

Oldenburg, D.W., One dimensional  
 invasion of natural source magneto-  
 telluric observations, Geophysics 44,  
 1218-1244, 1979.

conductivity  
 note resistivity  $\rho(z) = \frac{1}{\sigma(z)}$

$$\frac{d^2 E}{dz^2} + i\omega\mu_0\sigma(z)E = 0$$

$$R = \frac{B}{E} \quad B = \frac{-1}{i\omega} \frac{dE}{dz}$$

$$\frac{dR}{dz} - i\omega R^2 - \mu_0\sigma(z) = 0$$

homogeneous solution  $\sigma(z) = \sigma_0$

$$E = E_0 e^{(i-1)\xi z} \quad \xi = \left[ \frac{1}{2} \mu_0 \sigma_0 \omega \right]^{\frac{1}{2}}$$

$$\frac{dE}{dz} = E_0 (i-1)\xi e^{(i-1)\xi z}$$

$$R = \frac{-1}{i\omega} \frac{dE/dz}{E} = \frac{-1}{i\omega} \frac{(i-1)\xi}{1} \xi$$

$i-1 = \sqrt{2} e^{i\pi/4}$   
 $(i-1)^2 = -1 - 2i + 1 = -2i$   
 $\left(\frac{\sqrt{2}}{2} e^{i\pi/4}\right)^2 = -2i$   
 $\xi = \sqrt{-2i} \left[ \frac{1}{2} \mu_0 \sigma_0 \omega \right]^{\frac{1}{2}}$   
 $R = \frac{-1}{i\omega} \frac{(i-1)\xi}{1} \xi$

$$-\left(\frac{1-i}{i}\right) = -\left(\frac{-ii-i}{i}\right) = -(-i-1) = i+1$$

$$R = \frac{i+1}{\omega} \sqrt{\frac{\mu_0 \sigma_0}{2\omega}}$$

$$(i+1)^2 = -1 + 2i + 1 = 2i$$

$$R^2 = 2i \frac{\mu_0 \sigma_0}{2\omega} = \frac{i\mu_0 \sigma_0}{\omega}$$

$$\mathcal{L}\{R\} + i\omega R^2 = \frac{i\mu_0 \sigma_0}{\omega}$$

$$R^2 = -\mu_0 \sigma_0$$

$$\frac{dR}{dz} = i\omega R^2 + \mu_0 \sigma_0 = 0$$

so given a  $\sigma(z)$  over a homogenous halfspace, we can integrate up

$$\text{with } \frac{dR}{dz} = i\omega R^2 + \mu\sigma$$

with  $R = (i+1) \sqrt{\frac{\mu_0 \sigma_0}{2\omega}}$  at ~~the~~ halfspace interface



Numerical Methods for Scientists  
and Engineers, 2nd Edition,  
RW Hamming, McGraw Hill,  
1973, 721 pp.

