

**The Annual Cycle of Precipitation
over the Tropical Atlantic, South America, and
Africa.**

Mutual Influences of Land and Ocean.

Michela Biasutti

with David S. Battisti and Edward S. Sarachik

New York, October 3, 2003

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- Motivation
- Method
- Results
- Conclusions
- Implications

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 - Similarity between TAV and Annual Cycle (AC)
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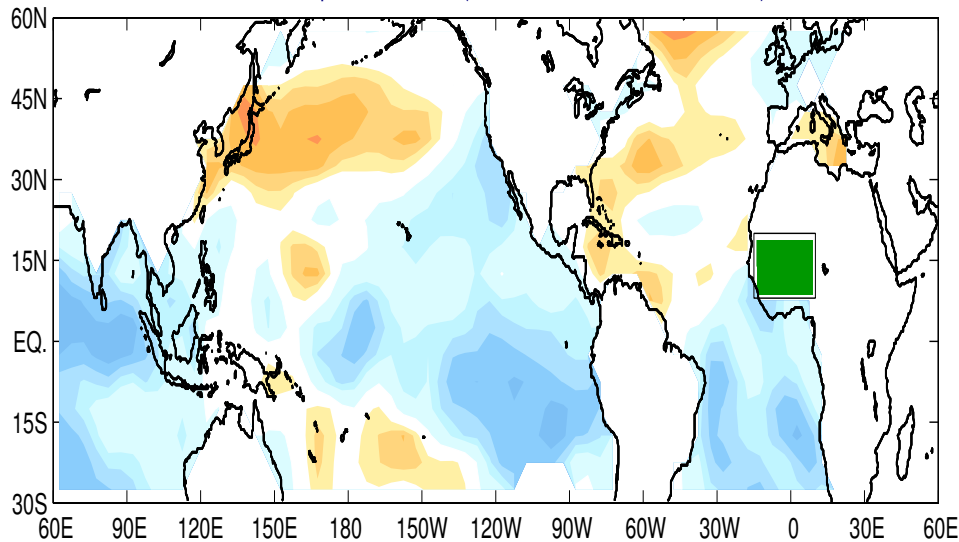
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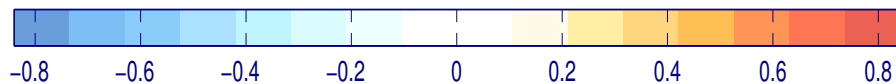
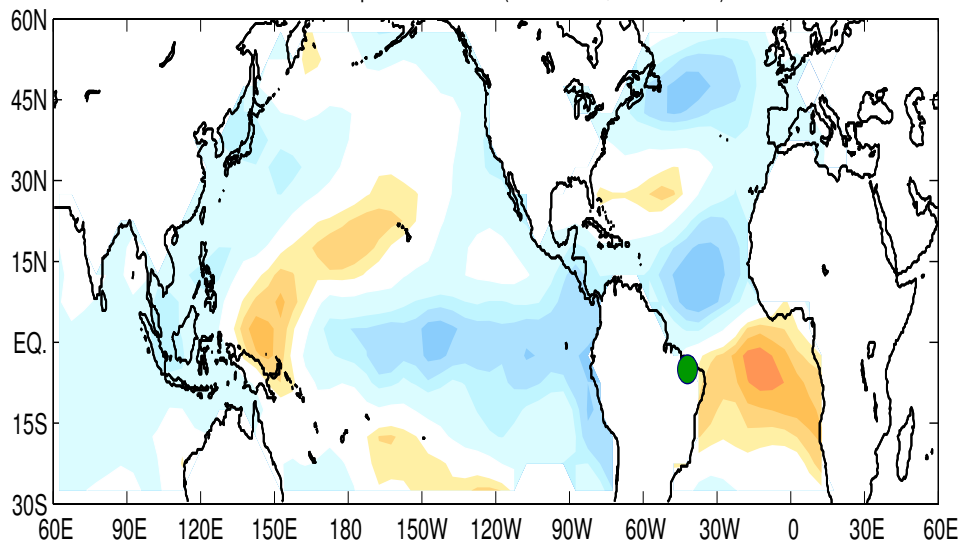
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Motivation: SST & LAND PRECIPITATION

Mitchell/Kaplan 1950:2000 (JJASO Sahel Rainfall, JJAS SST)



