

1                   **The marine hydrological cycle: the Ocean's floods and droughts**

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7    **Abstract:** The sea surface salinity (SSS) displays fluctuations that are not solely in  
8    response to local air-sea flux of freshwater, but also reflect ocean circulation and mixing  
9    processes. Ponte and Vinogradova (2016) using Estimating the Circulation and Climate  
10   of the Ocean output, estimate the relative roles of these forces for the global ocean. They  
11   find that the governing forces vary greatly across ocean regimes. Their research identifies  
12   features that will be addressed with enhanced SSS observations from orbiting satellites  
13   and *in situ* global arrays, which promise new insight into the marine water cycle and its  
14   place in the global hydrological system.

15   **1. Introduction:**

16           There is justified concern about floods and droughts on land, as water impacts  
17   agricultural production and the welfare of populations. The ability to predict changes in  
18   terrestrial freshwater availability in response to natural and anthropogenic induced  
19   climate change is of high priority, in that they enable preparedness and management of  
20   resources. It should come as no surprise that the ocean, amounting to 71% of the Earth's  
21   surface, also experiences floods and droughts, with periods of excess precipitation, or  
22   lack of it, reflected in changes of the sea surface salinity (SSS). Floods and droughts over  
23   land and ocean together mark a response, across a wide range of scales, of the global  
24   hydrological or water cycle (Schmitt, 2008; Schanze et al., 2010; Durack et al., 2012;  
25   Durack, 2015; Lagerloef et al., 2015). The terrestrial and marine components are joined  
26   through atmospheric moisture transport and continental runoff.

27           There have been many studies of the relationship of climate to sea surface  
28   temperature (SST), often presented in terms of climate indices (of which there are many).  
29   Orbiting satellite have provided synoptic global coverage of SST for the last half century,

30 defining the characteristics of SST anomalies and associated air-sea heat flux, which are  
31 dynamically linked as air-sea heat flux is governed by the difference between SST to air  
32 temperature. Specific SST anomalies have been related to terrestrial flood and drought,  
33 primary examples being the Pacific El Niño (Dai et al., 1998) and Atlantic SST  
34 anomalies linked to the Sahel drought (Giannini et al., 2003).

35 While there is satellite coverage of marine precipitation, with the Global  
36 Precipitation Mission, the direct detection of SSS anomalies, which reflects a broader  
37 range of air-sea and ocean factors, has been stymied by the lack of synoptic coverage.  
38 Blurry views of salinity, built upon many decades of ship-based data, depict a  
39 climatology that likely never exists at all points at the same time (Figure 1). We now  
40 have, at long last, near synoptic global views of SSS from orbiting satellites (Font et al.,  
41 2010; Lagerloef, 2012; Figure 1), as well as in situ coverage from the global array of  
42 Argo profilers and ocean drifters. These have finally “opened a window” to the marine  
43 component of the global hydrological cycle.

44 With the advent of near synoptic mapping of SSS from space we are now aware  
45 that SSS anomalies are also related to terrestrial flood and droughts (Li et al, 2016a,b,  
46 who link SSS to Sahel rainfall and to U.S. Midwest summer precipitation). Zhu et al.  
47 (2014) find that SSS variability plays an active role in ENSO evolution, essential to  
48 correctly forecasting the 2007/08 La Niña.

49 The global hydrological cycle in response to climate change is becoming more  
50 intense. Durack and Wijffels (2010) show that in the past five decades, the salty regions  
51 (e.g. the evaporation-dominant subtropical gyres) have become saltier and fresher regions  
52 (e.g., the tropical and high-latitude precipitation-dominant regions) have become fresher;  
53 the salinity contrast between the saltier Atlantic and fresher Pacific has also become  
54 larger, which has implication to the global thermohaline ocean circulation.

## 55 **2. Sea surface salinity governing processes:**

56

57 SSS responds to the amount of water exchange between the atmosphere and  
58 ocean: evaporation (E) and precipitation (P), as well as the formation and melting of sea

59 ice (sea ice having a salt content of ~20% of sea water). In this way SSS might be  
60 thought of as a rain gauge, but this is an over-simplification. Under the influence of E-P  
61 SSS would steadily increase or decrease were it not for compensating power of ocean  
62 processes. Ocean processes involve horizontal and vertical advection, including wind  
63 induced Ekman transport, the spreading of river runoff (R) and mixing by ocean eddies  
64 and vertical diffusion, that result in net divergence or convergence of freshwater to offset  
65 E-P. Air-sea exchange and the array of ocean processes achieve a quasi-steady state  
66 condition. However, the balance is not perfect, resulting in SSS variability across a wide  
67 range of spatial and temporal scales.