p 413 Runge-Kutta Methods, Sec 24.2

$$y' = f(x, y)$$

$$y(x_n) = y_n$$

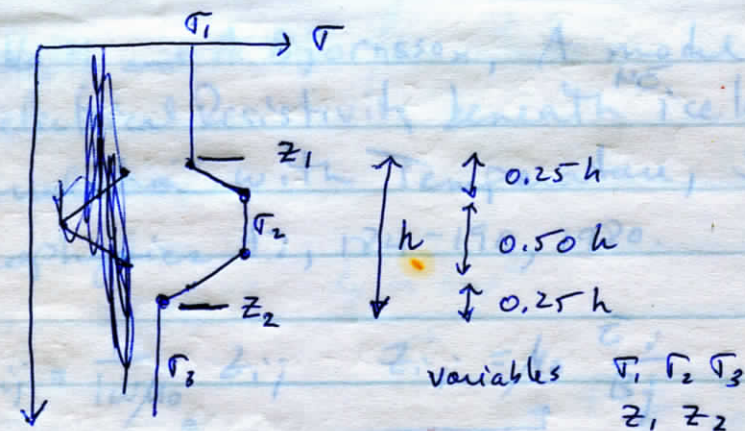
$$k_1 = hf(x_n, y_n)$$

$$k_2 = hf\left(x_n + \frac{h}{2}, y_n + \frac{k_1}{2}\right)$$

$$k_3 = hf\left(x_n + \frac{h}{2}, y_n + k_2\right)$$

$$k_4 = hf(x_n + h, y_n + k_3)$$

$$y_{n+1} = y_n + \frac{1}{6}(k_1 + 2k_2 + 2k_3 + k_4)$$



$$h = z_2 - z_1$$

$$z < z_1 \quad \sigma = \sigma_1$$

$$z > z_1 \text{ and } z < z_1 + 0.25h$$

$$\sigma = \sigma_1 + \frac{\sigma_2 - \sigma_1}{0.25h} (z - z_1)$$

$$z > z_1 + 0.25h \text{ and } z < z_1 + 0.75h$$

$$\sigma = \sigma_2$$

$$z > z_1 + 0.75h \text{ and } z < z_2$$

$$\sigma = \sigma_2 + \frac{\sigma_3 - \sigma_2}{0.25h} (z - (z_1 + 0.75h))$$

$$z > z_2 \quad \sigma = \sigma_3$$



Bello, M and A Bjornsson, A model of electrical Resistivity beneath <sup>NE</sup> Iceland, Correlation with Temperature, J. Geophysics 47, 184-190, 1980.

$$C_{ij} = \frac{1}{i\omega\mu_0} z_{ij} \quad z_{ij} = \mu_0 \frac{\sigma_j}{B_j}$$

B in nT

E in mV/km

problem here!

URD

43	8
11	12
25	

resistivity depth  
Ωm, km

period range  
20 - 2000 s

$$\rho_a = \frac{\mu_0 T}{2\pi} |z_{11}|^2 \quad \phi_{11} = \arg(z_{11})$$

$\rho = \Omega m \quad \phi \text{ in deg.}$

This must be wrong.

units MKSA for Jackson p 238

B tesla

E volts/m

$\rho$   $\Omega m$

$\mu_0$   $4\pi \times 10^{-7} \frac{kg \cdot m \cdot amp^2}{s^2}$

$z =$

$$f = \frac{1}{\sqrt{\epsilon}} = \frac{\omega}{2\pi}$$

$$z = \mu_0 \frac{E}{B} = \mu_0 R^{-1}$$

$$\omega = \frac{2\pi}{\epsilon}$$

in homogeneous media

$$R^{-1} = \frac{1}{(i+1)} \sqrt{\frac{2\omega}{\mu_0 \sigma_0}}$$

$$\sigma_0 = \rho_0^{-1}$$

$$z = \frac{1}{i+1} \left( \frac{2\mu_0 \omega}{\sigma_0} \right)^{1/2} = \frac{1}{(i+1)} \left( 2\mu_0 \rho_0 \omega \right)^{1/2}$$

$$|z|^2 = \frac{1}{2} 2\mu_0 \rho_0 \omega$$

$$\rho = \mu_0 \frac{I}{2\pi} |z|^2 = \mu_0 \frac{1}{\omega} \left| \frac{1}{i+1} \right|^2 (2\mu_0 \rho_0 \omega)$$

$$= \mu_0 \frac{1}{\omega} \frac{1}{2} 2\mu_0 \rho_0 \omega$$

$$= \mu_0^2$$

This must be wrong.



handbook of Phys. Prop. of R<sub>1</sub> p 230

graphite resistivity

$$36 - 100 \times 10^{-8} \Omega m \quad \parallel \text{ cleavage}$$

$$28 - 9900 \times 10^{-8} \Omega m \quad \perp \text{ cleavage}$$

ohm's law

$$j = \sigma E$$

↑      ↑      ↘ electric field  
current density      conductivity

$$E = R J$$

↓      ↗      ↖  
EMF      resistance      current

Volts       $\Omega$       amp

$$E/m = \frac{\Omega}{m} m^2 \frac{J}{m^2}$$

$$E = (\Omega m) J$$

↓      ↘      ↖  
 $\frac{V}{m}$       amp       $\frac{m^2}{m^2}$

resistivity  $\Omega m$

conductivity  $\Omega^{-1} m^{-1}$

Beblo, M and A. Bjornsson, Magneto-  
telluric investigation of The Lower  
crust and upper mantle  
beneath Iceland, J. Geophysics 45,  
1-16, 1977.

note he define  $z = \frac{E}{B}$

$$S_a = \frac{\mu_0 \Gamma}{2\Omega} |z|^2 = \frac{\mu_0}{\omega} |z|^2$$

$$z = R^{-1} = \frac{1}{(1+i)} \sqrt{\frac{2\omega}{\mu_0 \Gamma^2}}$$

$$|z|^2 = \frac{1}{2} \frac{2\omega}{\mu_0 \Gamma^2} = \frac{\omega \mu_0}{\mu_0}$$

$$S_a = \frac{\mu_0}{\omega} \frac{\omega \mu_0}{\mu_0} = \mu_0$$

This works, so  
def of  $z$  in Beblo + Bjornsson  
(1977) must be off by a  
factor of  $\mu_0$ . Looks  
like he tried patch to make  
impedance match usual  
def. + blew  
it.