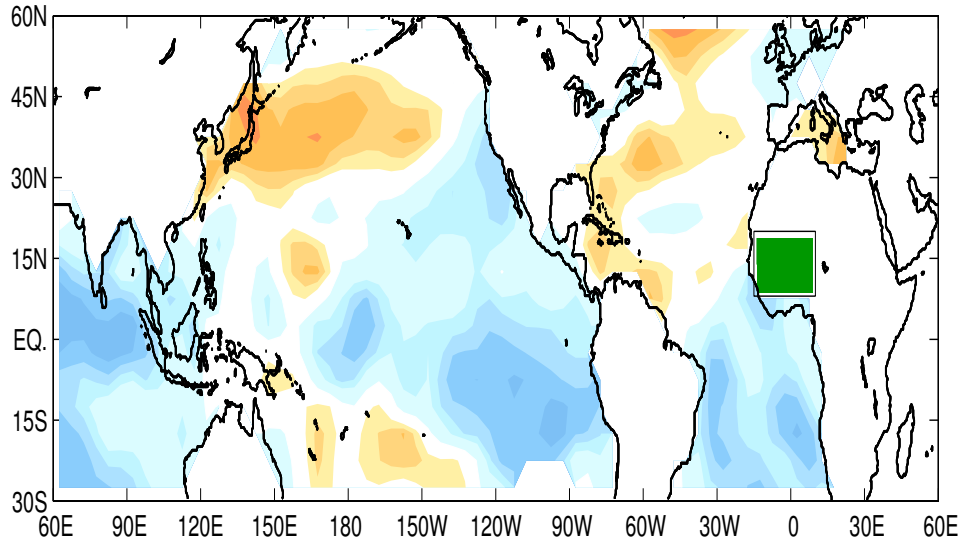
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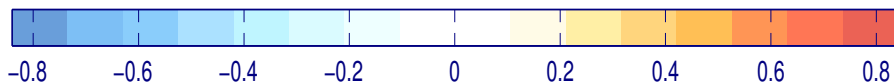
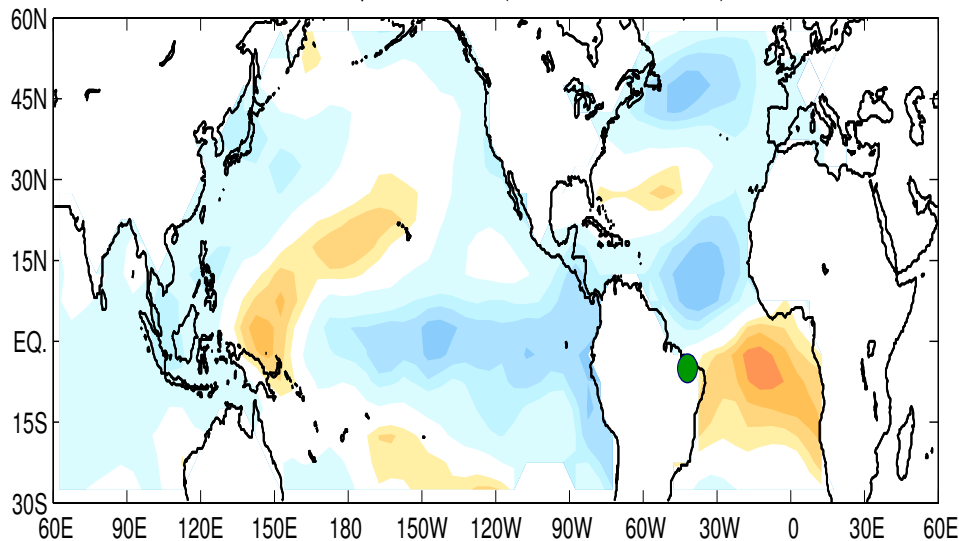
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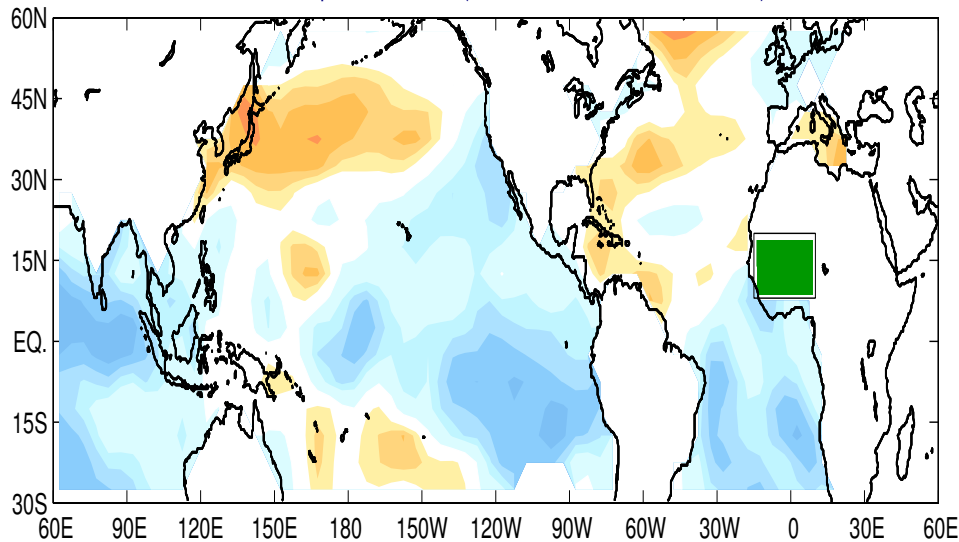
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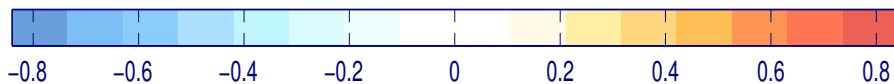
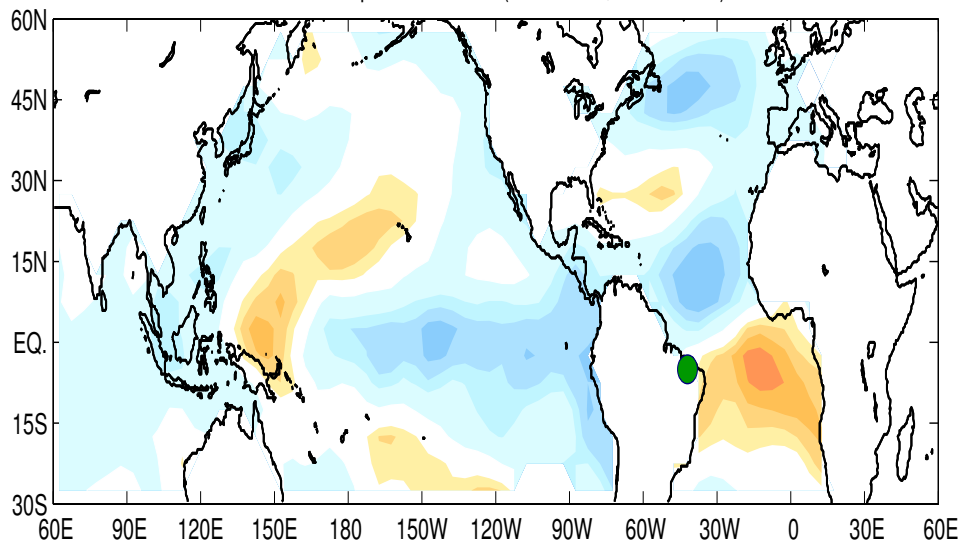
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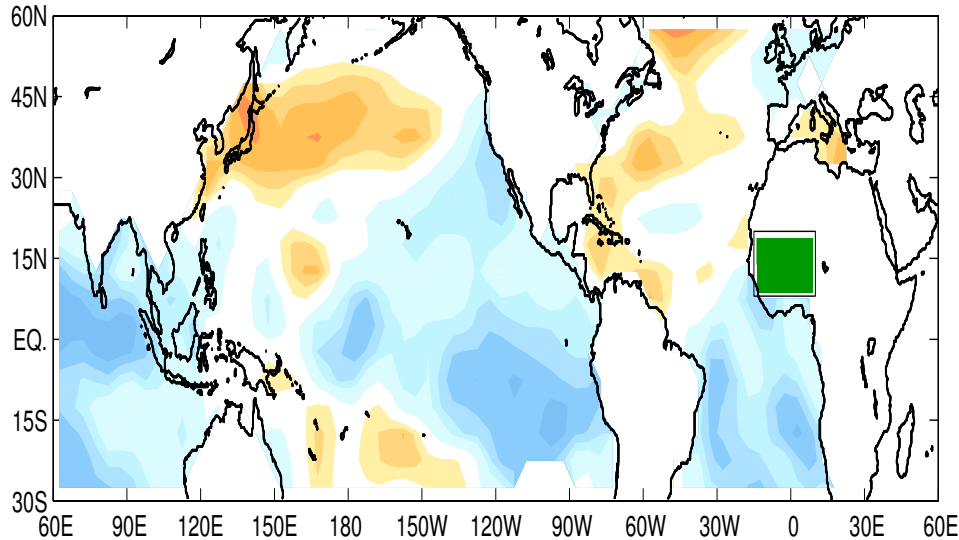
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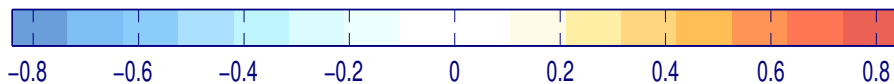
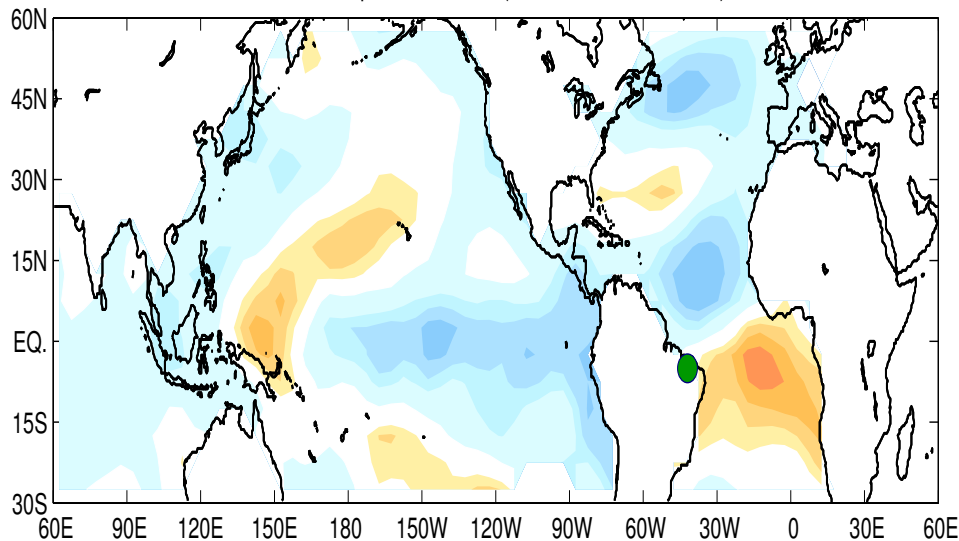
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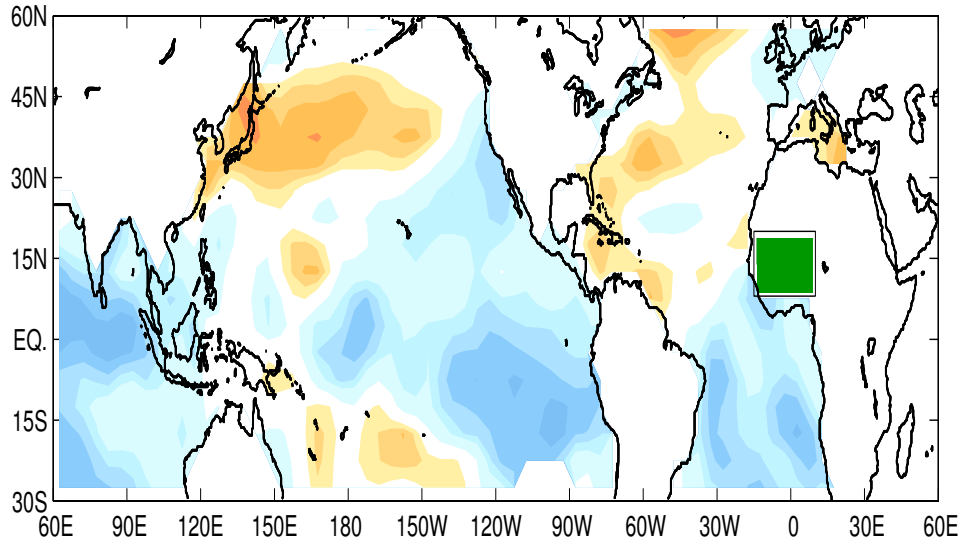
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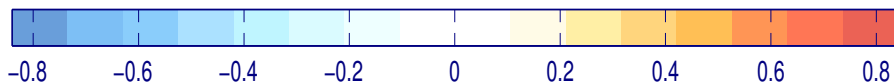
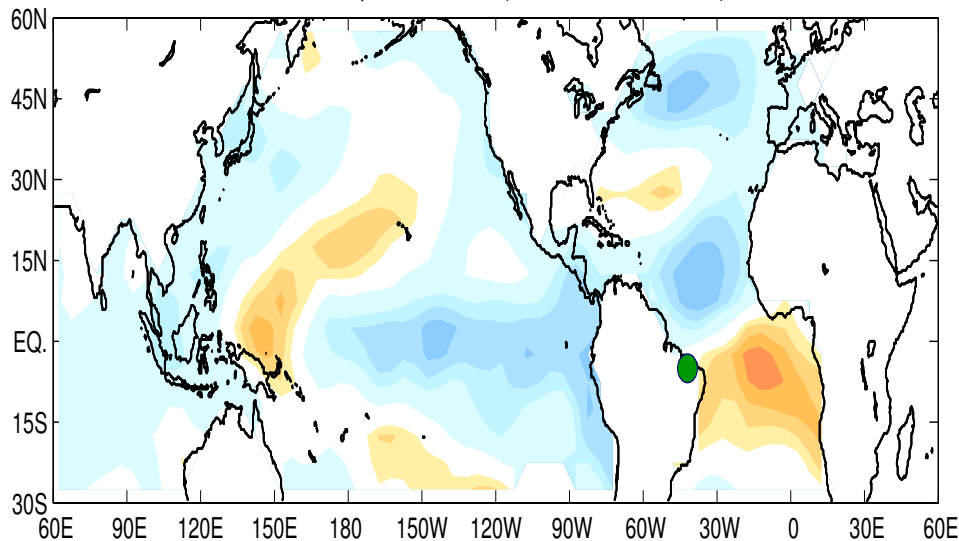
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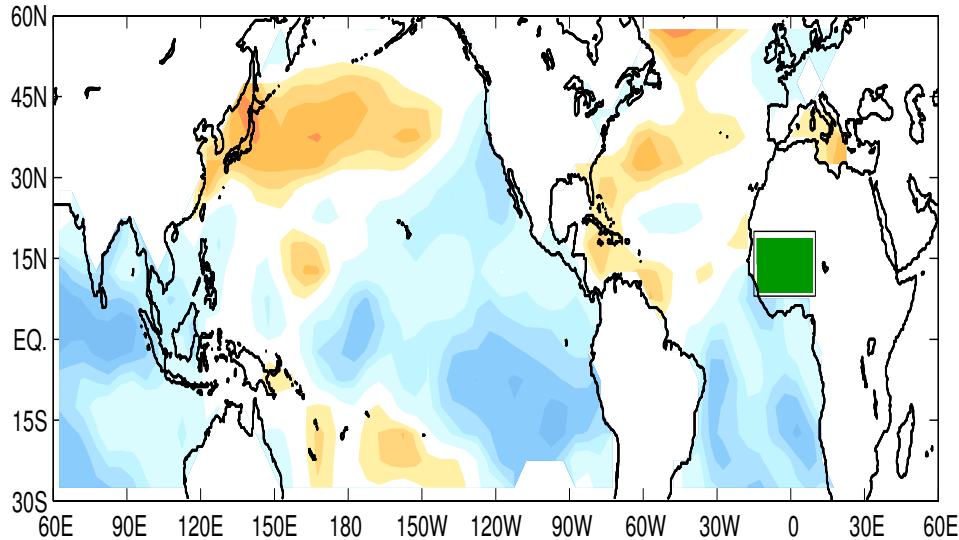
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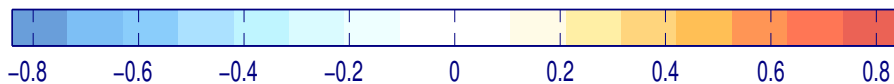
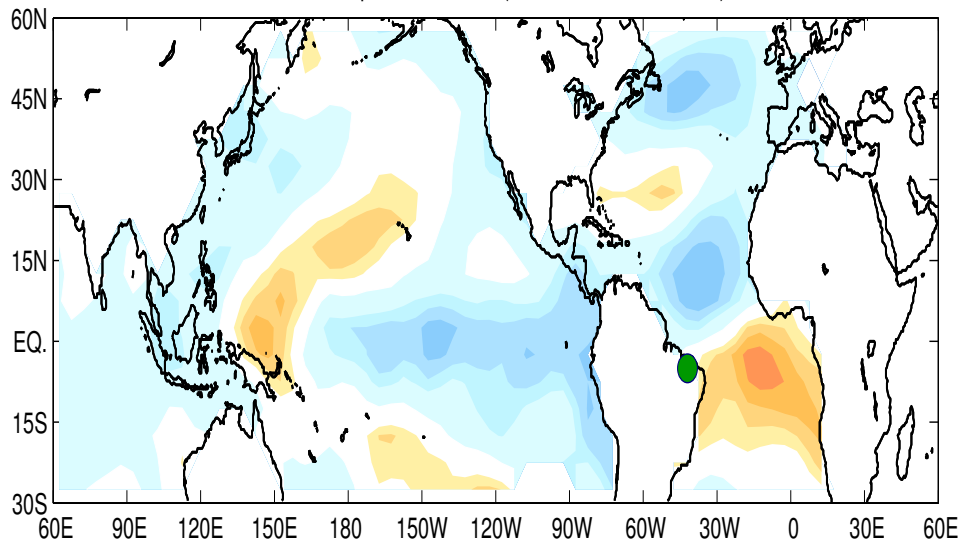
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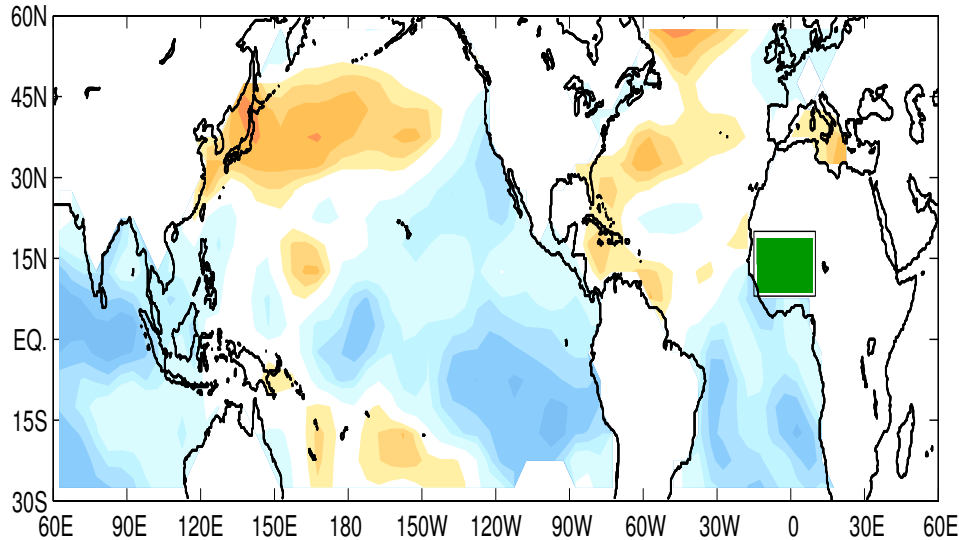
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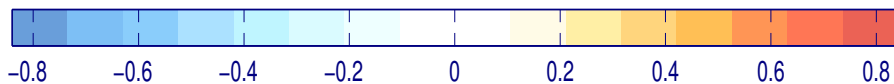
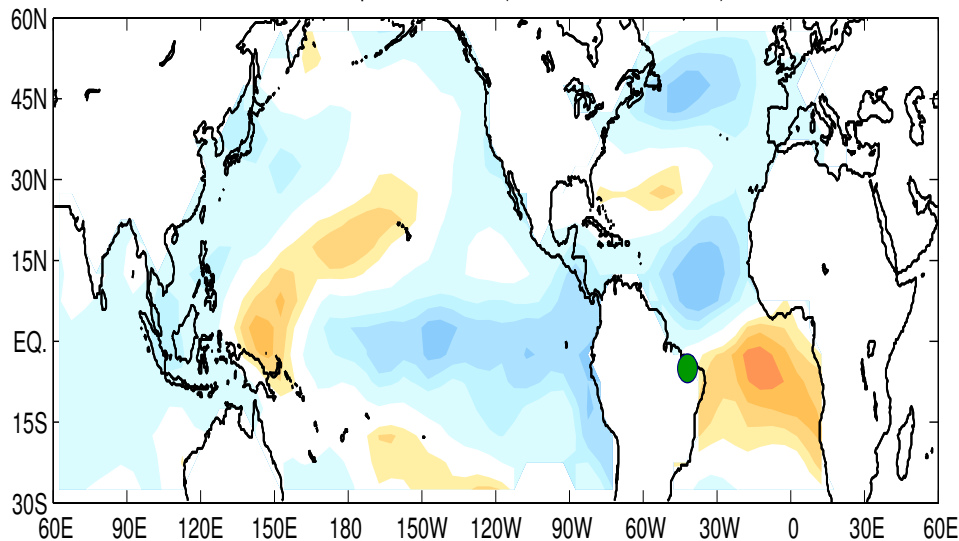
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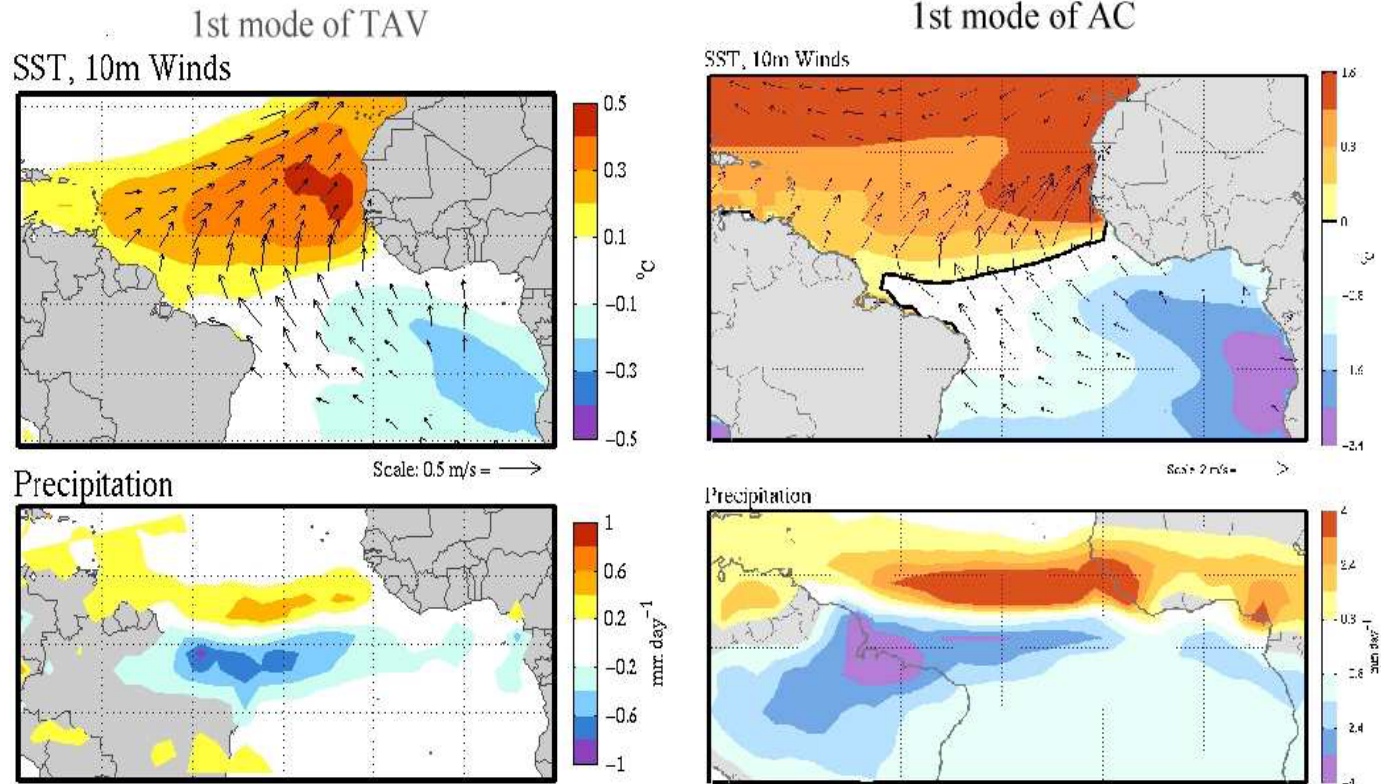