68

69 Regional SSS differences (Figure 1) reflect their place in the global climate  
70 system and ocean circulation. Even regions in the same climate belt, such as the salty  
71 subtropical regimes display differences from each other (Gordon et al., 2015), as they  
72 respond to local conditions. To understand the factors that affect SSS we must develop a  
73 quantitative grasp of the governing physics, which involve air-sea exchange and ocean  
74 processes: which are dominant? which are minor players?

75

76 Yu (2011) using an array of observationally based data compilations present a  
77 global view of the dominant process governing seasonal surface mixed layer salinity  
78 (effectively SSS), finding that freshwater fluxes across the air-sea interface, on a global  
79 basis, account for nearly 14% of the seasonal SSS changes, with the seasonal changes in  
80 Ekman advection accounting for 13%. The mean annual Ekman transport and  
81 entrainment at the base of the mixed layer are each about 7%. These 4 terms account for  
82 about 40% of the total SSS seasonal variance, with the remainder being a consequence of  
83 uncertainty, mainly in the form of mixed layer dynamics. The processes driving seasonal  
84 SSS variability, as expected, display large spatial variability. For example, E-P is the  
85 first dominant term, accounting for 40 to 70% of SSS seasonal variability, in the rainy  
86 intertropical convergence zones, and sections of evaporative subtropical regions, whereas  
87 closer to the equator and at higher latitudes Ekman advection and vertical entrainment  
88 take on larger roles.

89

90 Ponte and Vinogradova (2016; PV16) noting the uncertainty of observational data  
91 use "a model-data synthesis efforts that are both close to the observations and that allow  
92 for calculation of full, closed property budgets". They use ECCO (Estimating the  
93 Circulation and Climate of the Ocean; Wunsch, 2009) version 4, release 1, which is  
94 optimized to fit an array of observations, taking into account their uncertainties, to  
95 determine the contributions to SSS (upper 10 m) variability of horizontal and vertical  
96 advection (ocean currents, including Ekman transport), diffusion and the air-sea flux of  
97 freshwater (E-P, runoff and sea ice formation/melting). PV16 find that the sum of  
98 advection and diffusion and air-sea flux are close to zero for the 18-year (1993-2010)  
99 averaged ECCO output, as would be expected if they properly incorporate the essentials  
100 of the freshwater fluxes.

101

102 PV16 maps clearly show that the relative contributions of each component change  
103 significantly across the ocean regimes. Diffusion and air-sea flux generally compensate  
104 each other, while advection can boost or decrease SSS. Diffusion and air-sea flux are the  
105 dominant factors in the subtropics and subpolar regimes, while the tropics and polar  
106 regions advection becomes a more dominant component. Advection is particularly  
107 significant in the Atlantic Ocean. At low latitudes advection acts to decrease SSS, both in  
108 regions of excess P and in excess E. As PV16 say: "Given the variable nature of  
109 horizontal gradients in [SSS] and also surface currents, the large scale, single sign  
110 behavior of [advection] in the tropics is not a trivial result"; PV16 suggest this issue  
111 "merits future scrutiny". Within the Bay of Bengal, advection and diffusion balance each  
112 other, with air-sea fluxes playing a reduced role. In the polar areas of net sea ice  
113 production and ocean diffusion boost SSS, which are then compensated by advection.

114

### 115 **3. Summary and way forward:**

116

117 The Yu (2011) and PV16 results clearly show that variability in air-sea freshwater  
118 fluxes alone do not properly track changes in SSS: the full range of ocean processes are  
119 needed, to resolve the spatial and temporal changes of the marine hydrological cycle and  
120 its place in the global system. The near synoptic global view of the SSS from space, as

121 well as the enhanced in situ observational array, has opened the door to a more  
122 quantitative understanding and appreciation of the marine hydrological cycle and of its  
123 governing physics.

124

125 The SPURS (Salinity Processes in the Upper-ocean Regional Study) 1 and 2 field  
126 programs (Lindstrom et al., 2015; SPURS-2 Planning Group, 2015) are delving into these  
127 issues, made possible by our new views of SSS, which will no doubt sharpen as the  
128 observational data sets grow and model simulations incorporate the essential elements to  
129 finer resolution. We will finally acquire an understanding of the full, global hydrological  
130 cycle and its coupling to the natural and anthropogenic forced climate change. As often  
131 happens, we can expect new observations to open the door to discovery.

