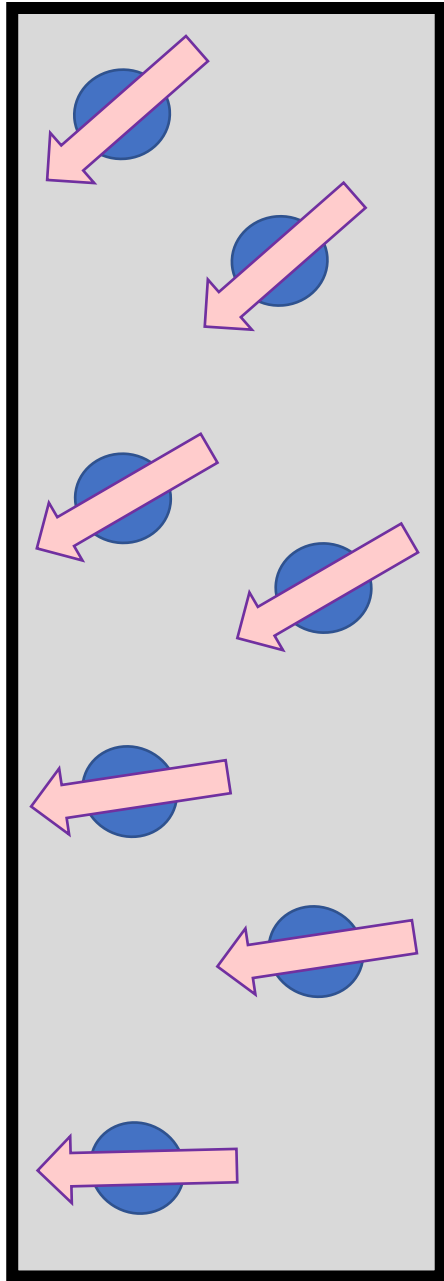




sediment core

what's happened here?