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Motivation: AC & TAV ANALOGUE

- Atmospheric processes bring the SST signal inland \Rightarrow the same fast atmospheric dynamics must be involved in both the AC and the TAV.
- The gross features of AC and TAV in the deep tropics look alike

Motivation: AC & TAV ANALOGUE



meridional mode

cross-equatorial SST gradient

cross-equatorial wind towards warm SST

meridional shift of ITCZ towards warm SST

Method: THE MODELS

- Uncoupled Community Climate Model Version 3 (CCM3) T42 resolution.
 - ⇒ The precipitation over Africa and South America is overestimated.
 - ⇒ Tropical land is too cold (especially the Sahara).
- Coupled CCM3 coupled to a Slab Ocean Model (SOM) in the tropical Atlantic only (prescribed SST elsewhere).
 - ⇒ SST responds to radiative and turbulent heat fluxes.
 - ⇒ No ocean dynamics (Ocean Heat Transport Convergence is parameterized with a flux adjustment, the Q-flux).
 - ⇒ Climatology of CCM3+SOM is (by construction) nearly identical to the climatology of CCM3.

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Method:
THE IDEA BEHIND THE EXPERIMENTS

We want to identify:

- what features of the AC over land and ocean are **locally** forced?
- what features of the AC over land and ocean are **remotely** forced?
- the mechanisms of mutual influences between land and ocean regions

IDEA: we suppress the AC of forcings over land and ocean regions **separately** \Rightarrow we can separate out local and remote responses.

Method: THE FORCINGS

- Insolation over land.
- Insolation over ocean.
- Q-flux. } or SST
- Elevated Condensational Heating
(Q_{cond}) in selected areas

Method: EXPERIMENTAL DESIGN

- Fixed insolation over land, AC of Ocean Forcings ($Q_{\text{flux}} + \text{Insolation}$ or SST)
- Fixed Ocean Forcings, AC of insolation over land

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AC over **LAND** is a **REMOTE** response to AC over ocean
AC over **OCEAN** is a **LOCAL** response

- Fixed Ocean Forcings, AC of insolation over land

Method: EXPERIMENTAL DESIGN

- Fixed insolation over land, AC of Ocean Forcings (Qflux+Insolation or SST)



AC over LAND is a REMOTE response to AC over ocean
AC over OCEAN is a LOCAL response

- Fixed Ocean Forcings, AC of insolation over land

Method: EXPERIMENTAL DESIGN

- Fixed insolation over land, AC of Ocean Forcings (Qflux+Insolation or SST)



AC over LAND is a REMOTE response to AC over ocean
AC over OCEAN is a LOCAL response

- Fixed Ocean Forcings, AC of insolation over land



AC over LAND is a LOCAL response
AC over OCEAN is a REMOTE response to AC over land

RESULTS

A. LOCAL LAND

- AC of insolation;
March SST
⇒ AC of land precip

B. LOCAL OCEAN

- AC of SST;
March Insolation
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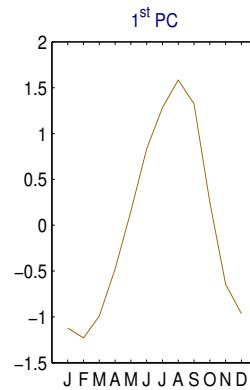
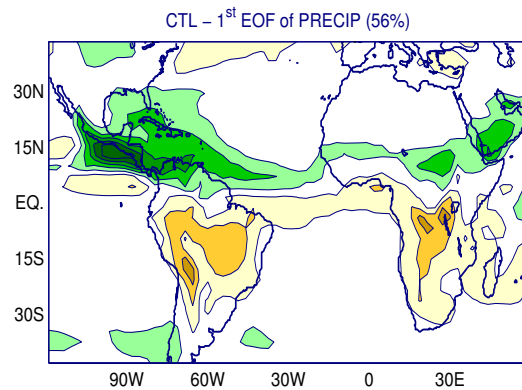
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D. REMOTE OCEAN

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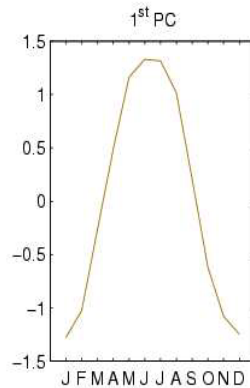
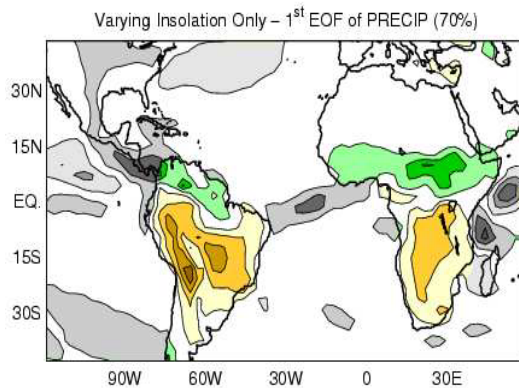
Results: LOCAL LAND



CTL:

North-South displacement over land and ocean

Extrema: February/August



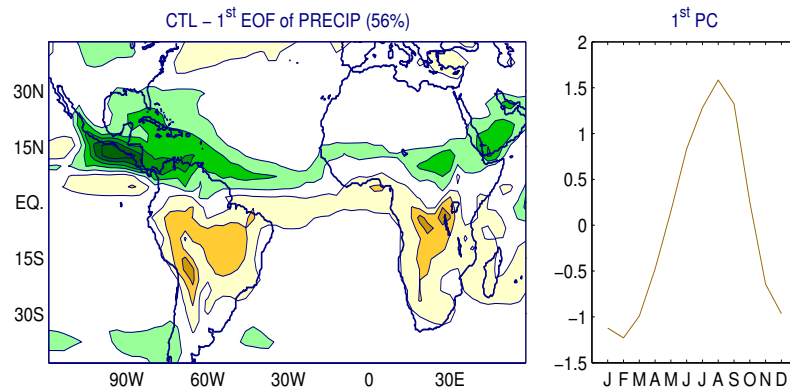
Local Land:

North-South displacement over land

Extrema: December/June

Gulf of Guinea is way off!

