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Journal of Geochemical Exploration xx (2005) xxx-xxx



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#### $\mathbf{2}$ Preliminary evidence of a link between surface soil properties and the 3 arsenic content of shallow groundwater in Bangladesh

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19	Dessived 26 March 2005, accorted 10 August 2005
15	Received 26 March 2005; accepted 19 August 2005
14	

#### Abstract 15

The extremely heterogeneous distribution of As in Bangladesh groundwater has hampered efforts to identify with certainty the 1617mechanisms that lead to extensive mobilization of this metalloid in reducing aquifers. We show here on the basis of a high-18resolution transect of soil and aquifer properties collected in Araihazar, Bangladesh, that revealing tractable associations between As concentrations in shallow (<20 m) groundwater with other geological, hydrological, and geochemical features requires a lateral 1920sampling resolution of 10-100 m. Variations in the electromagnetic conductivity of surface soils (5-40 mS/m) within a 500 21  $m \times 200$  m area are documented with 560 EM31 measurements. The results are compared with a detailed section of groundwater 22As concentrations  $(5-150 \text{ }\mu\text{g/L})$  and other aquifer properties obtained with a simple sampling device, "the needle-sampler", that 23builds on the local drilling technology. By invoking complementary observations obtained in the same area and in other regions of 24Bangladesh, we postulate that local groundwater recharge throughout permeable sandy soils plays a major role in regulating the As 25content of shallow aquifers by diluting the flux of As released from reducing sediments.

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Keywords: Arsenic; Groundwater; Well water; Bangladesh; Health 27

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#### 1. Introduction 29

30 The landmark survey of groundwater pumped from 31thousands of tube wells in Bangladesh conducted by 32DPHE/MMD/BGS (1999) and BGS/DPHE (2001) has demonstrated that the concentrations of As in aquifers 33

throughout Bangladesh is spatially highly variable on 34scales of 1-100 km. Even after depth and broad-scale 35 patterns associated with the geology of the country are 36 taken into account, the average arsenic content of wells 37 in a given village, let alone individual wells, is difficult 38to predict (BGS and DPHE, 2001; van Geen et al., 392003a; McArthur et al., 2004; Horneman et al., 402004). One of our group's goals for the past several 41 years has been to determine if tractable changes in 42aquifer properties that are related to groundwater As 43

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<sup>0375-6742/\$ -</sup> see front matter © 2005 Elsevier B.V. All rights reserved. doi:10.1016/j.gexplo.2005.08.106

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concentrations are observed if they are mapped at spatial scales <1 km. We summarize here some of these results by focusing on shallow (<20 m) aquifers and augment recent observations for one particular study area in Araihazar, Bangladesh, with a densely spaced set of measurements of the electromagnetic conductivtiv of surface soils.

# 51 2. Geological setting

52Electromagnetic conductivity data and profiles of 53groundwater and sediment properties were collected within a 25 km<sup>2</sup> region of Araihazar upazila where the 54distribution of groundwater As has previously been 5556documented by sampling and analyzing groundwater from over 6500 wells (van Geen et al., 2003a, in 5758press). Columbia University scientists and Bangladeshi 59partners have since 2000 been investigating the health 60 effects of elevated As on a cohort of 12000 recruited 61 from the 70 000 inhabitants of the area, as well as various 62 approaches to mitigation (van Geen et al., 2002, 63 2003a,b). Araihazar upazila straddles the present 64Meghna River floodplain to the southeast and older deposits of the uplifted Madhupur tract to the northwest 6566 (BGS and DPHE, 2001; Goodbred et al., 2003; Zheng et 67 al., submitted for publication). The majority of shallow wells in the area tap thick and relatively recent Holocene 68 69 deposits frequently associated with elevated groundwa-70 ter As concentrations (van Geen et al., 2003a).

71 This report focuses on a 500-m transect that extends 72 west of Lashkardi village where the vast majority of 73 wells contain <50 µg/L As (van Geen et al., 2004b). 74The western portion of the transect starts from a small 75village (Rishir Char), where a limited number of exist-76ing private wells and a nest of monitoring wells indicate 77elevated groundwater As concentrations in the shallow aquifer (van Geen et al., in press). The westernmost 7879profile of this section was collected on the margin of an 80 abandoned stream channel north of Rishir Char.