## Summary

$$I_a = \frac{\mu_0}{2\pi} \omega^{-1} |z|^{-2}$$

$$\phi = \arg(z)$$

$$z = R^{-1}$$

$$\frac{dR}{dz} = \omega R^2 + \frac{\mu_0}{\rho(z)}$$

$\rho$  = resistivity

$$R_0 = (i+1) \sqrt{\frac{\mu_0}{2\omega \rho_0}}$$



$$s = 3000 \text{ m}$$

skm depth  $\delta = \left( \frac{z}{\omega \mu \sigma} \right)^{1/2}$

$$= \left( \frac{2\rho}{\omega \mu} \right)^{1/2}$$

$$= \left( \frac{4\pi \rho}{T \cdot 4\pi \times 10^{-7}} \right)^{1/2}$$

$$= \left( \frac{\rho}{T} \right)^{1/2} 10^{7/2}$$

$$= \left( \frac{\rho}{T} \right)^{1/2} 1000 \sqrt{10}$$

$$\approx 3000 \left( \frac{\rho}{T} \right)^{1/2}$$

$$\rho = 20$$

$$T = 20$$

$$\delta = 3000 \text{ m}$$



derivation of Riccati eqn.

$$\frac{d}{dz} R = \text{Riccati eqn.}$$

$$R = \frac{B}{E} = -\frac{1}{i\omega} \frac{E_z}{E}$$

$$\frac{d}{dz} E_z + i\omega\mu\sigma E = 0$$

$$E_z = -i\omega ER$$

$$\frac{d}{dz} E_z = -i\omega ER_z - i\omega E_z R$$

$$-i\omega ER_z - i\omega E_z R + i\omega\mu\sigma E = 0$$

$$-i\omega R_z - i\omega \frac{E_z}{E} R + i\omega\mu\sigma = 0$$

$$R_z + \frac{E_z}{E} R - \mu\sigma = 0$$

$$R_z - i\omega R^2 - \mu\sigma = 0$$

$$R_z = i\omega R^2 + \mu\sigma$$

Oldenberg P 1228

$$P = 10 \text{ cm} \cdot 10$$

$$390 \quad 240$$

10

T    R  $\times 10^5$

ny mzf. f program

0.1	0.5359	0.53539
0.178	0.83065	0.83004
0.316	1.2829	1.28361
0.562	1.9563	1.95719
0.100	2.9283	2.92835
0.178	4.2979	4.29571
0.316	6.1806	6.18332
0.562	8.7539	8.75708
1.0	12.234	12.23440
1.78	16.916	16.90856
3.16	23.154	23.16348
5.62	31.517	31.52690
10	42.694	42.69418
17.8	57.625	57.59622
31.6	77.448	77.47714
56.2	103.96	103.99398
100	139.35	139.35499



$$E = e^{(i-1)\xi z}$$

$$\xi = \sqrt{\frac{\mu\sigma\omega}{2}}$$

$$\frac{dE}{dz} = (i-1)\xi e^{(i-1)\xi z}$$

$$\frac{d^2E}{dz^2} = (i-1)^2 \xi^2 e^{(i-1)\xi z}$$

$$\begin{aligned} (i-1)(i-1) &= \\ -1 - 2i + 1 &= -2i \end{aligned}$$

$$\frac{d^2E}{dz^2} + i\omega\mu\sigma E = 0$$

$$-2i \frac{\mu\sigma\omega}{2} e^{(i-1)\xi z} + i\omega\mu\sigma e^{(i-1)\xi z} = 0$$

$$(-i\omega\mu\sigma + i\omega\mu\sigma) e^{(i-1)\xi z} = 0$$

yes.

$$A^- A^- A^+ A^+ \\ E, E_z E, E_z$$

• •  
|  
A

~~EE~~  
(a, b)

$E, E_z$

$$E = a e^{+(i-1)\xi z} + b e^{-(i-1)\xi z}$$

$$E_z = (i-1)\xi a e^{(i-1)\xi z} - (i-1)\xi b e^{-(i-1)\xi z}$$

$$E(0) = a + b \quad \text{compare}$$

$$E_z(0) = (i-1)\xi a - (i-1)\xi b \quad \text{compare}$$

$$\begin{pmatrix} E \\ E_z \end{pmatrix}_0 = \begin{pmatrix} 1 & 1 \\ (i-1)\xi & -(i-1)\xi \end{pmatrix} \begin{pmatrix} a \\ b \end{pmatrix}$$

$$\begin{pmatrix} a \\ b \end{pmatrix} = \begin{pmatrix} 1 & (i-1)\xi & -1 \\ -2(i-1)\xi & & \end{pmatrix} \begin{pmatrix} E \\ E_z \end{pmatrix}_0$$