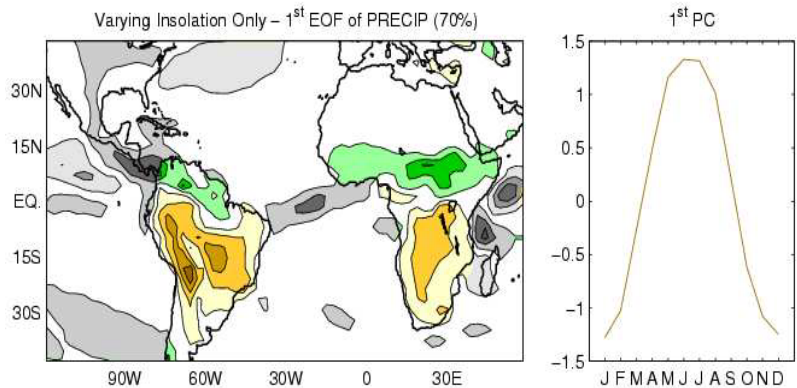
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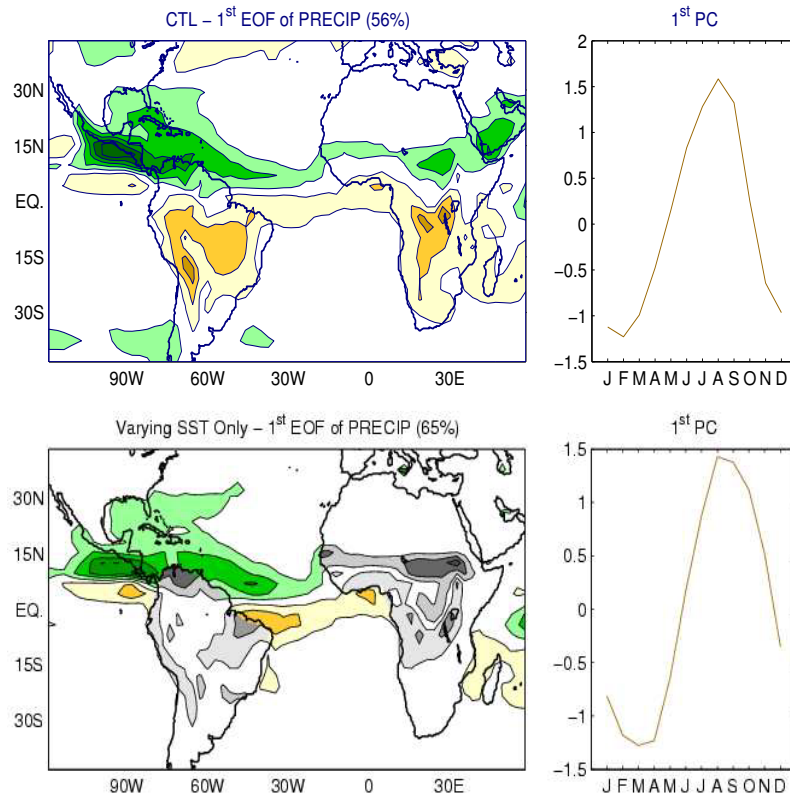
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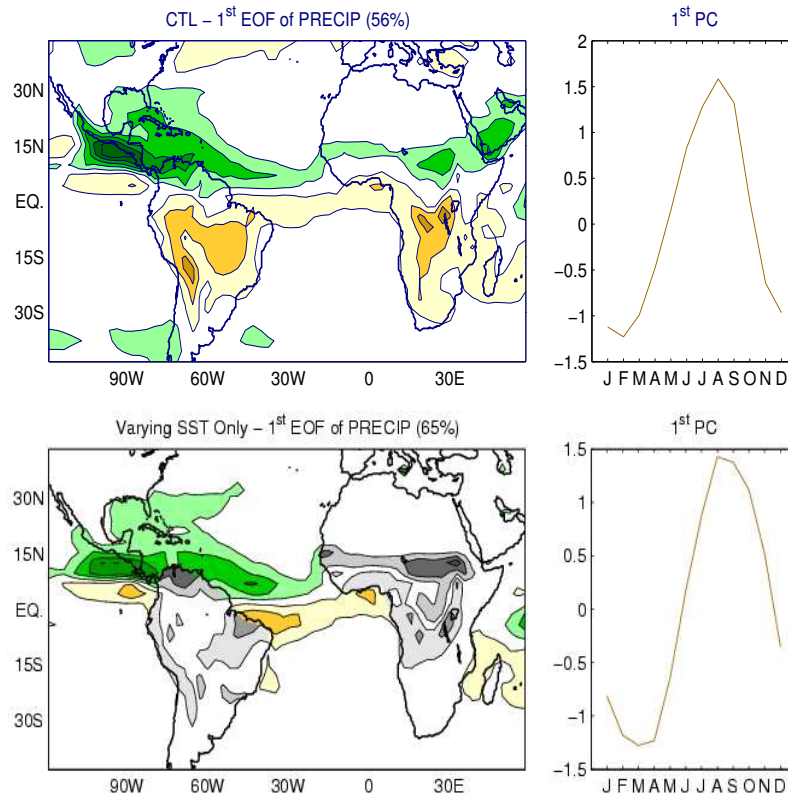
Extrema: March/August

Local control of ITCZ position

Non-local control of ITCZ intensity

What's the role of elevated condensational heating (Q_{cond}) in generating surface winds and convergence?

Results: LOCAL OCEAN



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North-South displacement over land and ocean

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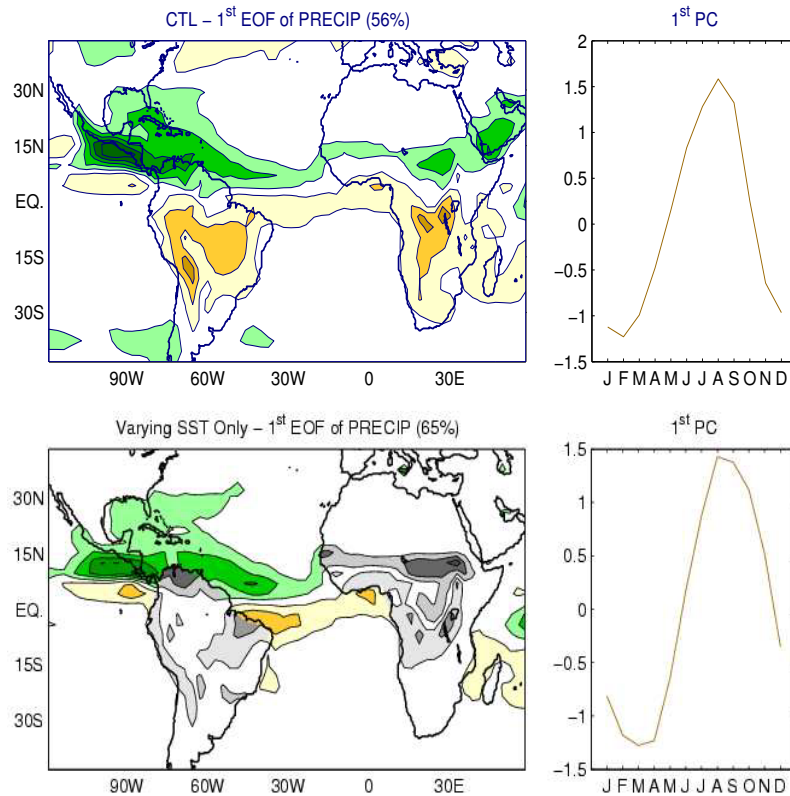
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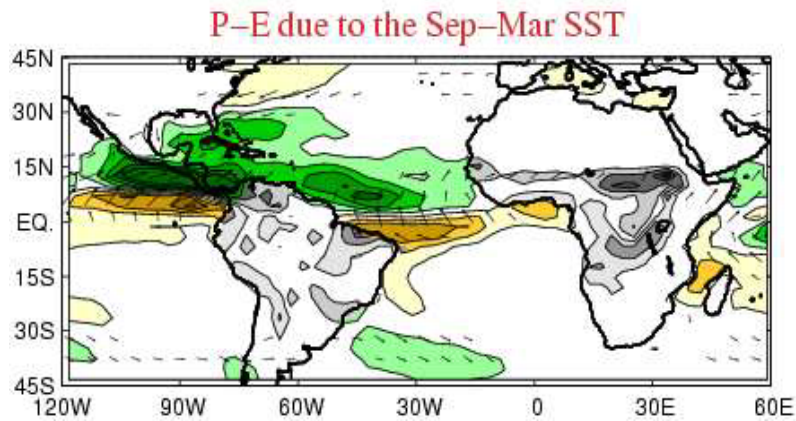
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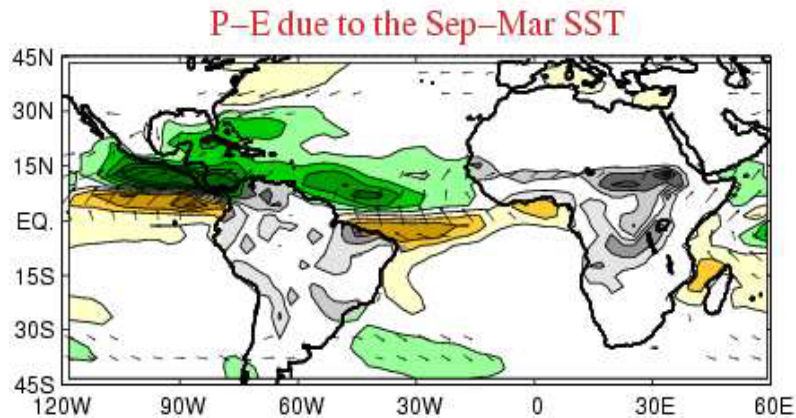
What's the role of elevated condensational heating (Q_{cond}) in generating surface winds and convergence?

Results: **LOCAL OCEAN**: the role of Q_{cond} .

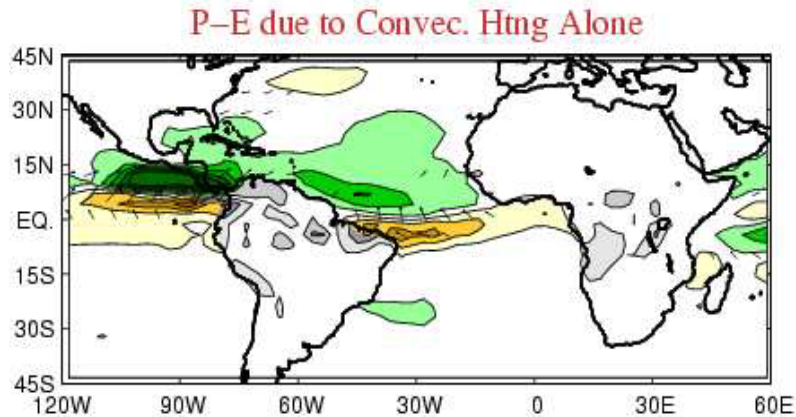


Moisture Convergence (P-E) &
Surface Wind Response to SST
changes.

Results: LOCAL OCEAN: the role of Q_{cond} .



Moisture Convergence (P-E) & Surface Wind Response to SST changes.



Moisture Convergence (P-E) & Surface Wind Response to Q_{cond} changes in the ITCZ. The ITCZ Q_{cond} drives a circulation that sustains the original SST-induced displacement of the ITCZ.

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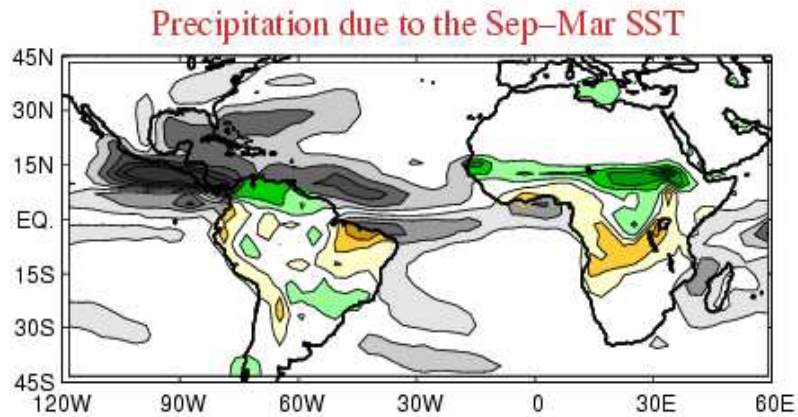
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Results: REMOTE LAND.