### 81 3. Methods

# 82 3.1. EM induction survey

83 The geophysical survey of the study area was con-84 ducted with a Geonics® EM31 instrument (McNeill, 85 1980). By generating an eddy current in the ground, the primary field generates a secondary electromagnet-86 ic field that is recorded by the receiving coil. The 87 88 intensity of the secondary field increases with the 89 conductivity of the ground. When the boom is held horizontally at waist height, which corresponds to a 90

vertical dipole orientation, 50% of the signal is generated in the upper 90 cm of the soil whereas only 92 27% of the signal reflects the conductivity of layers 93 below 180 cm depth (McNeill, 1980; Doolittle et al., 94 2001). 95

### 3.2. Needle-sampler profiles 96

The needle-sampler consists of an evacuated sample 97 chamber capped with a silicone stopper, a long needle 98 and plunger assembly to transfer groundwater and sed-99 iment from a depth slightly greater than that of the drill 100 hole to the sample chamber, and a housing unit that 101 connects the needle and plunger assembly to the sample 102 chamber (van Geen et al., 2004b). The device is 103 deployed by drilling to the targeted depth using the 104 entirely manual method that has been used to install 105the vast majority of the millions of wells in Bangladesh 106(Horneman et al., 2004). The chamber fills almost 107entirely with groundwater and some sediment during 108deployment. Groundwater samples filtered under nitro-109gen were acidified and analyzed for As by high reso-110 lution inductively coupled plasma mass spectrometry 111 (HR ICP-MS, Cheng et al., 2004). The effective detec-112tion limit of the method for As in unfiltered samples is 113 $\sim 1 \ \mu g/L$ ; the precision is on the order of  $\pm 2\%$  (van 114 Geen et al., 2005). 115

### 4. Results

EM conductivities measured in the fields of this 117 portion of Araihazar span a range of 5 to 40 mS/m 118(n=560). The highest conductivities were measured in 119 the northwestern portion of the study area, near Rishir 120Char, and gradually declined towards the east (Fig. 1). 121 A diagonal swath of particularly low conductivities 122(<10 mS/m), ~100 m wide, characterizes the rice 123paddies near Lashkardi. 124

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Dissolved As concentrations measured in groundwa-125ter obtained with the needle-sampler are consistent with 126As levels measured in shallow wells of Rishar Char and 127Lashkardi. The spatial features of a transition from 128elevated groundwater As concentrations in the west to 129low As concentrations in the east are illustrated with a 130 contour plot based on 5 profiles ~100 m apart from 131each other (Fig. 1). The contour corresponding to 50 132µg/L As, the Bangladesh standard for drinking water, is 133crossed at a depth of 50 ft (15 m) at NS-F, with 134Lashkardi. Groundwater As concentrations barely 135reach this level at the same depth in the next two 136profiles to the west (NS-1, NS-2), and then actually 137 decrease below this depth. Further west along the tran-138

# **ARTICLE IN PRESS**

A. van Geen et al. / Journal of Geochemical Exploration xx (2005) xxx-xxx



Fig. 1. Comparison of the distributions of surface EM conductivity and groundwater As concentrations across the rice paddies that separate two contrasting villages of Araihazar, Bangladesh. The background for the two upper panels is a portion of a  $60 \text{ km}^2$  IKONOS satellite image of the area (van Geen et al., 2003a). Color-coded circles in the same panels show the location and As concentrations in privately owned tube wells. Color-coded triangles in the upper panel indicate the location and range of EM31 measurements. The false color of the satellite image in the middle panel emphasizes in purple areas of high water content. The numbered triangles indicate where 5 needle-sampler profiles of shallow aquifer properties were obtained (van Geen et al., in press). The two white circles show the locations of nests of monitoring wells where profiles of the age of groundwater since recharge have been determined by the  ${}^{3}\text{H}$ - ${}^{3}\text{H}$  dating method (Stute et al., submitted for publication). The lower panel shows a cross-section of groundwater As concentrations that is based on the 5 needle-sampler profiles. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

139 sect, the minimum in dissolved As content of the 140 shallow aquifer that is located roughly mid-way be-141 tween the two villages is bound by the sharp transition 142 to NS-3 where concentrations exceed 100  $\mu$ g/L from 30 143 to 60 ft depth (9–18 m). At the western end of the 144 transect (NS-4), the depth range of intervals where As 145 concentrations exceed 100  $\mu$ g/L broadens to include all 146 but the deepest sample.