132

133

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211 **Figure**

212

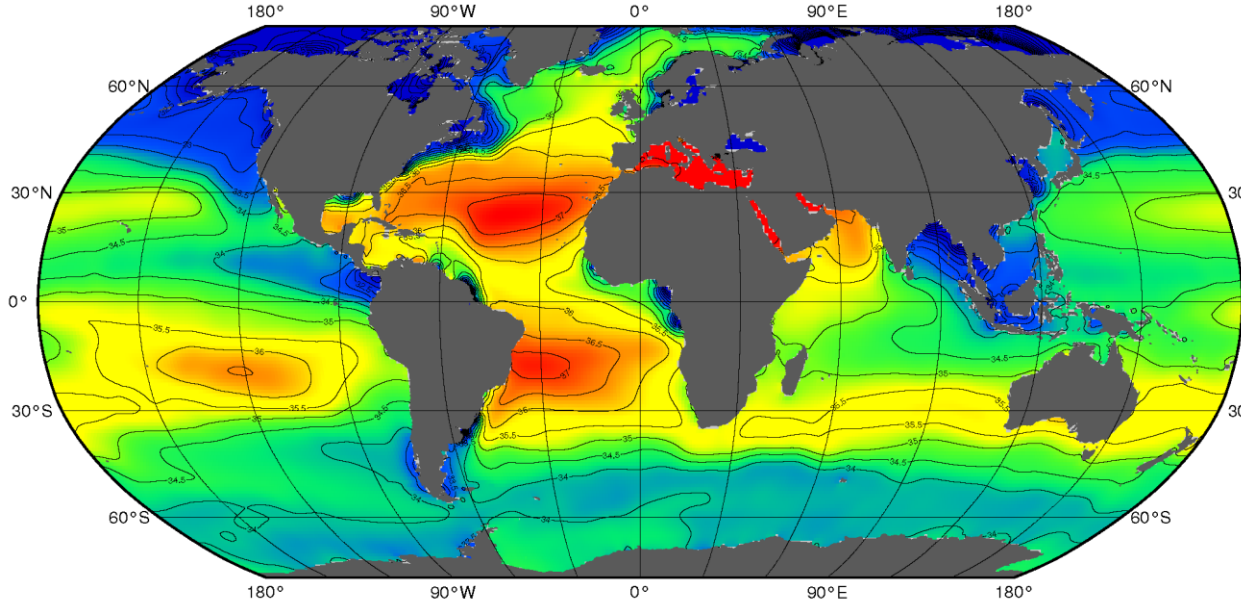
213 Figure 1. *Left Panel:* Sea surface salinity (SSS; Antonov et al., 2009) constructed from  
214 ship data obtained over many decades depict SSS climatology that likely never exists at  
215 all points at the same time.

216 *Right Panels:* With the advent of orbiting satellite views of SSS (NASA/CONAE  
217 Aquarius/SAC-D satellite mission; ESA's Soil Moisture and Ocean Salinity, SMOS; and  
218 Soil Moisture Active Passive, SMAP) we now have, at long last, near synoptic, global  
219 views of the SSS. The Aquarius satellite data, <http://podaac.jpl.nasa.gov/aquarius>, enable  
220 construction of global SSS maps at 7-day resolution. *Upper right panel:* the first week of  
221 January 2014; *lower right panel:* the first week of July 2014. The ever changing SSS field  
222 provide insight to the complex realm of the global water cycle, and terrestrial and ocean  
223 flood and droughts.

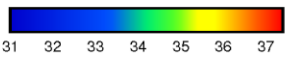
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**Figure 1. Left Panel: Sea surface salinity (SSS; Antonov et al., 2009) constructed from ship data obtained over many decades depict SSS climatology that likely never exists at all points at the same time. Right Panels: With the advent of orbiting satellite views of SSS (NASA/CONAE Aquarius/SAC-D satellite mission; ESA's Soil Moisture and Ocean Salinity, SMOS; and Soil Moisture Active Passive, SMAP) we now have, at long last, near synoptic, global views of the SSS. The Aquarius satellite data, <http://podaac.jpl.nasa.gov/aquarius>, enable construction of global SSS maps at 7-day resolution. Upper right panel: the first week of January 2014; lower right panel: the first week of July 2014. The ever changing SSS field provide insight to the complex realm of the global water cycle, and terrestrial and ocean flood and droughts.**