plate is moving away from the equator



sediment core

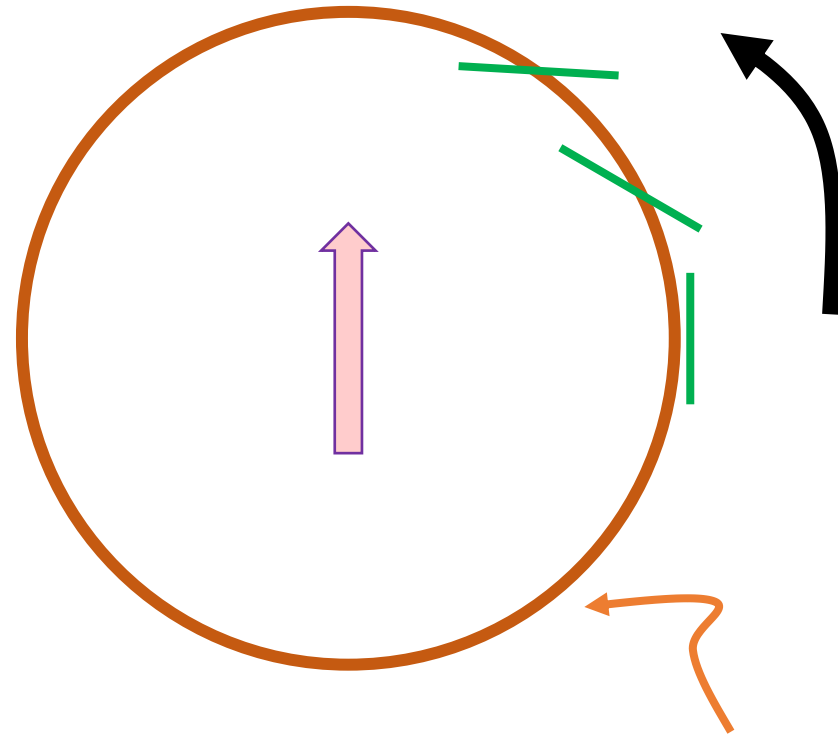


plate motion

earth

Site at Oregon – California Border

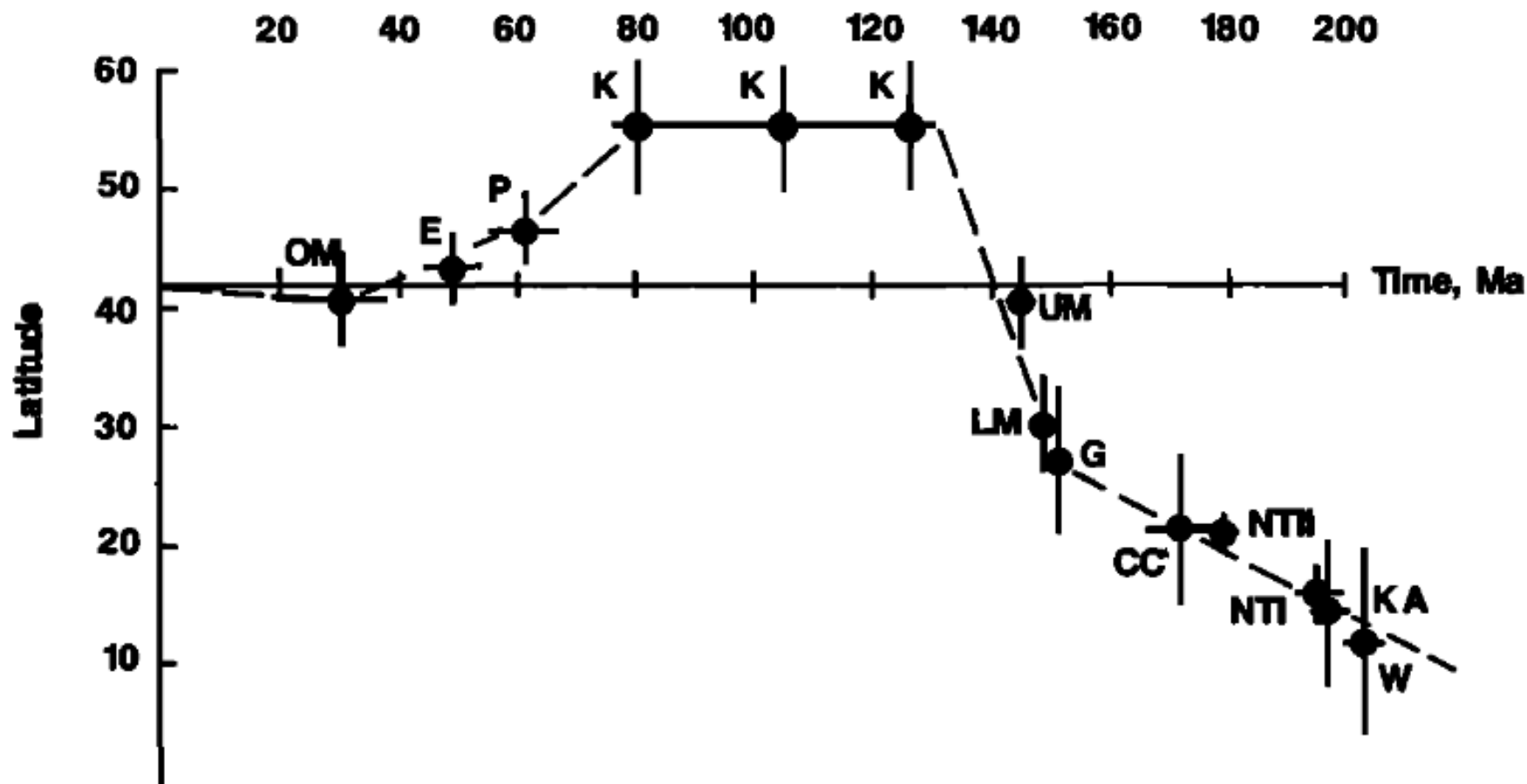
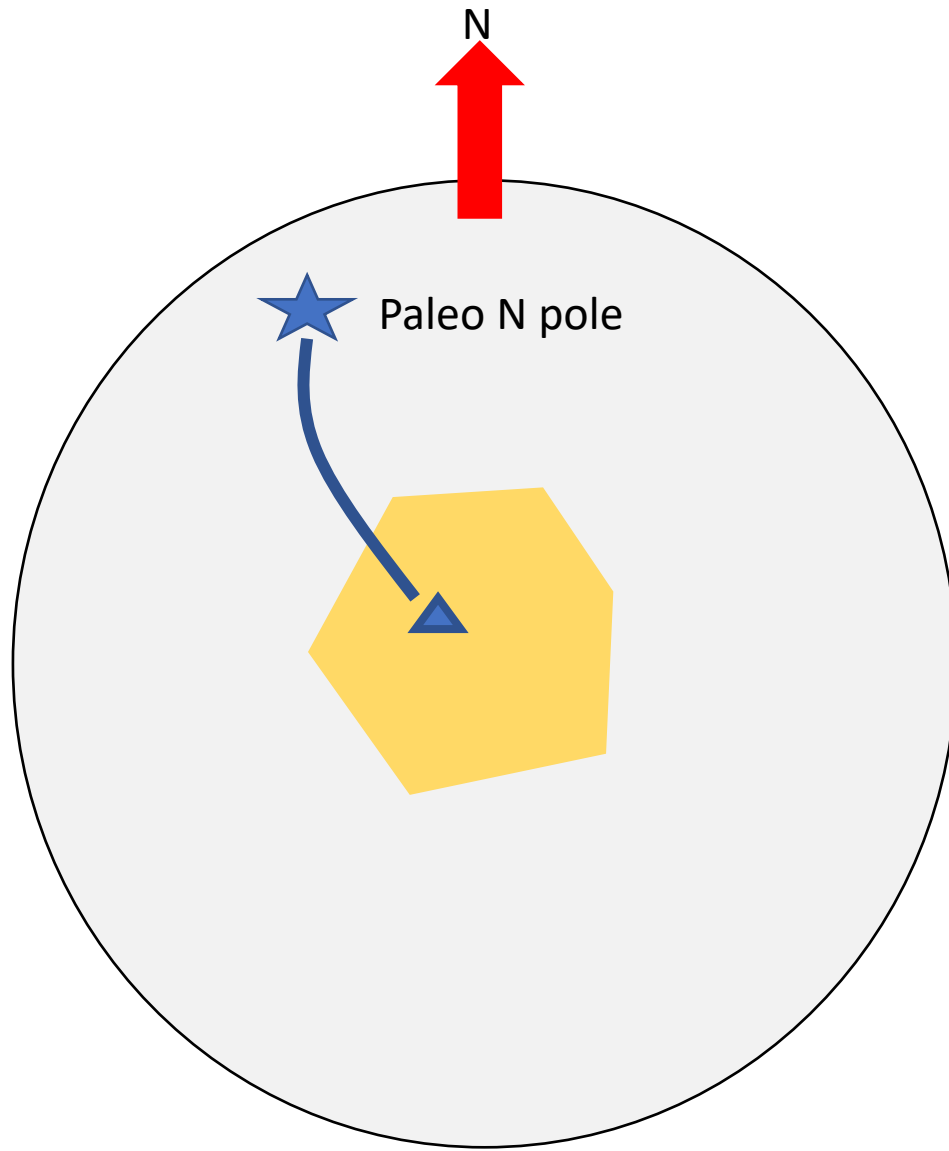
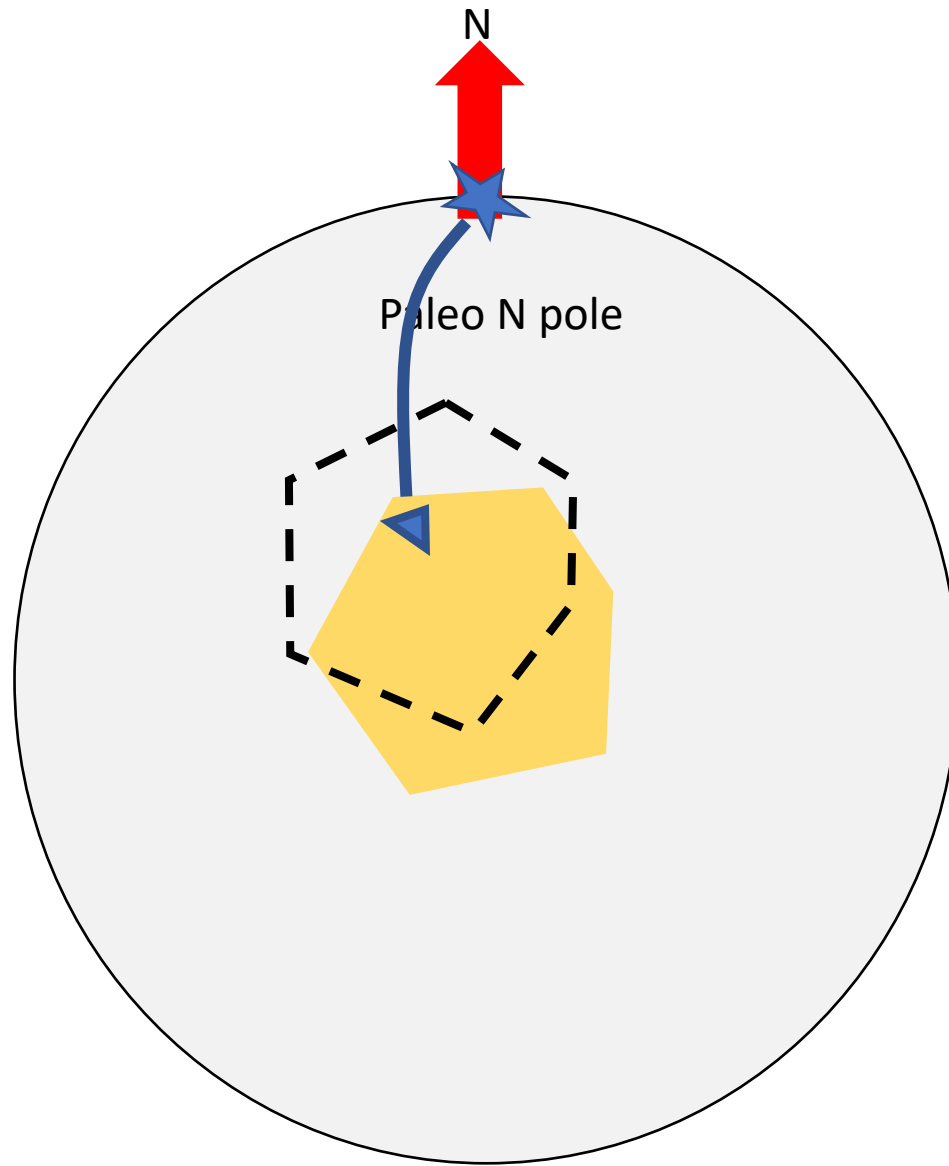


Fig. 2. Paleolatitude of "observation site" (42°N, 124°W) as a function of time, calculated from reference poles shown in Figure 1. Note the rapid increase in paleolatitude during the Late Jurassic-Early Cretaceous. Symbols as in Figure 1.