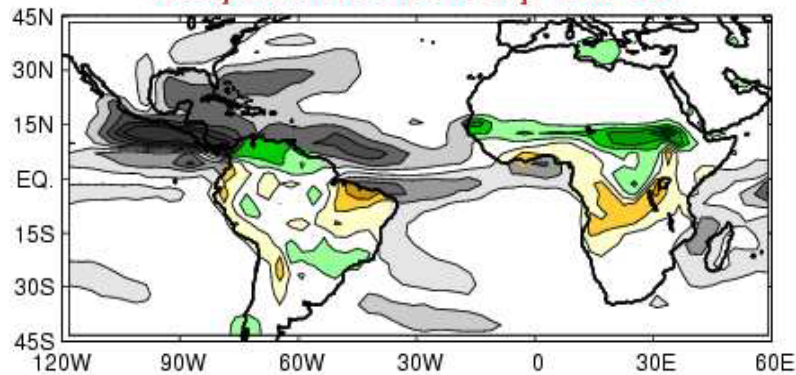


Precipitation response to SST changes.

- Northeast Brazil /Guiana
- Equatorial Africa & Guinea /Sahel

Results: REMOTE LAND: the role of Q_{cond} .

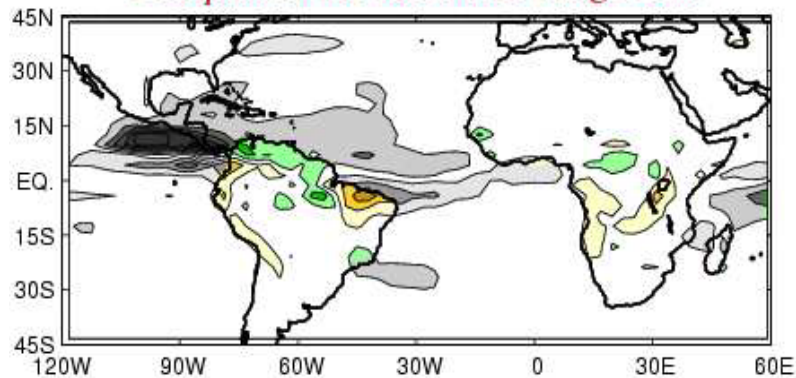
Precipitation due to the Sep–Mar SST



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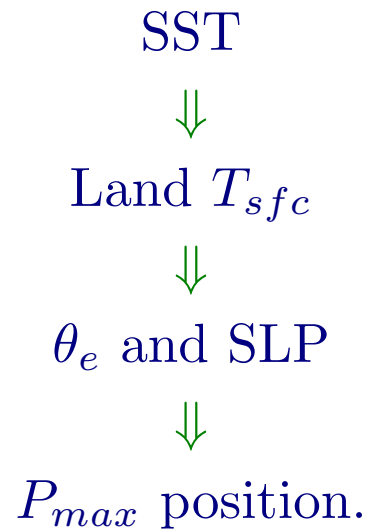
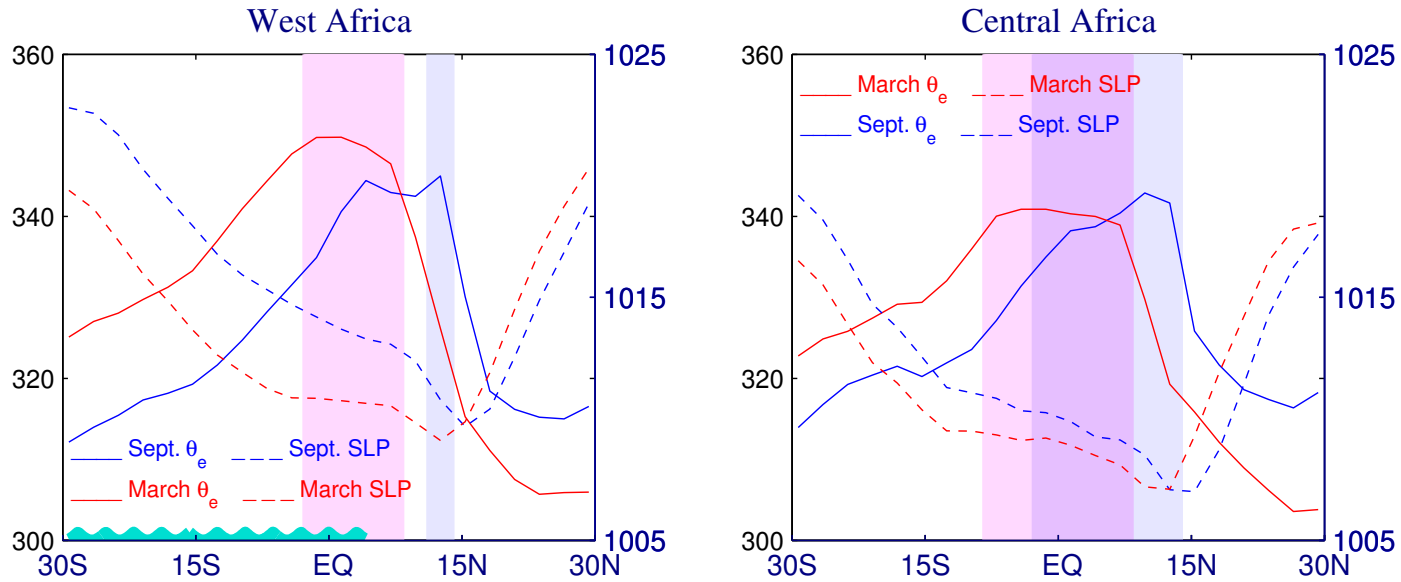
Precipitation due to Convec. Htng Alone



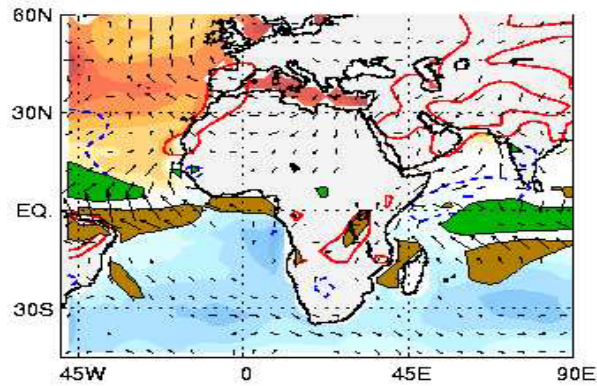
Precipitation response to Q_{cond} changes in the ITCZ.

- YES: South America
- NO: Africa

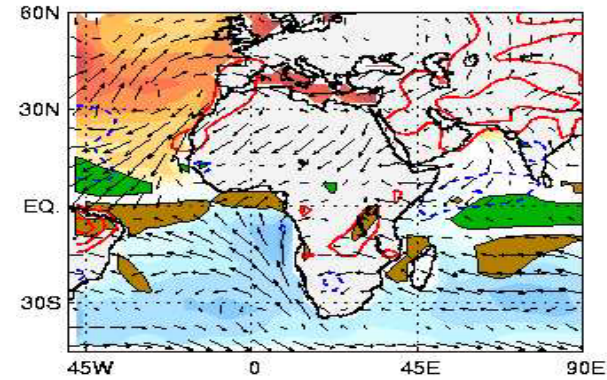
Results: **REMOTE LAND**: the role of T_{sfc}



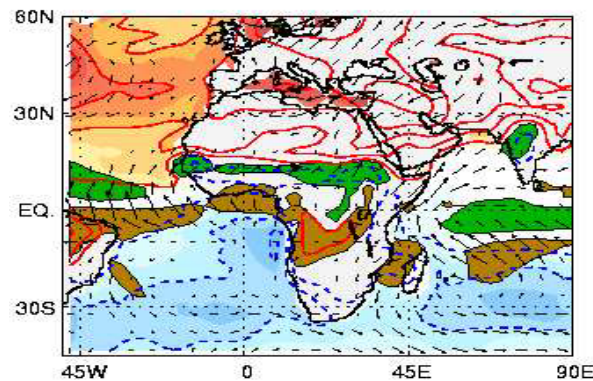
Results: REMOTE LAND: Inland Advection of SST.



+



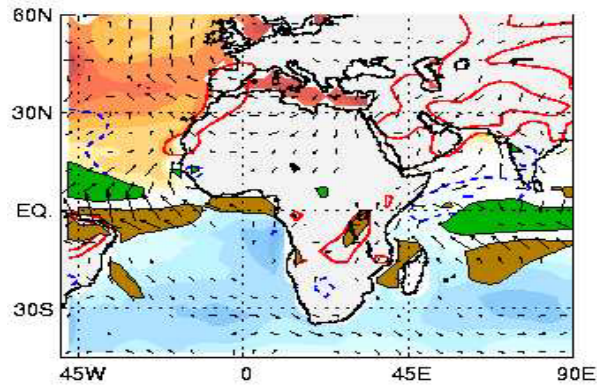
- SST ⇒ ITCZ
- ITCZ ⇒ wind in Africa + wind and SAT in Near East
- mean westerlies ⇒ Arabian SAT
- mean easterlies + wind anomalies ⇒ Sahara SAT



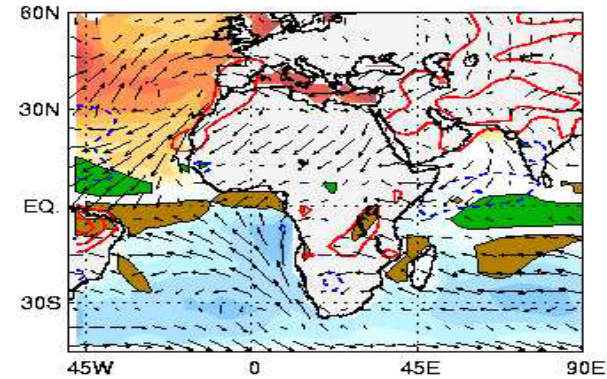
- African SLP ⇒ southerly wind
- ⇒ moisture convergence into the

Sahel

Results: REMOTE LAND: Inland Advection of SST.