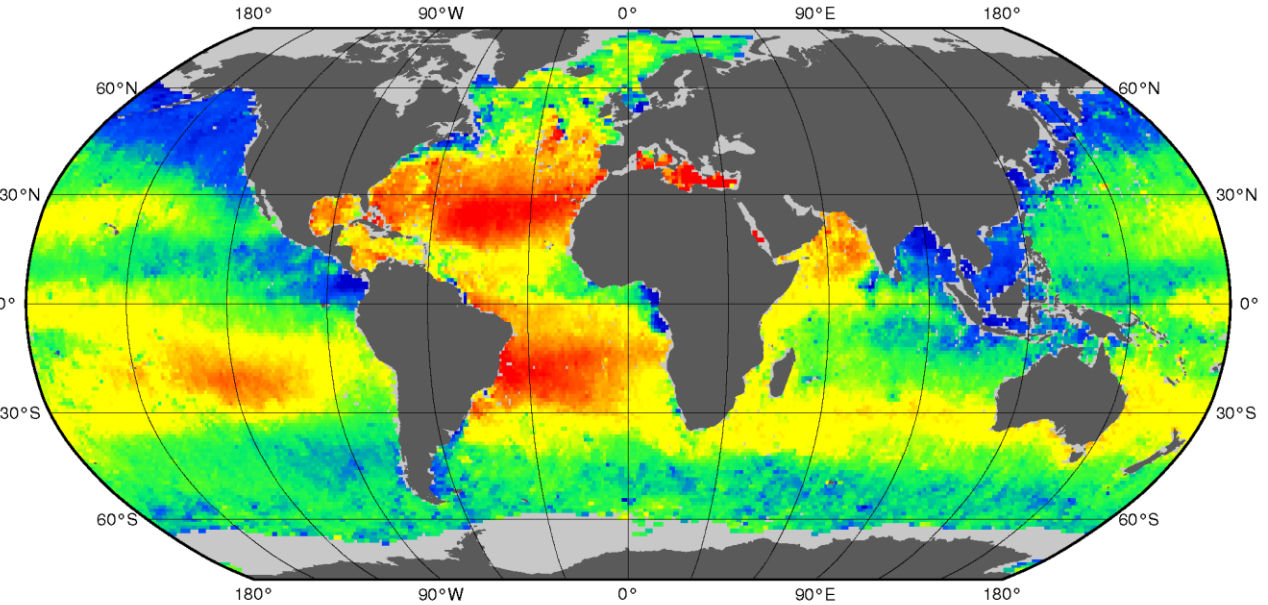


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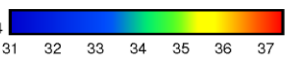


A horizontal color scale for salinity, ranging from 31 (dark blue) to 37 (dark red), with intermediate values 32, 33, 34, 35, and 36.

Annual mean salinity

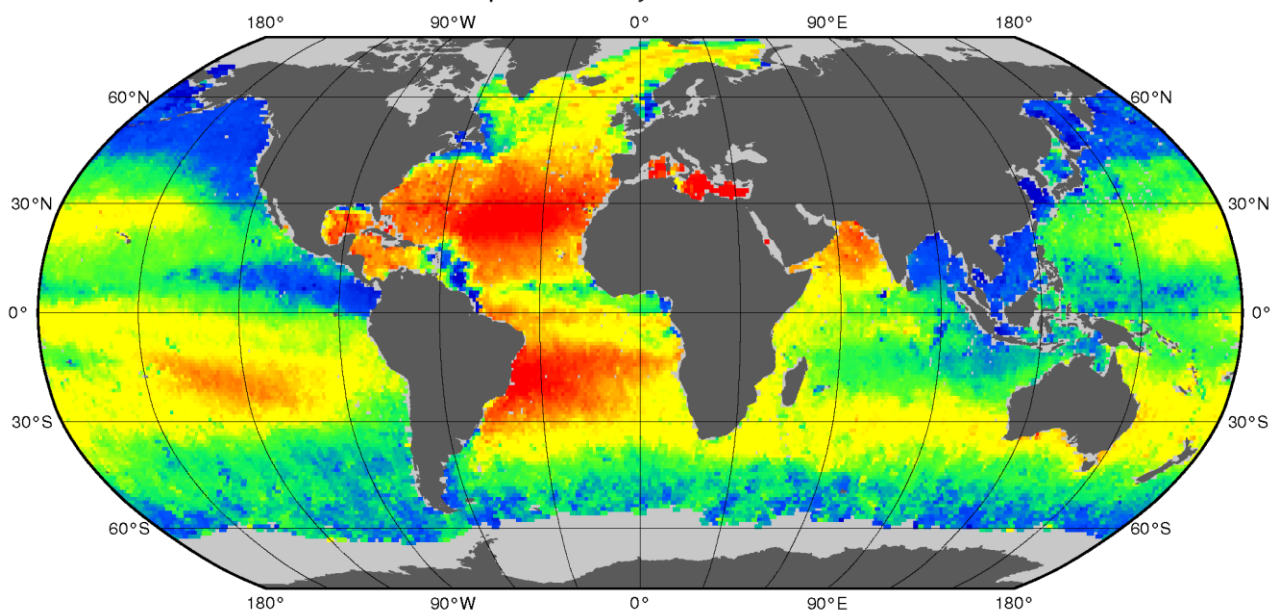


SSS: Aquarius V4

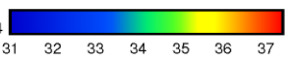


A horizontal color scale for salinity, ranging from 31 (dark blue) to 37 (dark red), with intermediate values 32, 33, 34, 35, and 36.

Aquarius salinity: 1-7 Jan 2014



SSS: Aquarius V4



A horizontal color scale for salinity, ranging from 31 (dark blue) to 37 (dark red), with intermediate values 32, 33, 34, 35, and 36.

Aquarius salinity: 2-8 Jul 2014