200 my ago
According to a site on the plate
The pole was here



200 my ago
According to a site on the plate
The pole was here

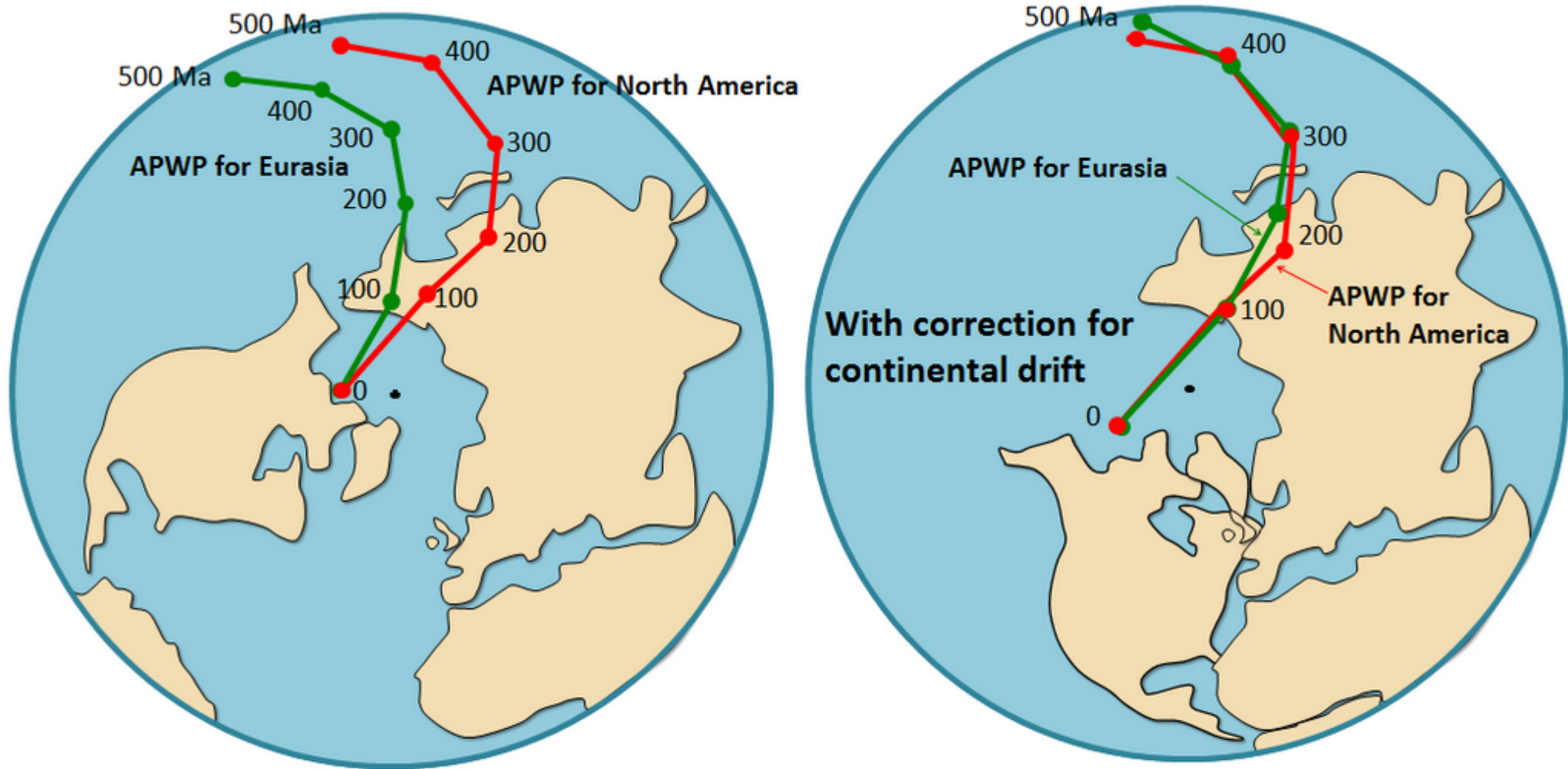


Figure 4.2.2 Polar wandering curves. Curves from Eurasia and North America seem to show that the north magnetic pole was located in two places simultaneously throughout history (left). However, if the continents are rearranged into Pangaea, the two curves overlap, showing that it is the continents that have moved, not the pole (right) (Steven Earle, “Physical Geology”).

Dynamo

(a) multipole expansion

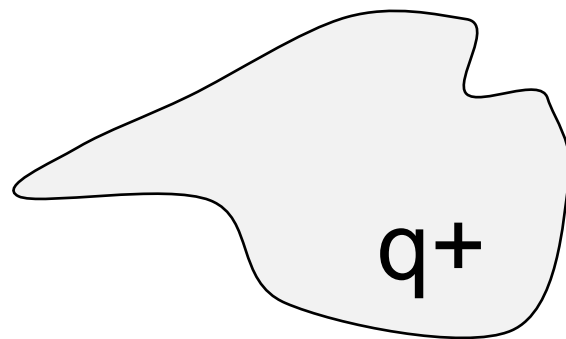
(b) frozen flux approximation

(c) magnetic diffusion

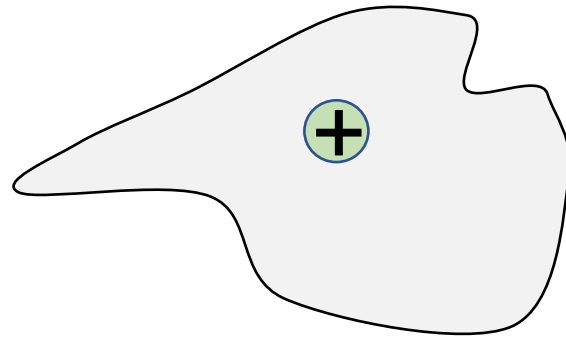
Dynamo

(a) multipole expansion

field of complicated objects

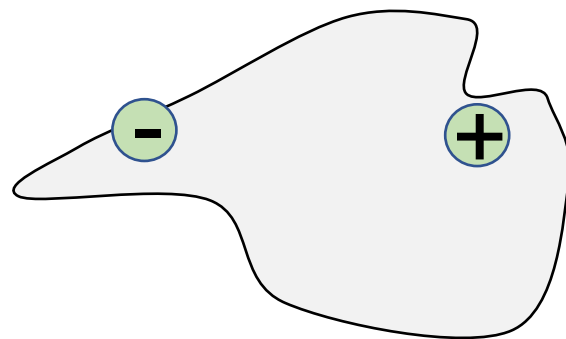


field of complicated objects, made by summing



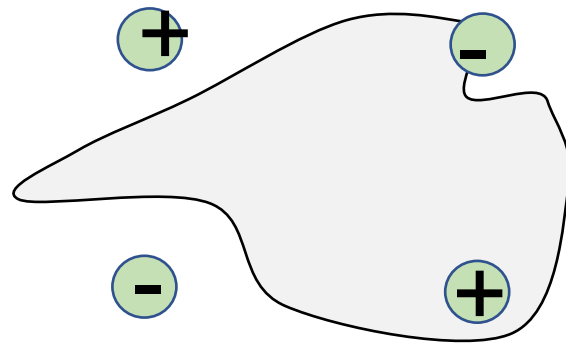
monopole
(point charge)