# 147 **5. Discussion**

148 A number of studies have shown that hand-held EM 149 instruments can be used to determine the spatial variability of soil salinity (Doolittle et al., 2001). It has, 150however, typically been difficult to disentangle the direct 151contribution of fine-grained particles to EM conductivity 152from an indirect contribution due to the accumulation of 153salt by evaporation at the surface of impermeable depos-154its. The likely contribution of both factors to EM con-155ductivity in Araihazar was recently shown by collecting 156a series of hand-auger cores and analyzing soil samples 157for grain-size and slurry conductivity (Aziz et al., in 158preparation). Layers of fine silt dominate the upper 2 m 159of soil in the western portion of the study area. In 160contrast, the walls of shallow pits in the eastern part of 161the study area are composed mostly of fine sand. 162

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163Additional data obtained from the same area are 164 important for interpreting subsurface patterns of 165 groundwater As and other aquifer properties. These 166 are measurements of groundwater age relative to re-167 charge obtained by the tritium-helium  $({}^{3}H-{}^{3}He)$  dating 168 technique for nests of monitoring wells located at NS-F 169 in Lashkardi and near NS-5 in Rishir Char (Stute et al., 170 submitted for publication). The measurements divide 171 the 2 areas into two distinct hydrologic regimes. Near 172 NS-F, estimated groundwater ages increase from <1 173 year at  $\sim$ 25 ft (8 m) to  $\sim$ 4 years at 60 ft (18 m). At 174 Site G, instead, groundwater appears to be already 175 several years old in the shallowest monitoring wells 176 and averages 30 years at a depth of 60 ft (18 m). The simplest, though still tentative, explanation for 177178 the observed contrast in groundwater ages is that the 179 sandy deposits surrounding Lashkardi village permit 180 local recharge during the wet season whereas recharge 181 is inhibited by the finer deposits capping the aquifers 182 near Rishir Char. Groundwater dating in several addi-183 tional villages of Araihazar supports the notion that 184 recharge of shallow aquifers is more rapid in those 185 villages fortunate enough to have low As concentrations 186 in their wells (Stute et al., submitted for publication). 187 The high-resolution transect of profiles connecting Site 188 F and Site G indicates that the effect on groundwater As 189 concentrations of contrasting recharge rates extends at 190 least 100 m laterally beyond the location of the nests of 191 wells where  ${}^{3}H-{}^{3}He$  data were obtained (Fig. 1).

# 192 6. Conclusions

193The concentration of As in groundwater pumped 194 from any particular well in Bangladesh is the result of 195 multiple and still poorly known processes (Nickson et 196 al., 1998; Harvey et al., 2002; McArthur et al., 2004; 197 Zheng et al., 2004). The results presented here illustrate 198 for one particular location an apparently intimate con-199 nection between the permeability of surface soils, 200 which can be mapped by EM conductivity, and the 201 As content of shallow groundwater, which can be 202 established with the needle-sampler without installing 203 a well. Establishing causation from spatial correlations 204 of dissolved As concentrations with other properties of 205 Bangladesh aquifers is likely to require more systematic 206 use of these or similar tools, as well as coordination of 207 combined hydrological, geochemical, and microbiolog-208 ical investigations.

#### 209 7. Uncited reference

210 van Geen et al., 2004a

### Acknowledgments

Our arsenic research in Bangladesh has been supported principally by NIEHS Superfund Basic Research212Program grant NIH 1 P42 ES10349 and NSF grant214EAR 03-45688.215

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#### A. van Geen et al. / Journal of Geochemical Exploration xx (2005) xxx-xxx

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