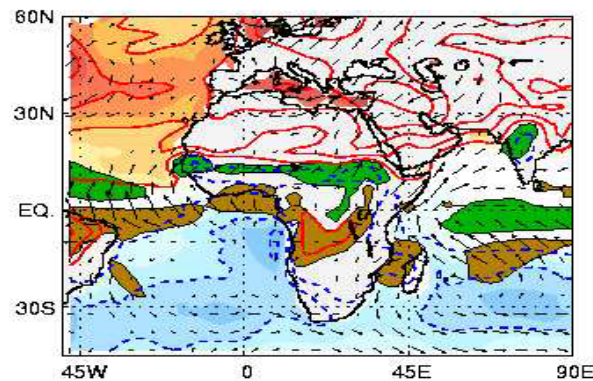


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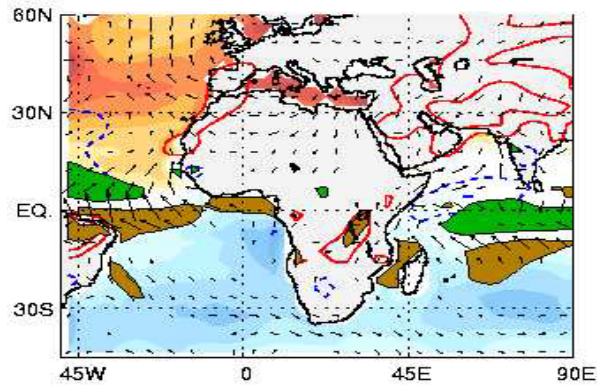
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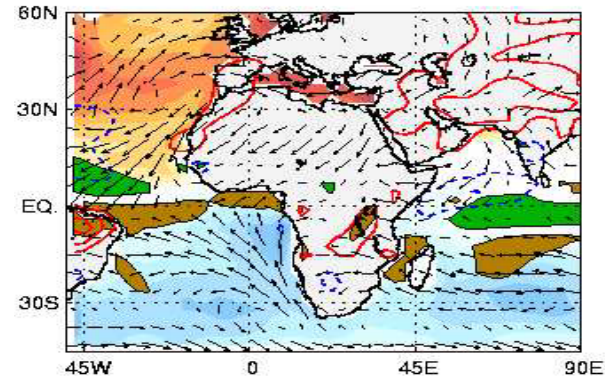
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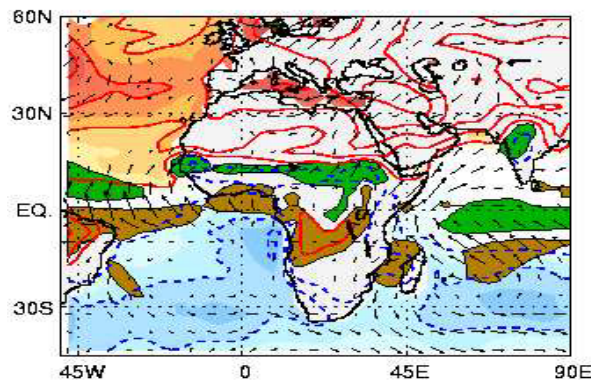
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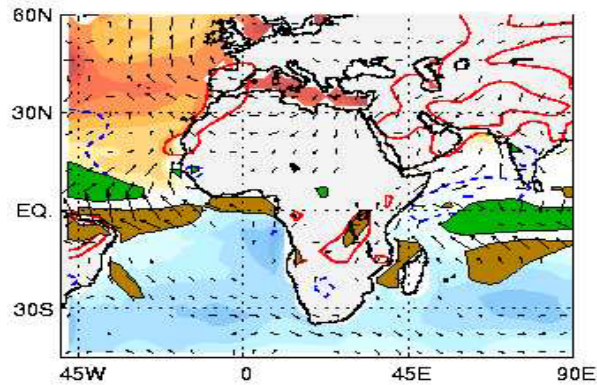
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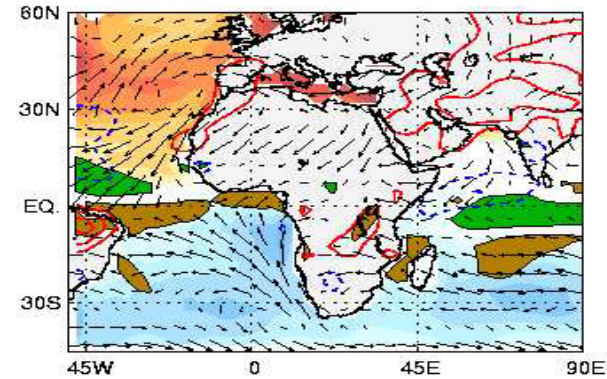
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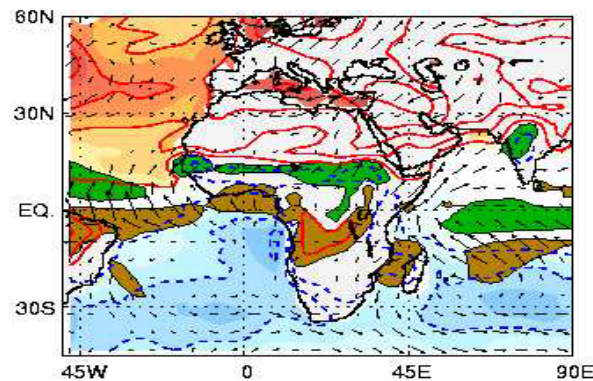
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RESULTS

A. LOCAL LAND

- AC of insolation;
March SST
⇒ AC of land precip

B. LOCAL OCEAN

- AC of SST;
March Insolation
⇒ AC of ocean precip

C. REMOTE LAND

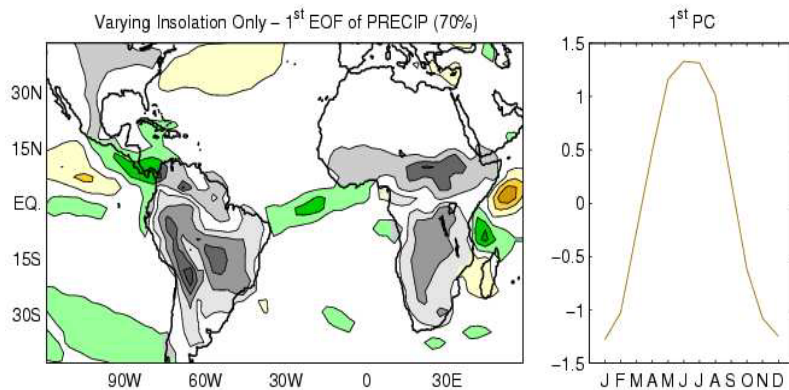
- AC of SST;
March Insolation
⇒ AC of land precip

D. REMOTE OCEAN

- AC of land insolation;
March SST
⇒ AC of ocean

Results: **REMOTE OCEAN:** ITCZ Intensity Change

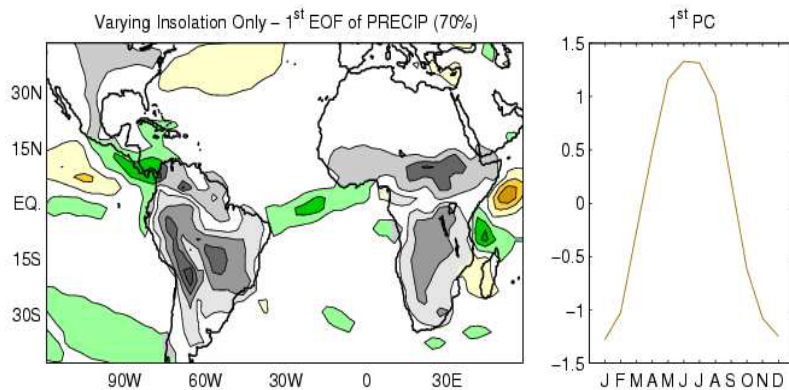
Uncoupled response: Land climate forces rainfall intensity anomalies in the Atlantic ITCZ.



- Does the ITCZ respond to changes in land surface temperature?
- Does it respond to changes in land precipitation (Q_{cond})?

Results: REMOTE OCEAN: ITCZ Intensity Change

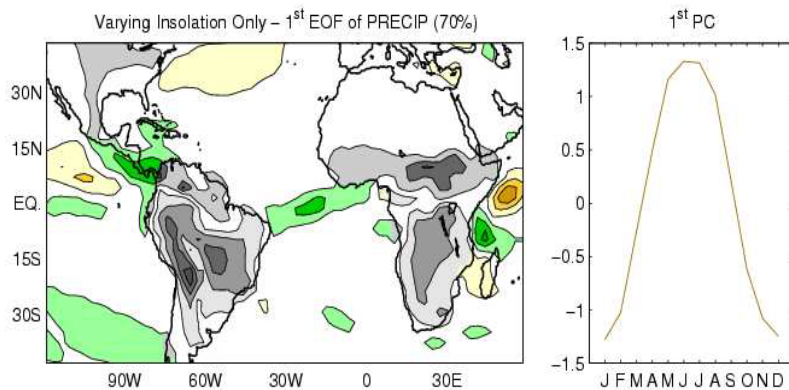
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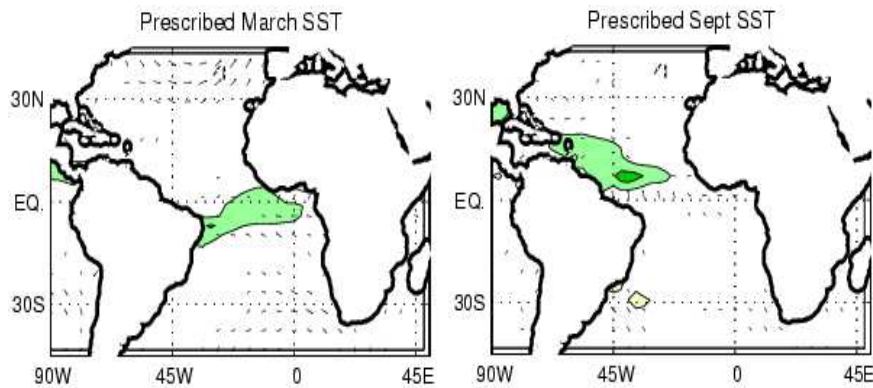
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Results: **REMOTE OCEAN:** Response to Land Q_{cond}

The response to imposed steady forcing in elevated condensational heating over Africa and South America.

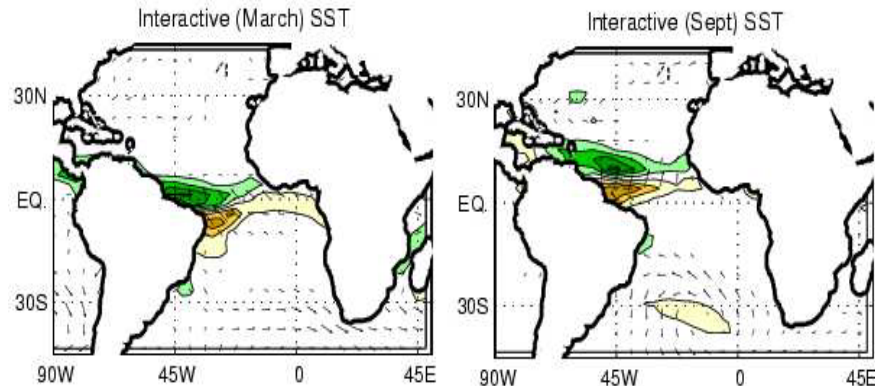
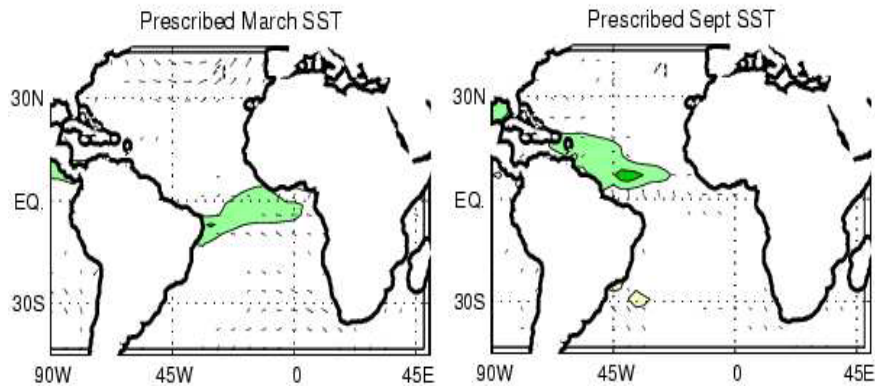


Uncoupled response

- ITCZ intensity responds to remote Q_{cond} anomalies.
- Anomalies are co-located with the mean ITCZ.
- Surface wind anomalies over ocean.

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The response to imposed steady forcing in elevated condensational heating (Q_{cond}) over Africa and South America.



Uncoupled response

- ITCZ intensity responds to remote Q_{cond} anomalies.
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Coupled response

- wind anomalies \Rightarrow evaporation anomalies
- latent heat flux anomalies \Rightarrow SST anomalies
- anomalous SST gradient \Rightarrow the ITCZ shift
- wind/evaporation/SST/ITCZ feedback.

COUPLED RESULTS

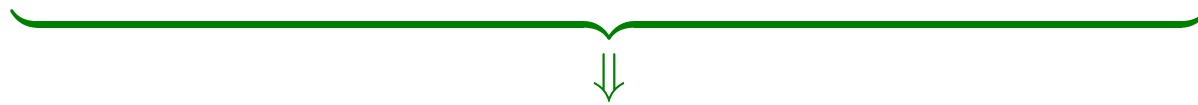
D. REMOTE OCEAN (coupled!)

CTL

- AC of Insolation over Land.
- AC of Insolation over Ocean.
- AC of Q-flux.

Fixed Insolation over Land

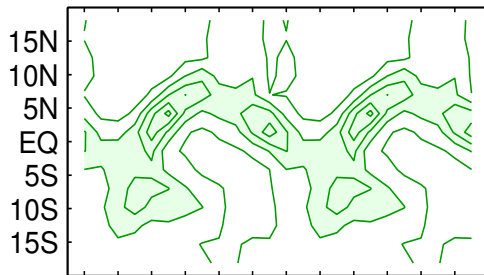
- March Insolation over Land.
- AC of Insolation over Ocean.
- AC of Q-flux.