plus



dipole

plus



quadrupole

monopole

$\oplus q$

r

$$\frac{q}{-r^2}$$

dipole

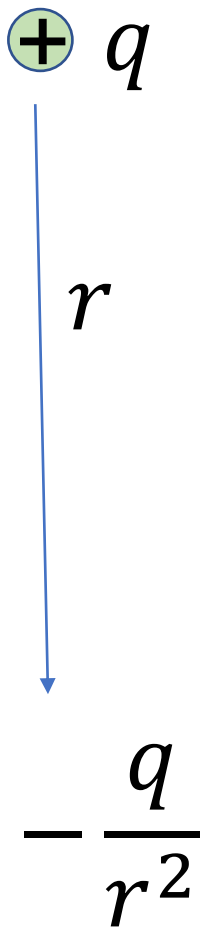
$\ominus -q$

$\oplus +q$

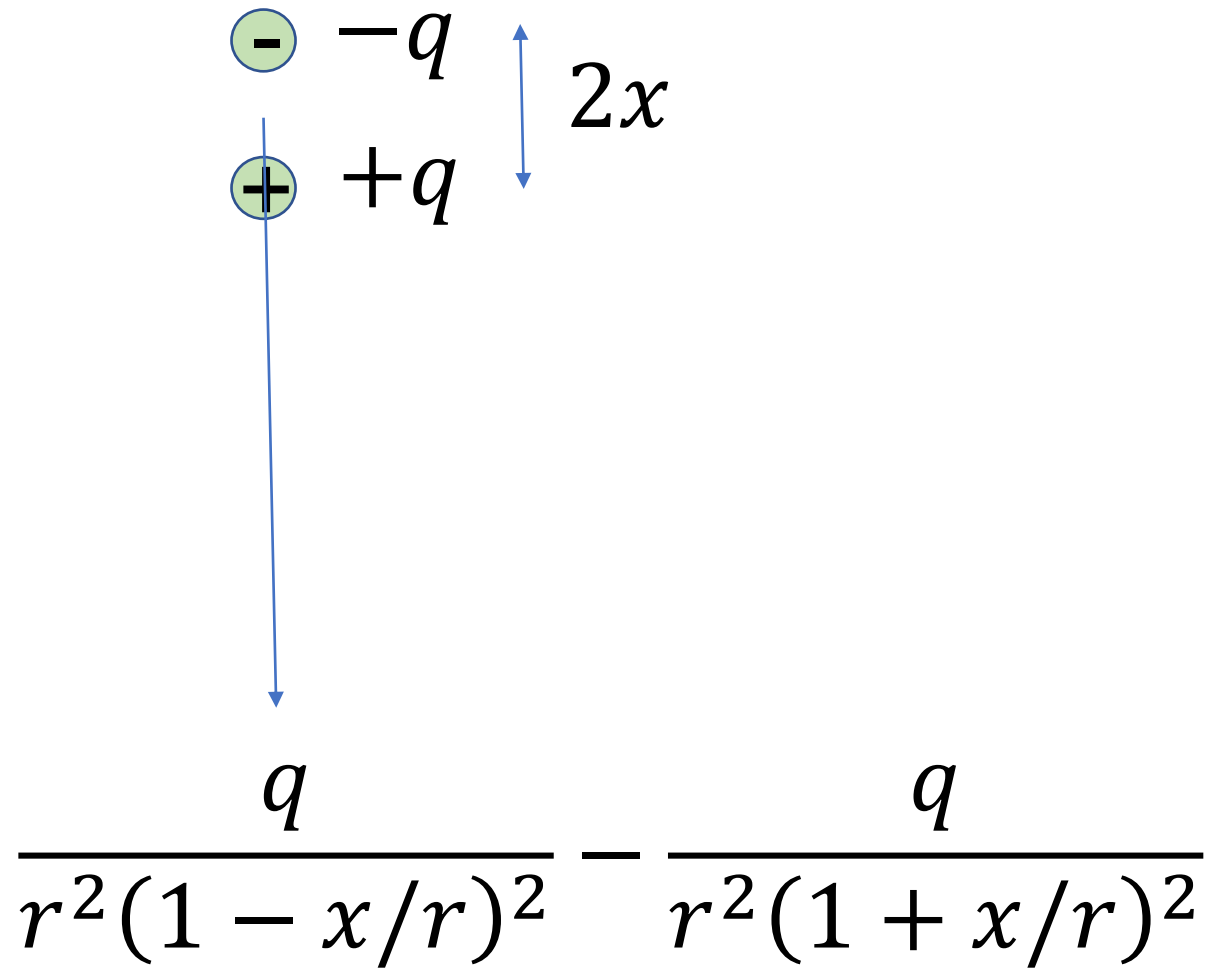
$2x$

$$\frac{q}{(r-x)^2} - \frac{q}{(r+x)^2}$$

monopole



dipole



monopole

$\oplus q$

r

$$\frac{q}{r^2}$$

dipole

$\ominus -q$

$\oplus +q$

$2x$

$$\frac{q(1 - 2x/r)}{r^2} - \frac{q(1 + 2x/r)}{r^2}$$

monopole

$\oplus q$

r

$$\frac{q}{r^2}$$

dipole

$\ominus -q$

$\oplus +q$

$2x$

$$\frac{q(2x/r)}{r^2} \quad \frac{q(2x/r)}{r^2}$$

monopole

$\oplus q$

r

$$\frac{q}{r^2}$$

dipole

$\ominus -q$

$\oplus +q$

$2x$

$$\frac{4qx}{r^3}$$

monopole

$\oplus q$

r

$$\frac{q}{r^2}$$

dipole

$\ominus -q$

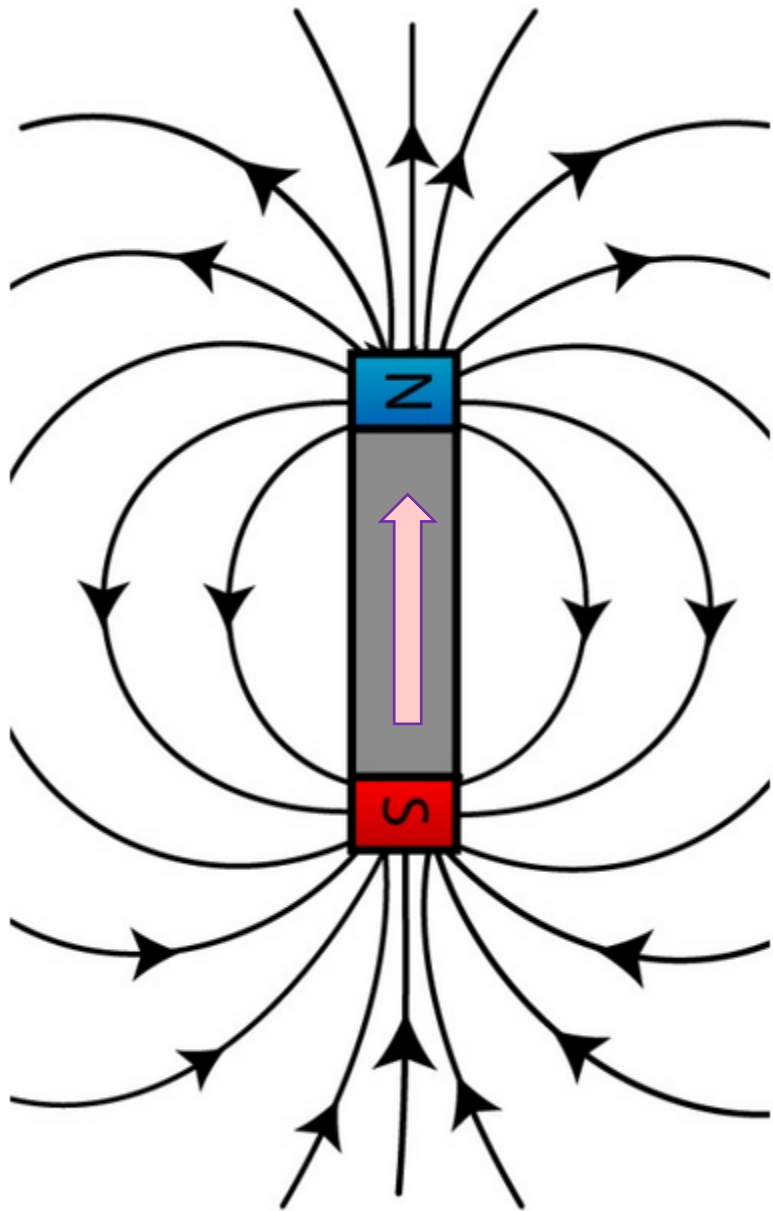
$\oplus +q$

$2x$

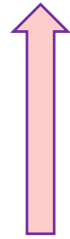
$$\frac{4qx}{r^3}$$

dipole moment m_2

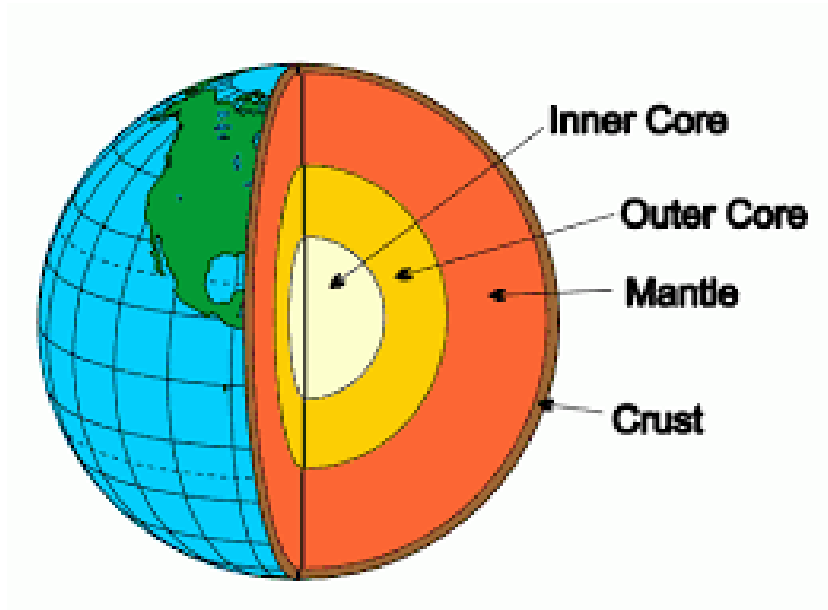
faster falloff



magnetization vector \mathbf{M}



is just the overall dipole
moment of the object



Lesson

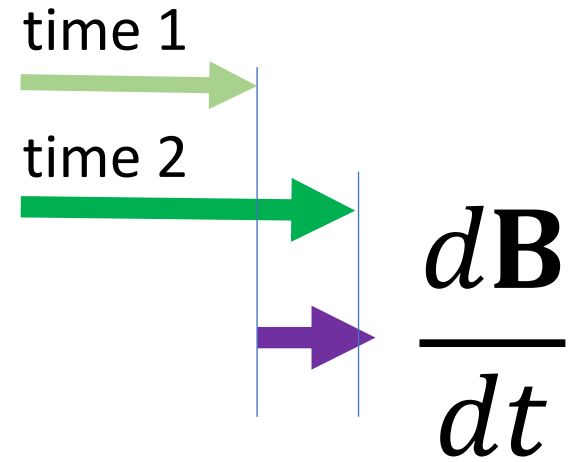
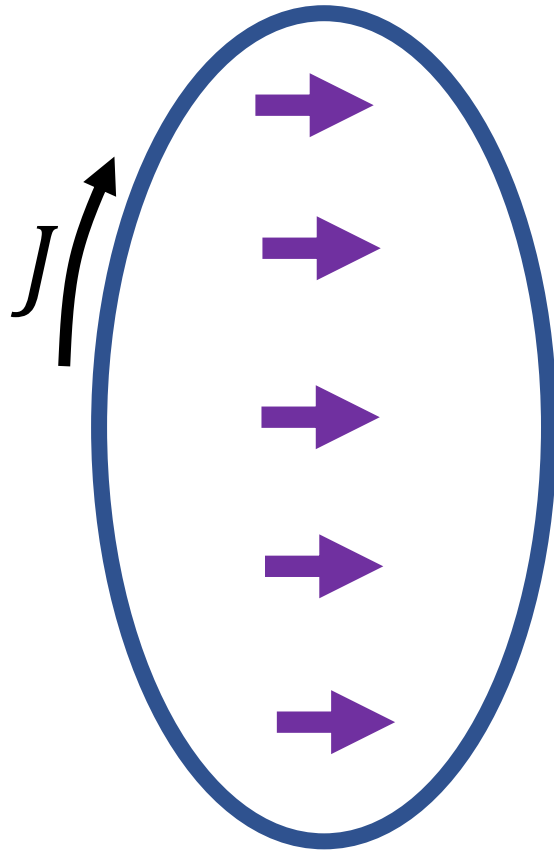
field does not have to be perfect dipole in core