$$\det M_1 = -(i-1)\xi - (i-1)\xi$$

$$= -2(i-1)\xi$$



$$\begin{pmatrix} E \\ E_z \end{pmatrix}_h = \begin{pmatrix} e^{(i-1)\xi h} & e^{-(i-1)\xi h} \\ (i-1)\xi e^{(i-1)\xi h} & -(i-1)\xi e^{-(i-1)\xi h} \end{pmatrix} \begin{pmatrix} a \\ b \end{pmatrix}$$

$$\begin{pmatrix} E \\ E_z \end{pmatrix}_h = \begin{pmatrix} M_z \\ M_1 \end{pmatrix} \begin{pmatrix} E \\ E_z \end{pmatrix}_0$$

mtz program based on this algorithm compares well with oldenberg's 3-layer result.

GRD DATA from Section by Johnson, 1970

28 8 39 18 11 2.6 5.07

26 8 39 18 11 2.5 5.3

26 8 39 15 11 2.1 4.59

50 11 26 55

---

26 8 39 15 10 1.97 4.59

100 10 27.8 50

150 9.5 22.9 46

200 11 26 45

250 10.5 24.9 43

300 10.3 24.5 46

350 12.5 29.6 45

400 10.5 24.9 47

450 11 26 43

2.3025



URB DATA from Beblø & Björnson, 1980

T	<sup>26.5</sup>	$\rho_a$	$\phi$
25	13	30.9	50
35	12.5	29.6	55
50	11	26.0	55
70	<del>9.5</del> 10	23.8	52
100	10	23.8	50
150	9.5	22.8	46
220	11	26	45
350	10.5	24.9	43
480	10.3	24.5	46
700	12.5	29.6	45
1000	10.5	24.9	57
1500	11	26	53

Z. 3025

$$\frac{43}{25} (20-5)$$

$$8 (5-20)$$

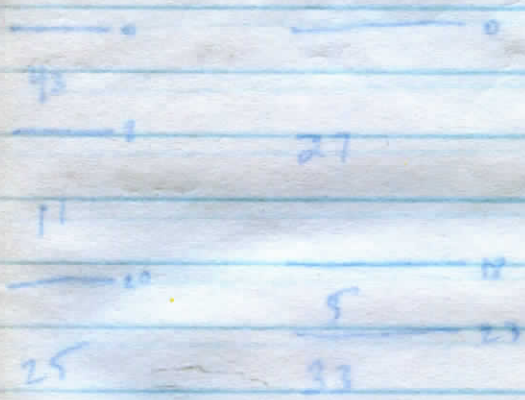
$$\frac{11}{25} (1-15)$$

$$4 (1-20)$$

$$25 (10-50)$$

initial error  $\rho_a = 2.00$   $\phi = 4.66$

about 80 s for inner 3 loops



$$43 + 11 = 54 / 2 = 27$$



$$\begin{array}{r} 43 \cdot 8 \\ \underline{11 \quad 12} \\ 25 \end{array}$$

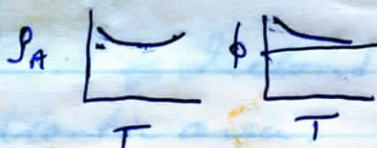
using exact  $E_{sa}$  (with  $\phi$ )

					$E_{sa}$	$E_{\phi}$
<del>26</del>	<del>5</del>	<del>2</del>	<del>15</del>	<del>3</del>	<del>1.9</del>	<del>3.93</del>
26	5	30	17	4	1.8	4.42
26	6	33	17	6	1.97	4.26
27	5	28	17	5	1.82	4.19
27	5	38	18	5	1.95	4.59
27	6	44	14	7	1.89	4.24

$$\begin{array}{r} \text{---} \circ \\ 43 \\ \text{---} 8 \\ 11 \\ \text{---} 20 \\ 25 \end{array} \qquad \begin{array}{r} \text{---} \circ \\ 27 \\ \text{---} 18 \\ 5 \\ \text{---} 23 \\ 33 \end{array}$$

$$43+11 = 54/2 = 27$$

43 8  
 11 4



25 *double: D. Reid* using exact data (mt4)  $P_a$   $\phi$

27	10	25	20	5	1.79	3.62
27	11	25	25	5	1.95	4.15
20	10	26	30	3	2.49	6.55

*Becked B. using MTH*

in relation to the mid-Atlantic Ridge based on *Journal of Plan Sci*, 2, 25, 1974

to *Geophysical J*, 87, 176-183, 1980  
 map of features of heat flow

Then had with the Gen of  
 Phil Tron May 20 to  
 11 283-646 Brown