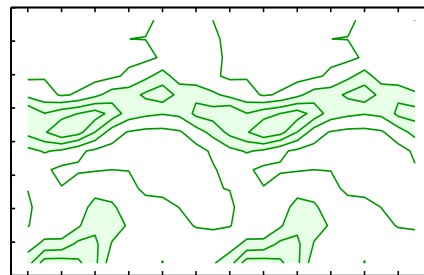
Effect of AC of Insolation over Land

Results: **REMOTE OCEAN**: Central Atlantic response to AC of Insolation over Land

Control
PRECIPITATION



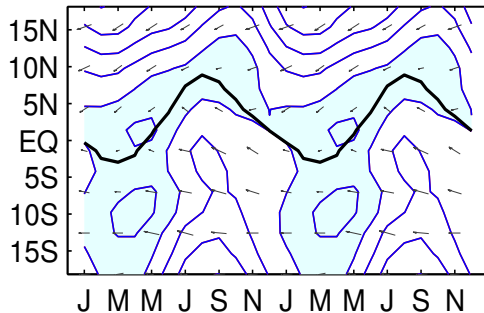
Oceanic Forcings Only
PRECIPITATION



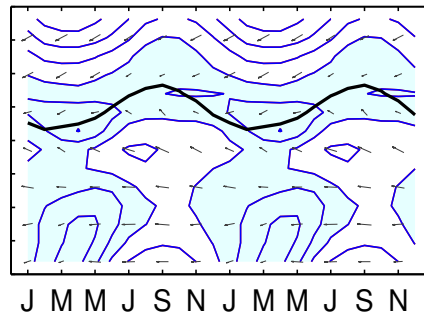
suppressed AC over land \Rightarrow suppressed meridional annual march of ITCZ.

Concomitant changes in SST.

SST & V



SST & V

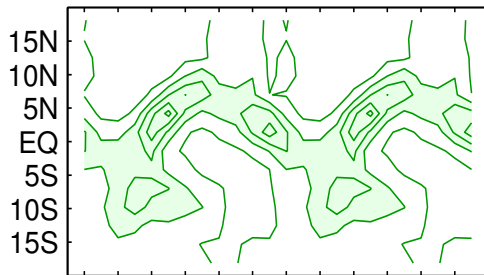


What forces the SST & ITCZ anomalies?

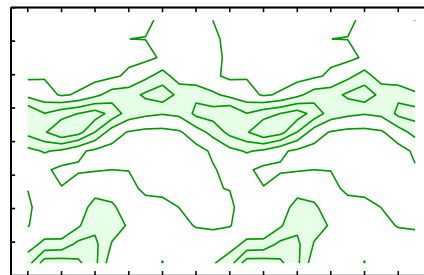
- Land surface temperature?
- Land precipitation?

Results: **REMOTE OCEAN**: Central Atlantic response to
AC of Insolation over Land

Control
PRECIPITATION



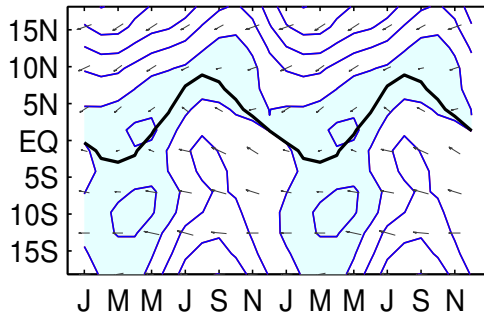
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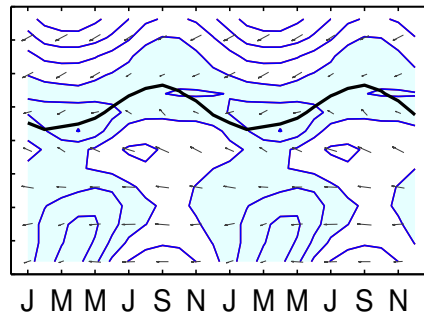
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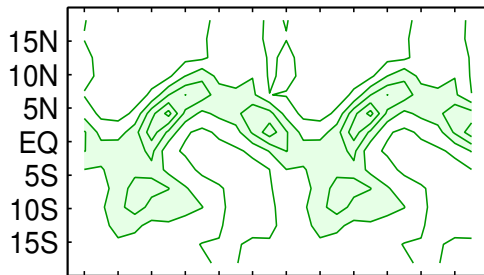


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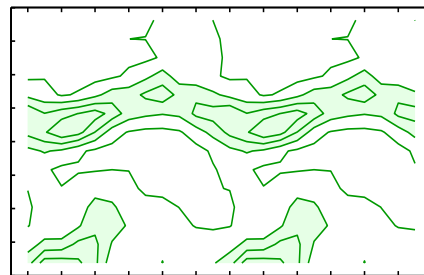
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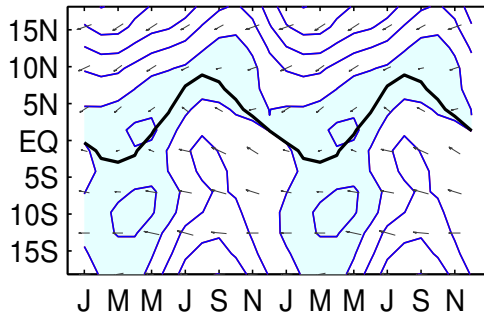
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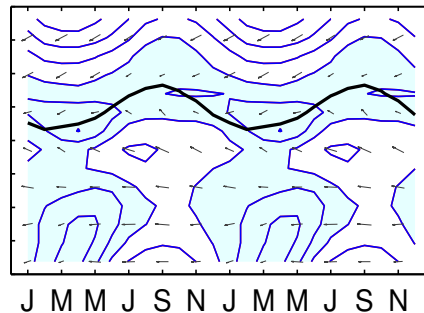
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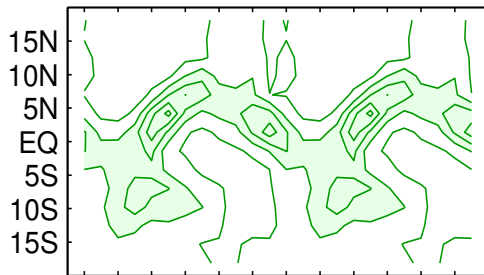


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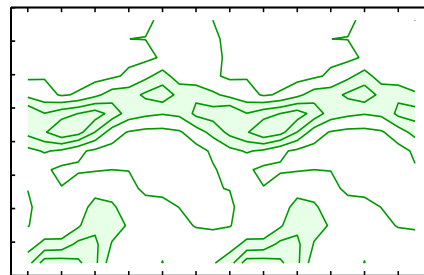
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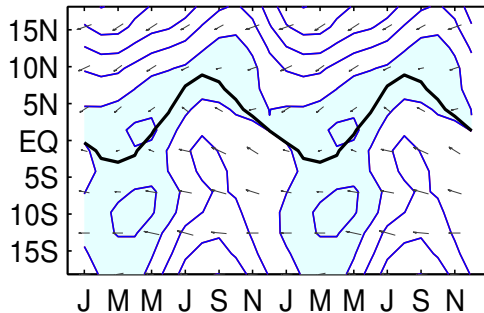
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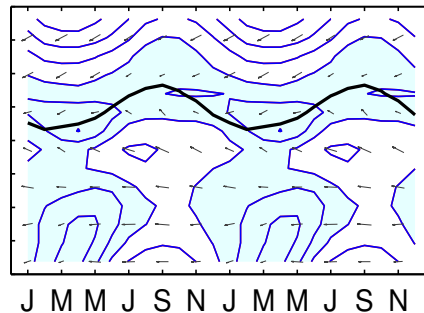
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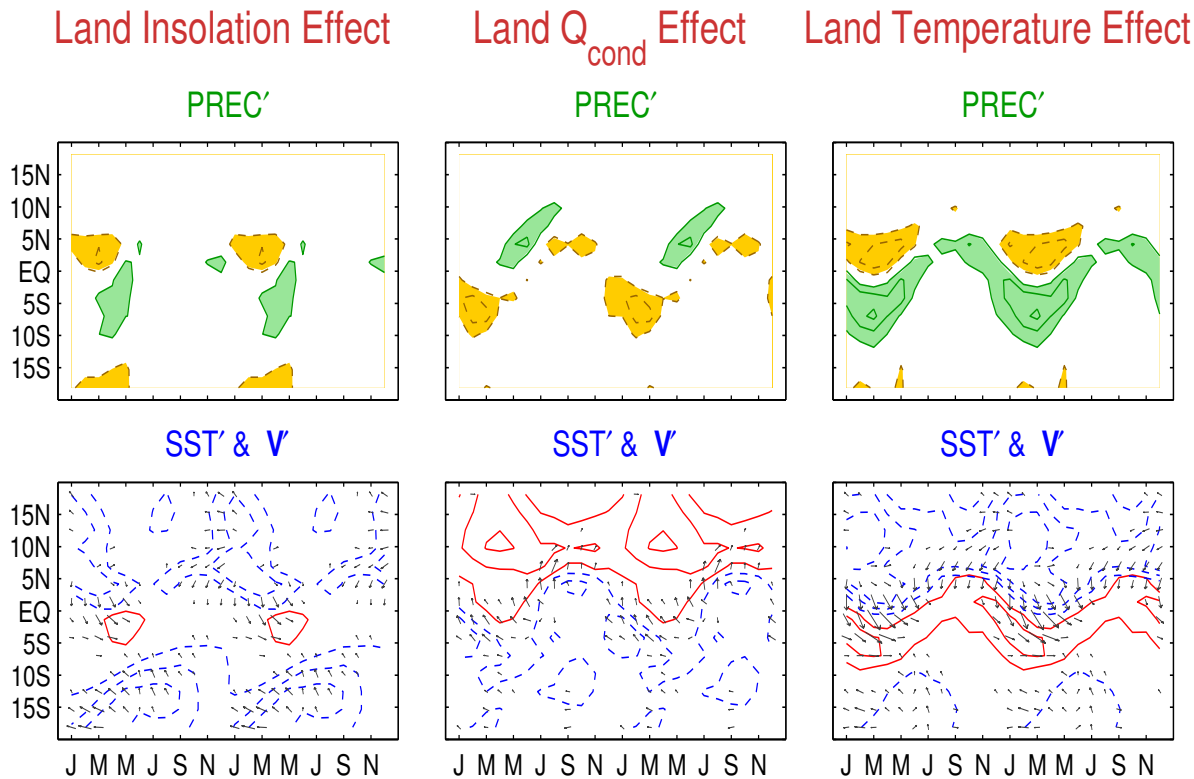


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Results: **REMOTE OCEAN**: Effect of Land Temperature and Precipitation

Compare the control simulation to the simulation with fixed insolation over land and the simulation with fixed elevated condensational heating over Africa and South America (Q_{cond})

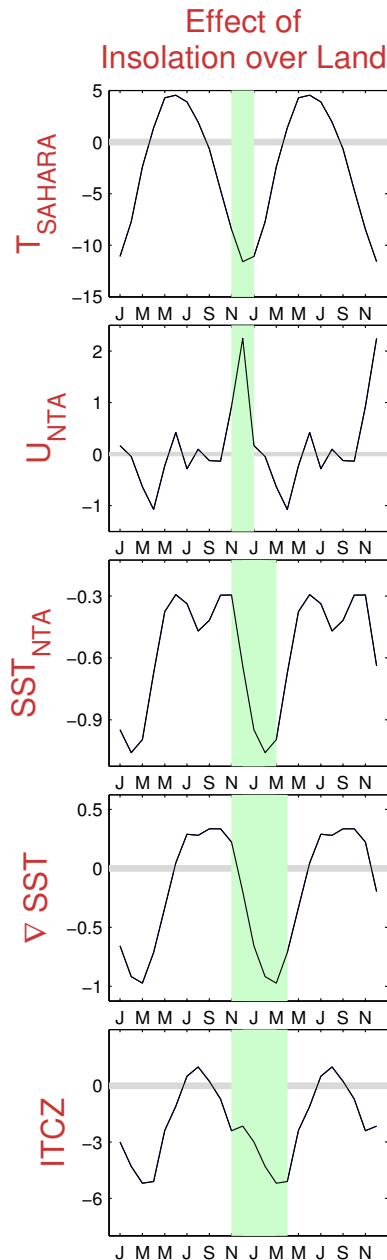


Land Q_{cond} and land T_{sfc} have large and opposite effects.