to be reasonably dipolar at the surface of the Earth

Dynamo

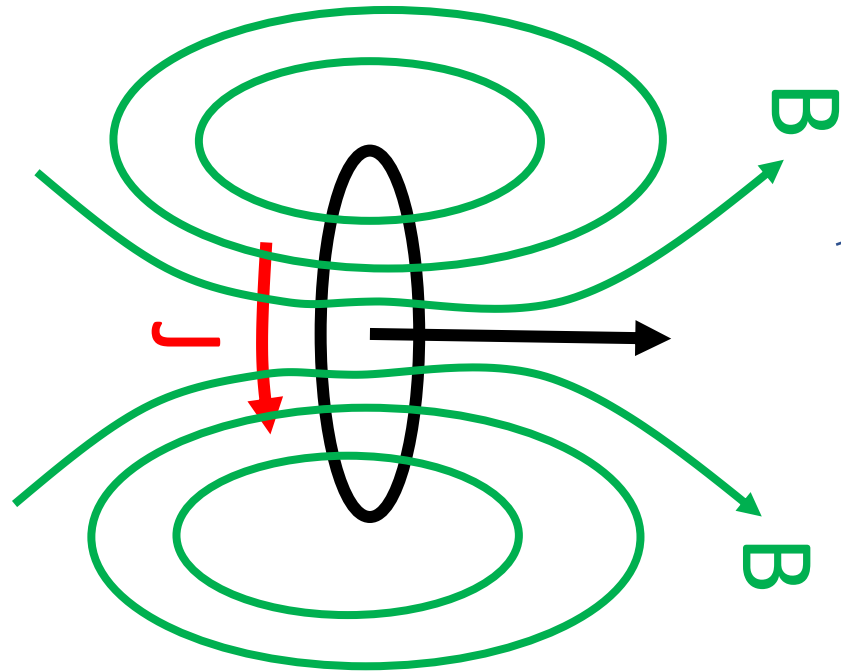
(b) frozen flux approximation

Generator principle



$$J \propto -\frac{d\mathbf{B}}{dt}$$

Solenoid principle

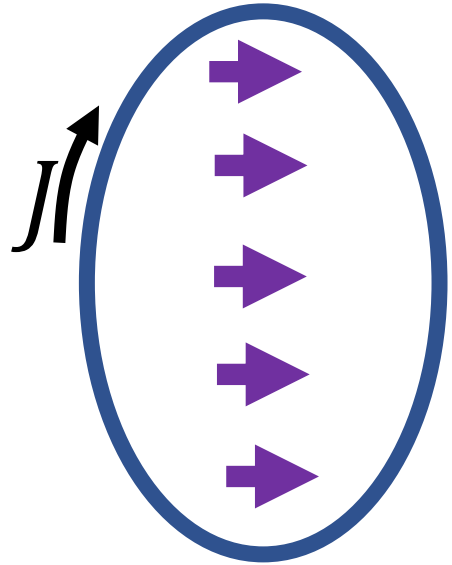


inside loop

$$\mathbf{B} \propto \mu_0 J$$

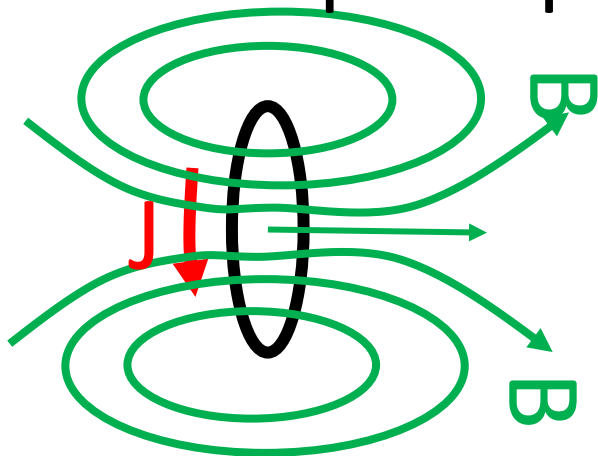


Generator principle

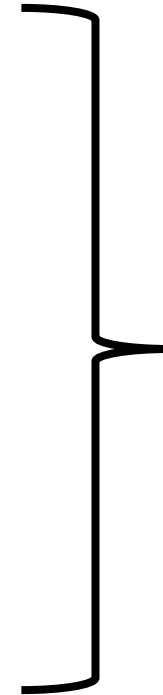


$$J \propto -\frac{d\mathbf{B}}{dt}$$

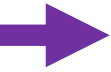
Solenoid principle

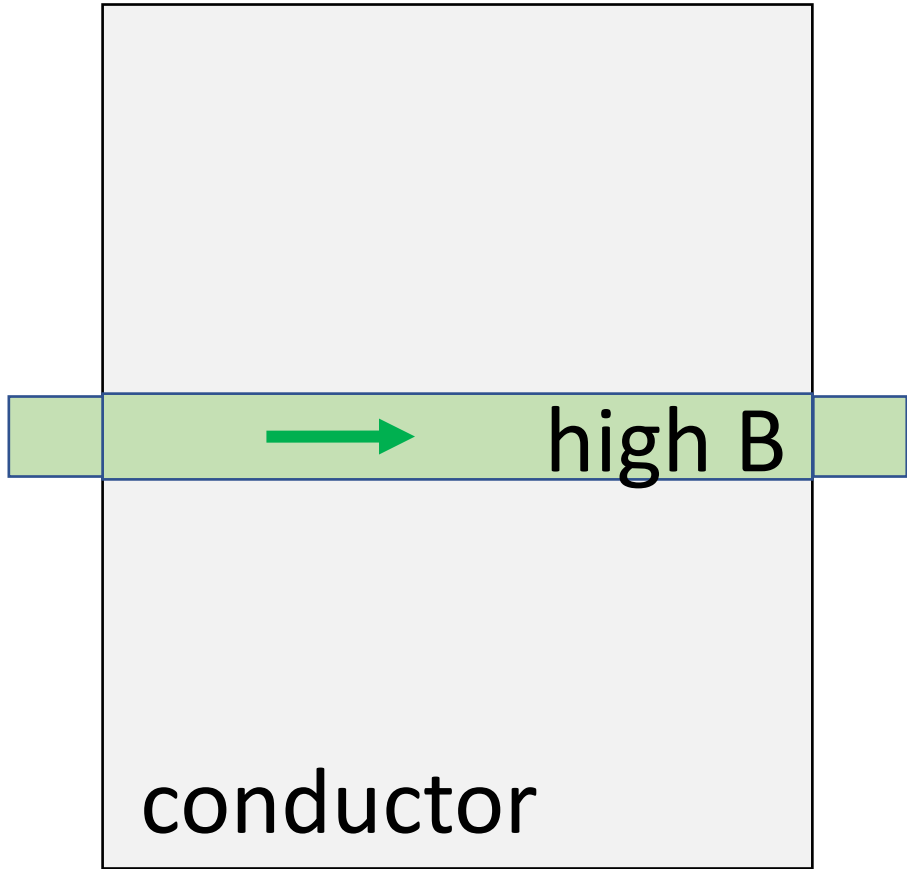


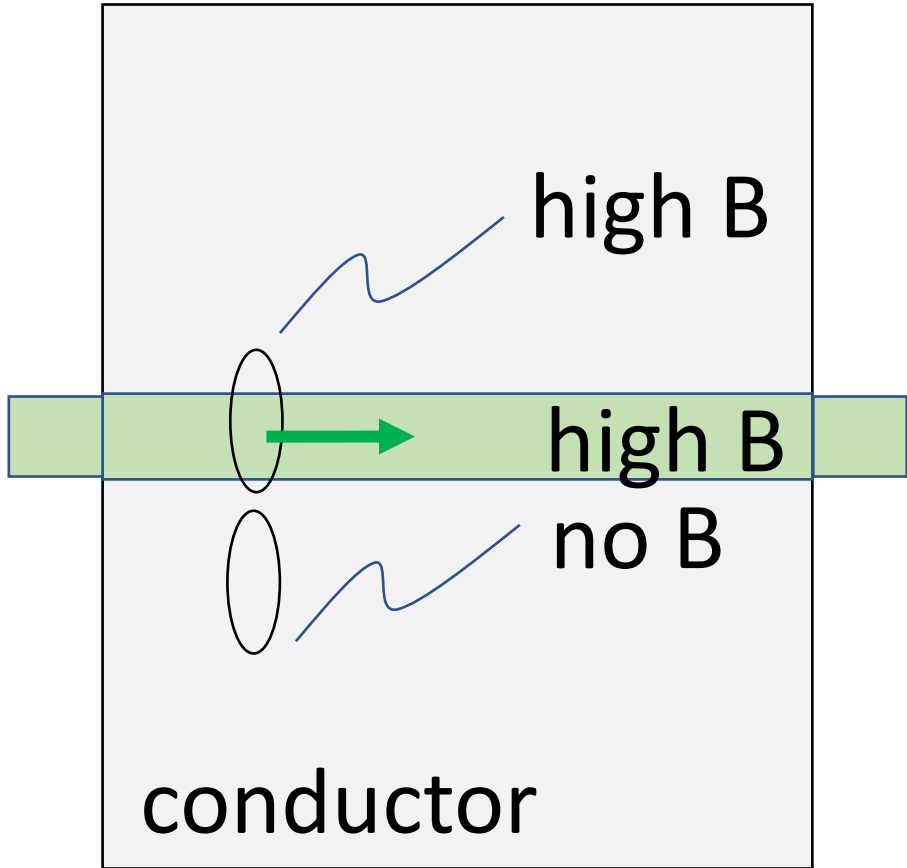
$$\mathbf{B} \propto \mu_0 J$$



$$\mathbf{B} \propto -\mu_0 \frac{d\mathbf{B}}{dt}$$



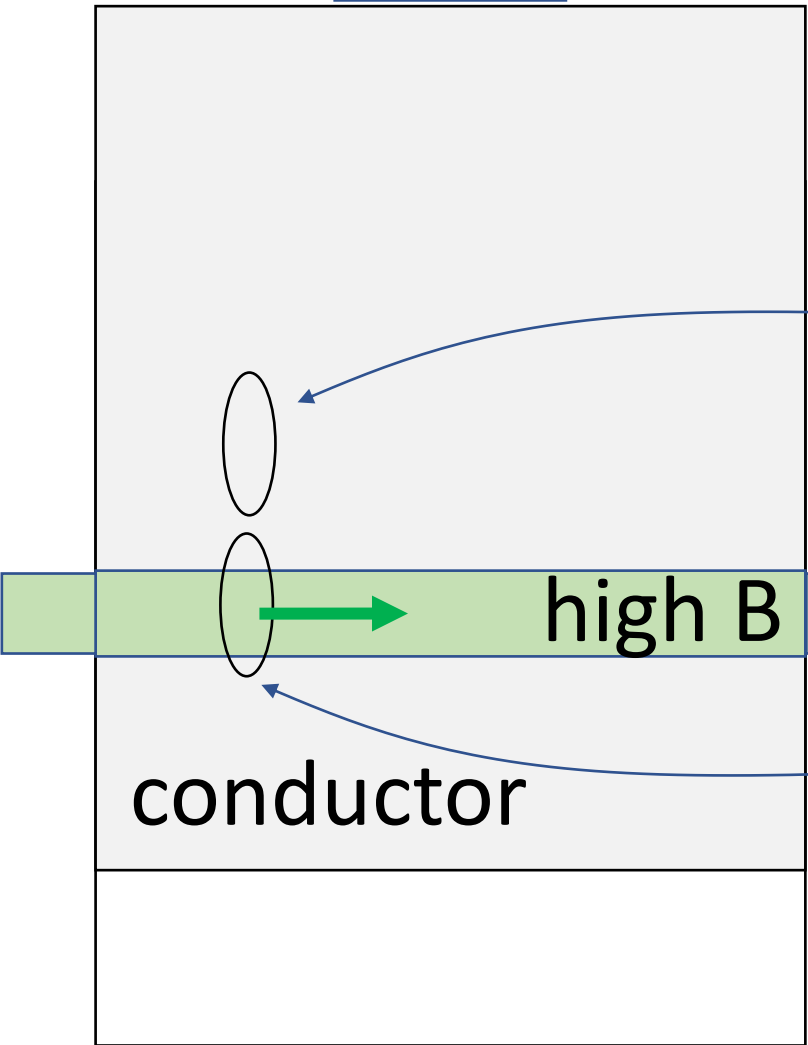




imaginary loops



move conductor

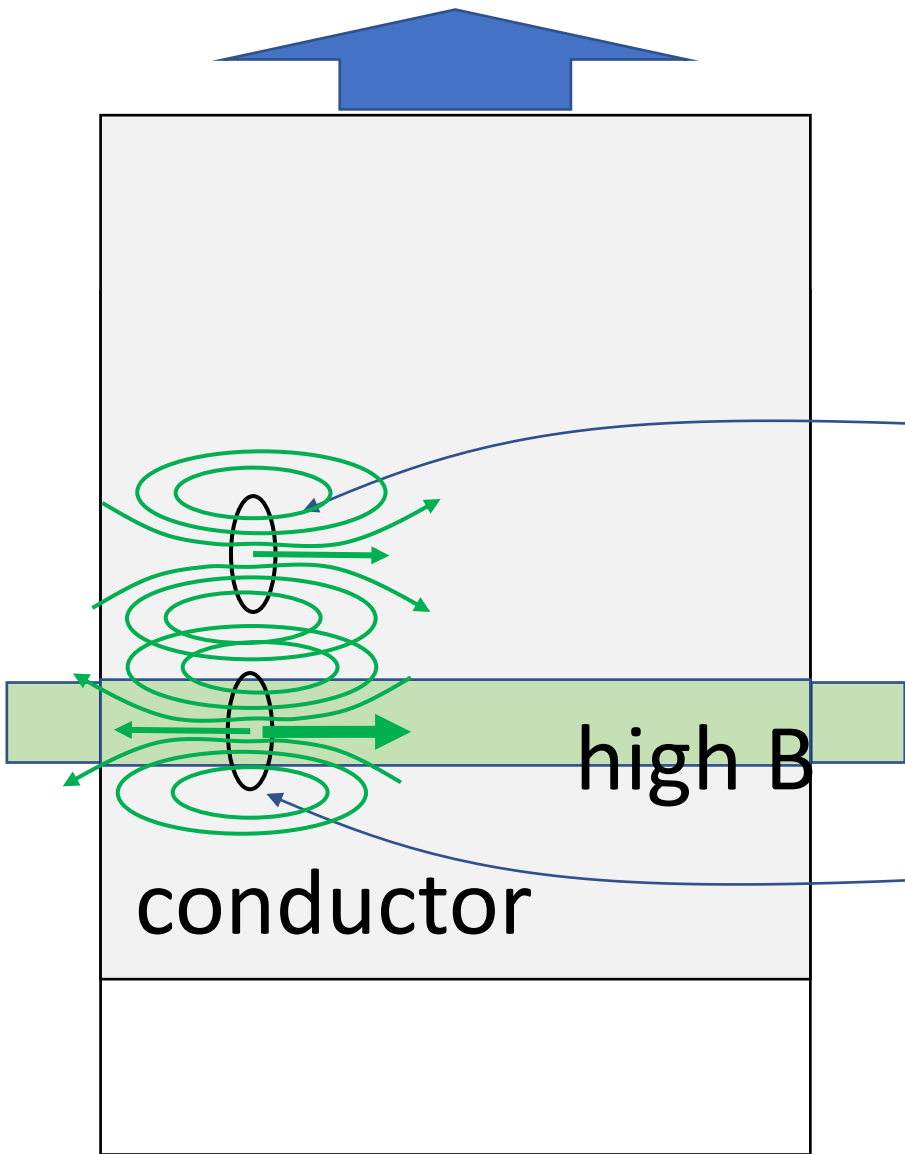


B went down

$$\frac{d\mathbf{B}}{dt} \quad -$$

B went up

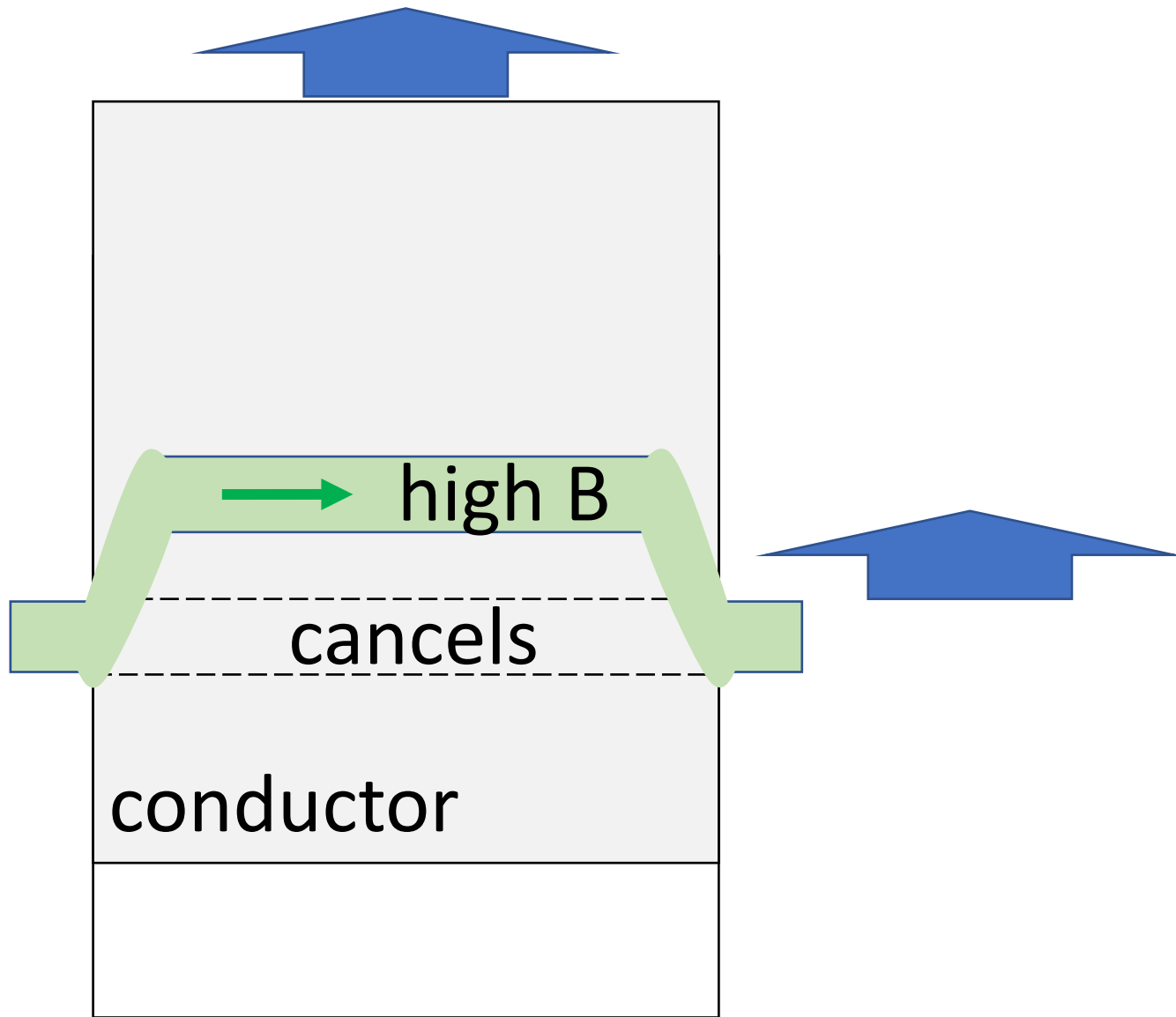
$$\frac{d\mathbf{B}}{dt} \quad +$$



$$\mathbf{B} \propto -\mu_0 \frac{d\mathbf{B}}{dt}$$

B+ because $\frac{d\mathbf{B}}{dt} -$

B- because $\frac{d\mathbf{B}}{dt} +$



net results
magnetic field
acts like it is "frozen in"

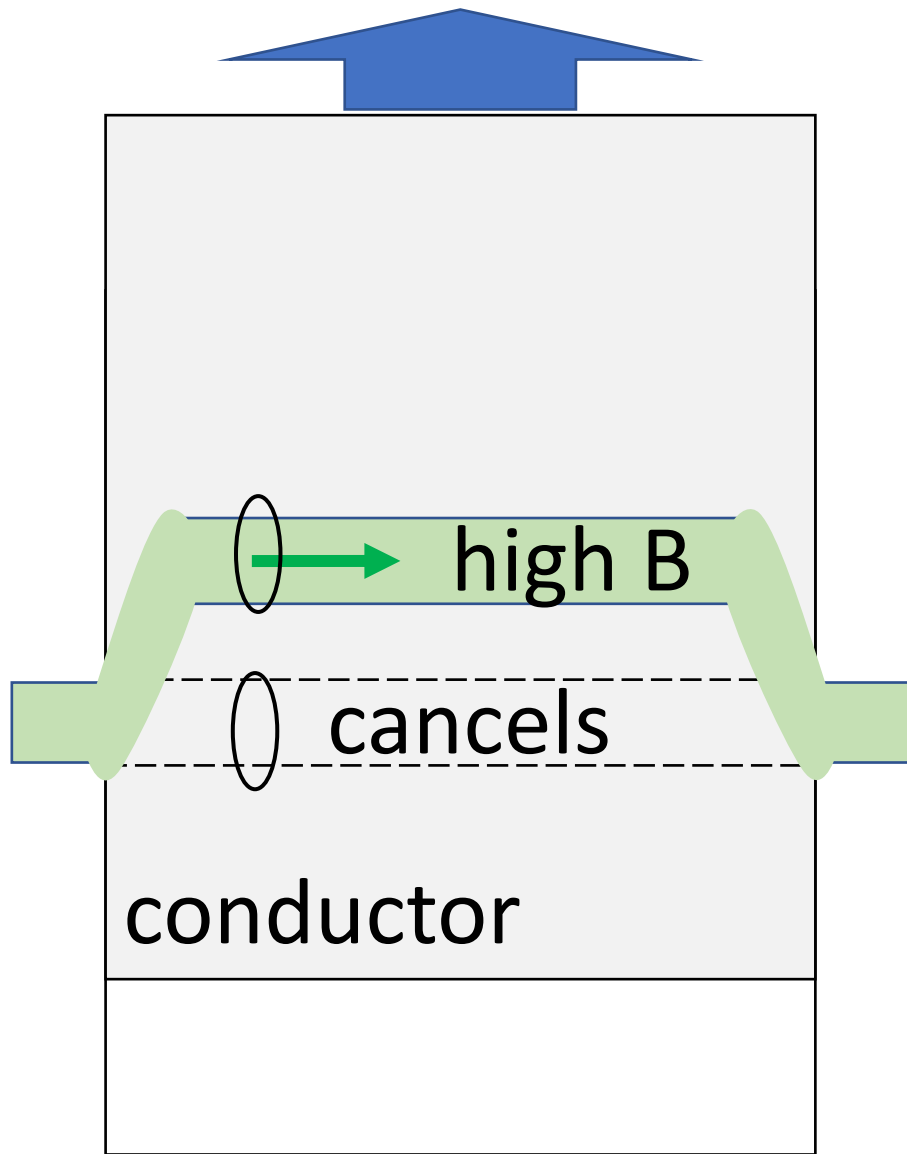
Lesson

Convection drags the magnetic field with it

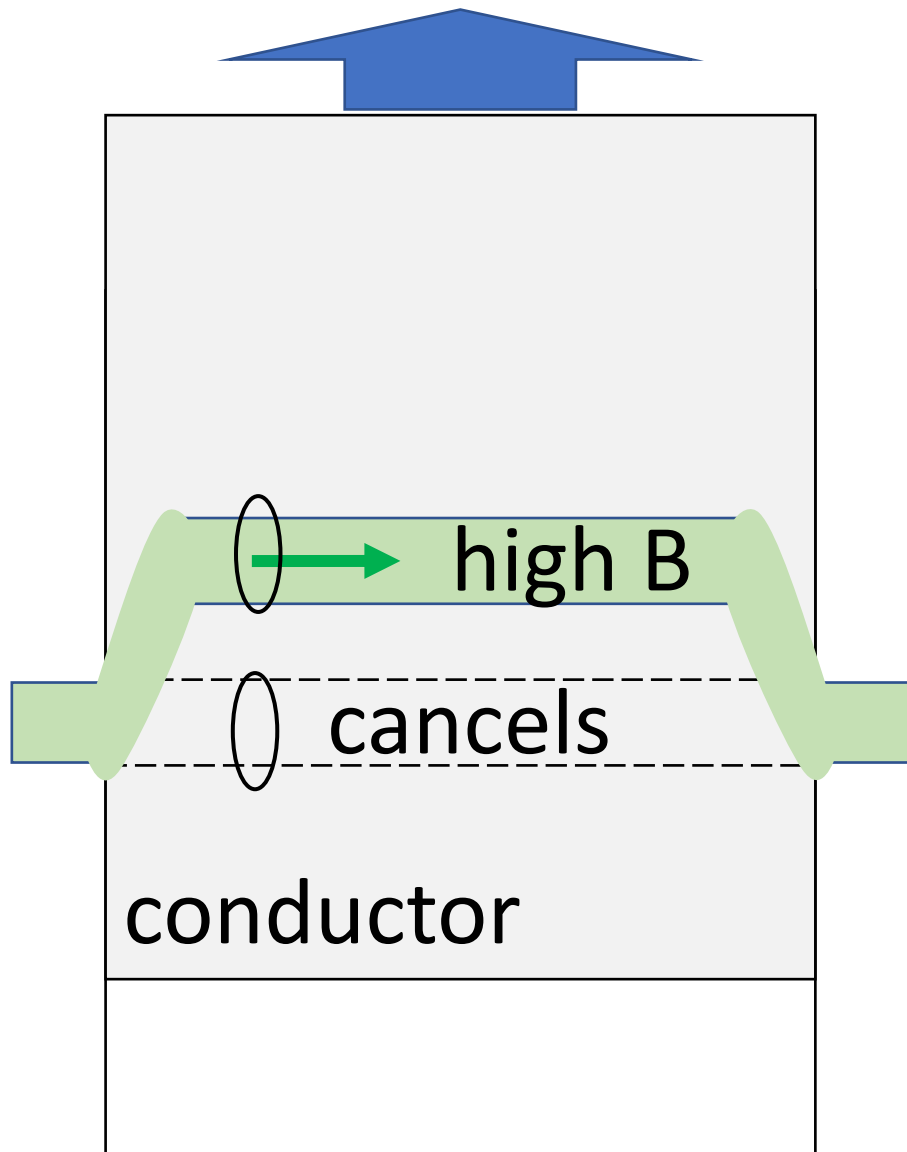
creates a field that changes with time

Dynamo

(c) magnetic diffusion

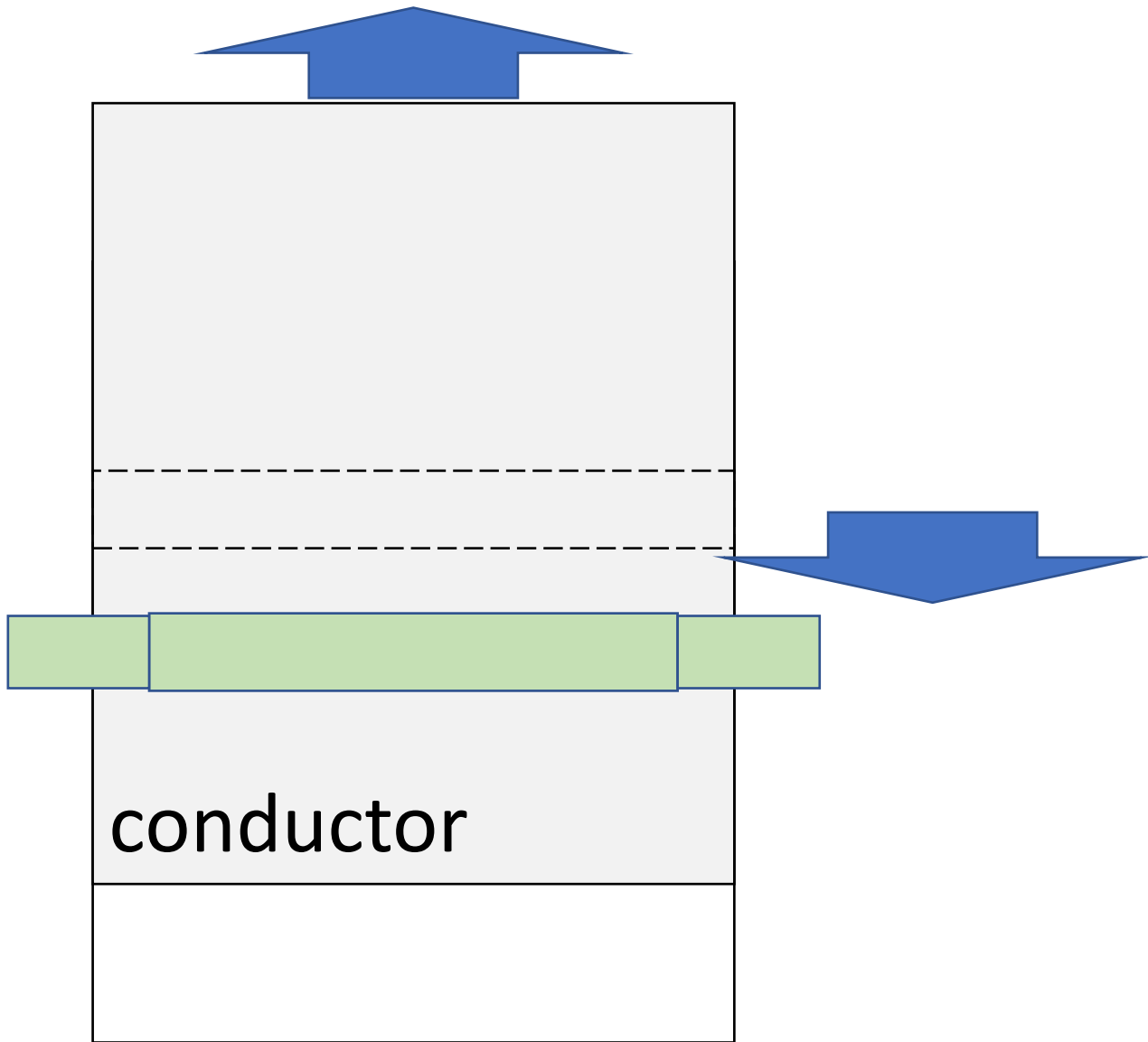


shape being sustained
by currents in the loops

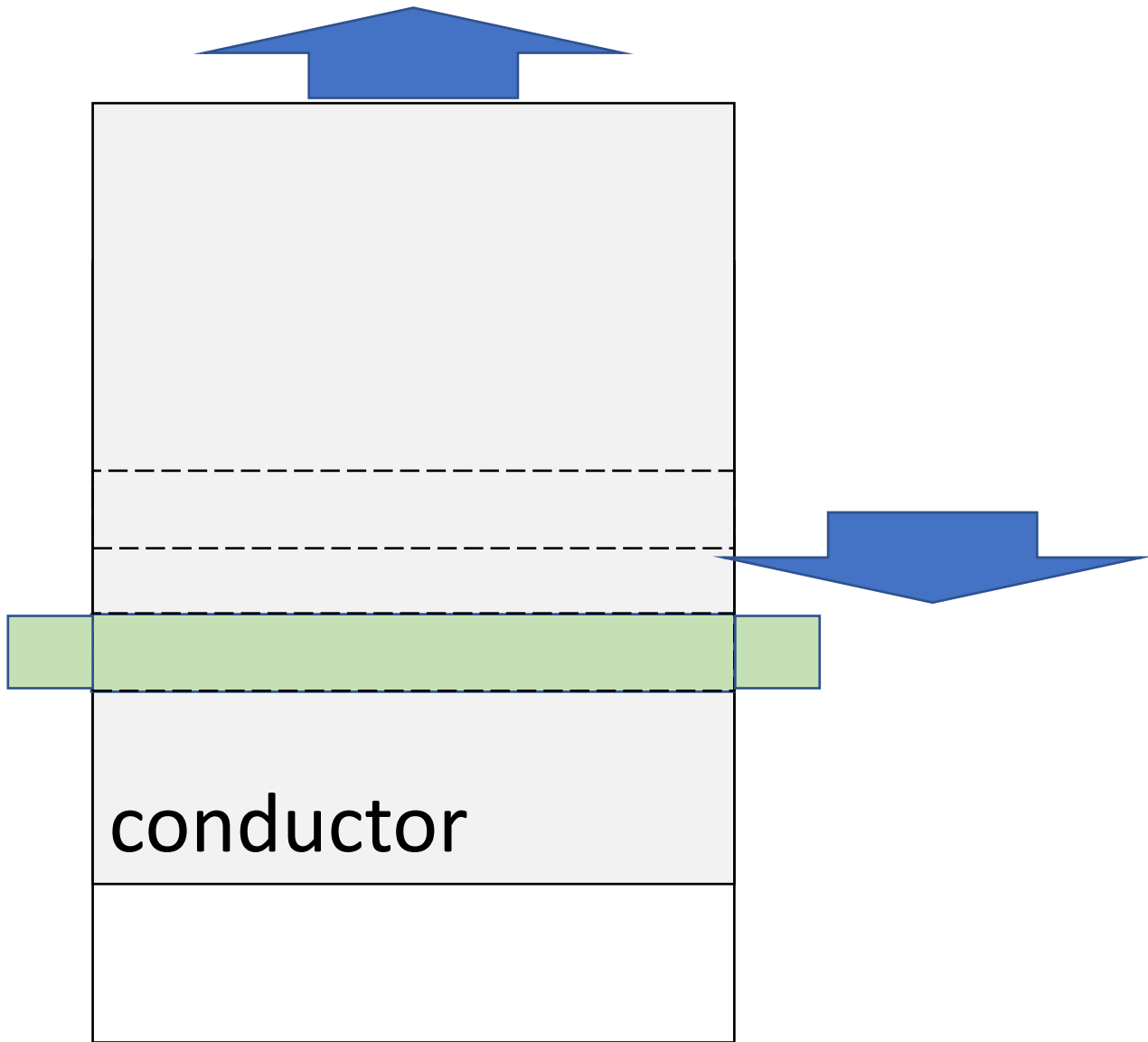


if not perfect conductor
loops lose energy

acts like heat flow with
a diffusivity of $\kappa = \frac{1}{\sigma\mu_0}$



Important question
what is the time scale
of diffusion



Temperature solution

$$\operatorname{erf}\left(\frac{z}{\sqrt{4\kappa t}}\right)$$

$$\frac{z}{\sqrt{4\kappa t}} = 1$$

$$t = \frac{z^2}{4\kappa} = \frac{z^2}{4\sigma\mu_0}$$

$z =$ length scale

$$\sigma = 10^5 \text{ S/m}$$

$$z = 400 \text{ km} \quad t = 10,000 \text{ years}$$

Lesson

field not completely “frozen in”

so field in core cannot become
absurdly complicated