Land T_{sfc} dominant.

How?

Results: **REMOTE OCEAN:** Development of ITCZ Anomalies in Response to Land Forcing.



December cold “anomalies” in the Sahara \Rightarrow thermal high



stronger NTA Trades \Rightarrow stronger evaporation



colder NTA



feedbacks @ equator

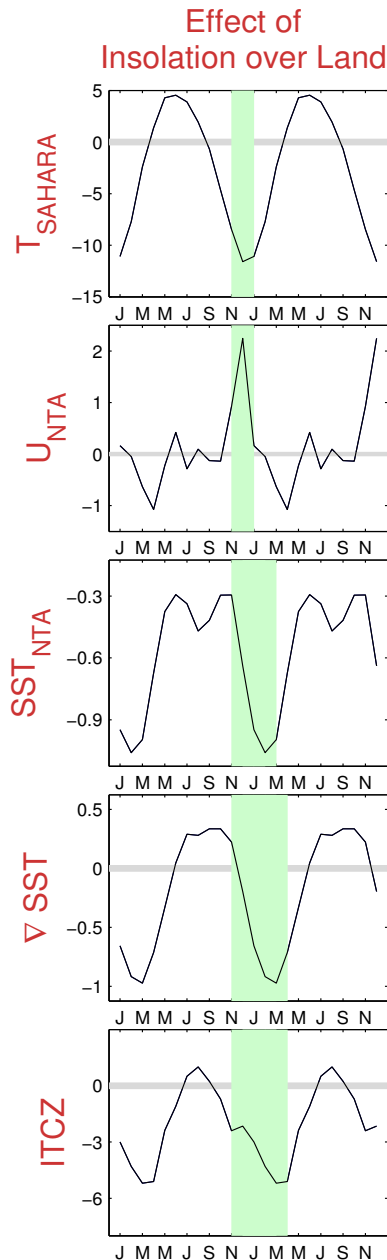


equatorial SST gradient



ITCZ shift

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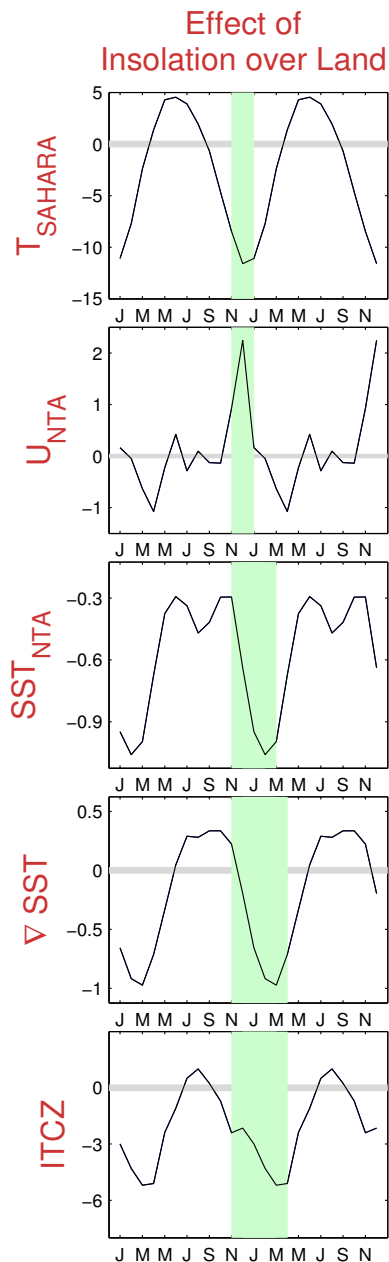


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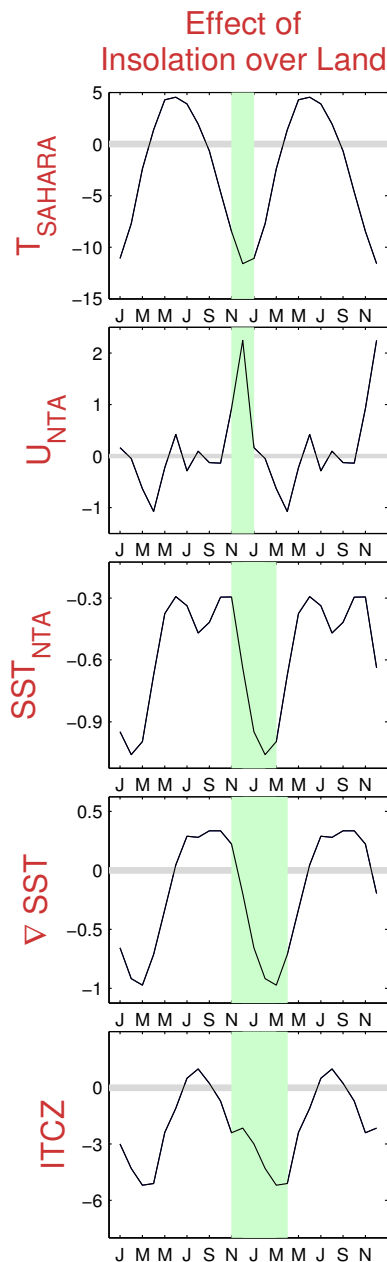


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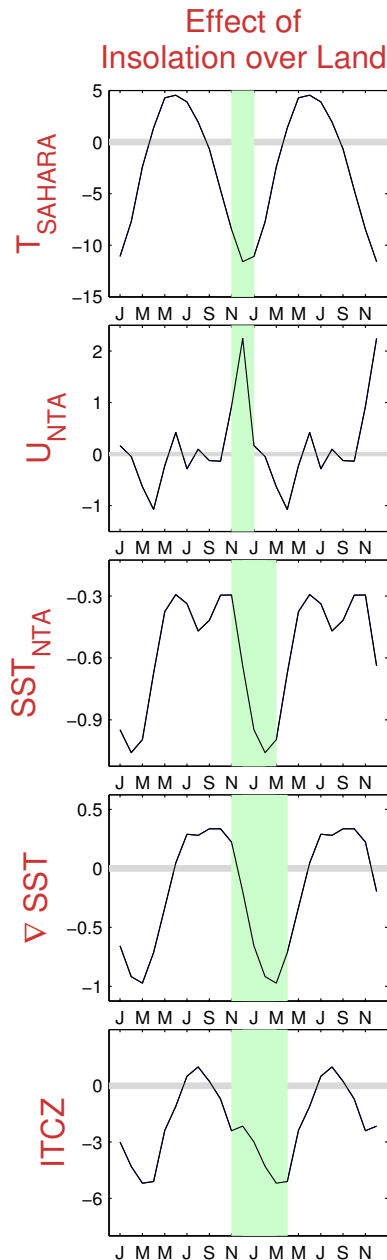


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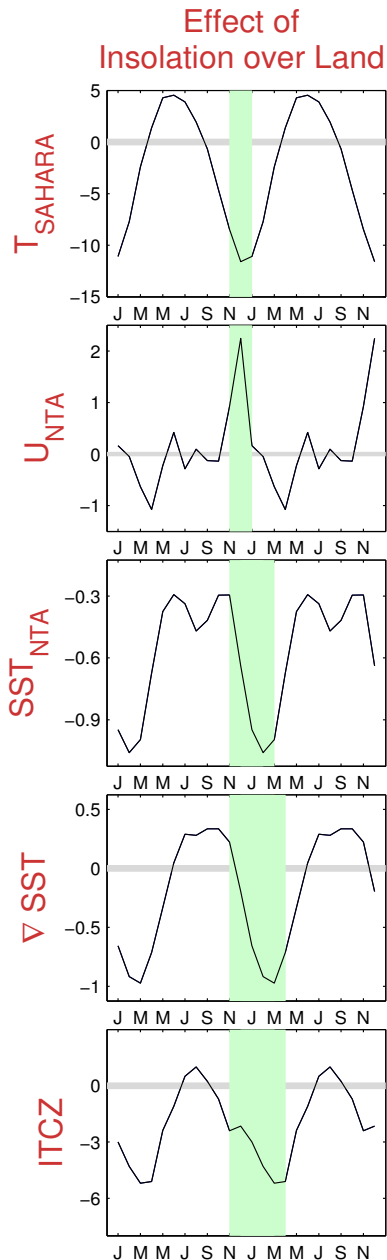


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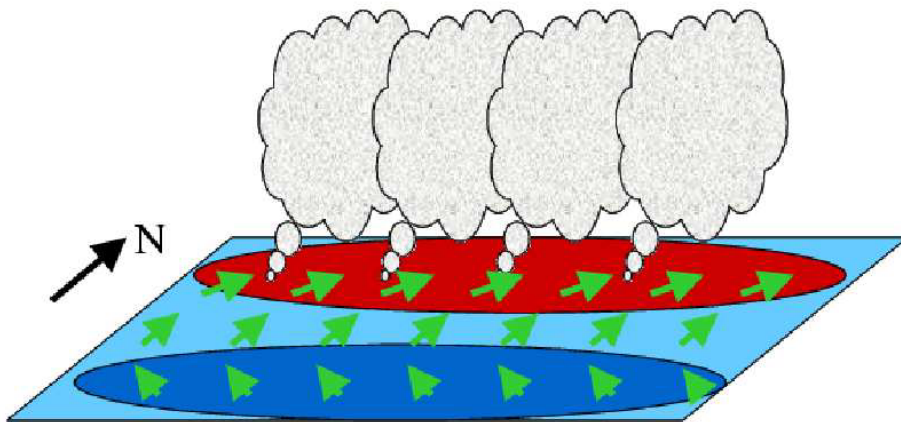
equatorial SST gradient



ITCZ shift

Equatorial Feedbacks

- SST gradient \implies cross-equatorial flow
 $\leftarrow \implies ?$



Feedback on SST gradient

- Cld-rad: -ve
- Winds : +ve
- (• OHT : -ve?)

John Chiang

CONCLUSIONS:
Mutual Influences of Land and Ocean

Conclusions: **REMOTE INFLUENCE OF SST ON LAND**

The AC of SST influences the AC of land precipitation

in coastal areas: SST \Rightarrow ITCZ \Rightarrow Northeast Brazil and Guiana precipitation.

in the Sahel: SST \Rightarrow Sahara T_{sfc} \Rightarrow SLP gradient \Rightarrow Sahel precipitation.

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Land climate influences the ITCZ intensity and position.

intensity: Land $Q_{cond} \Rightarrow$ free tropospheric temperature \Rightarrow stability over the ocean \Rightarrow ITCZ intensity.

position: directly and indirectly forced wind \Rightarrow latent heat loss \Rightarrow SST \Rightarrow ITCZ position.

- Sahara $T_{sfc} \Rightarrow$ north tropical Atlantic Trades.
- Land $Q_{cond} \Rightarrow$ equatorial wind.
- Land $Q_{cond} \Rightarrow$ ITCZ intensity \Rightarrow surface wind at the edge of the ITCZ.

Changes in surface wind trigger coupled feedbacks among wind, SST, and ITCZ.

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IMPLICATIONS:
for models, TAV, and paleo

Implications: MODELING THE ANNUAL CYCLE

- CCM3 biases: reduce the overestimate of precipitation over land (by changing albedo?) \Rightarrow solve the underestimation of precipitation in the ITCZ.
- CGCM biases: correct precipitation over equatorial coastal areas and Sahel \Leftrightarrow correct SST.
- CGCM biases: correct march of Atlantic ITCZ \Leftrightarrow correct AC of temperature in the Sahara and of precipitation in Africa and South America.

Implications: TROPICAL ATLANTIC VARIABILITY & CLIMATE CHANGE

- SST influences land precipitation at the annual timescale in the same way it does at the interannual timescale \Rightarrow the AC is indeed a useful analogue for TAV.
- AC of land influences AC of ocean $\stackrel{?}{\Rightarrow}$ continental variability influences maritime variability.
- Change in African climate (e.g. due to deforestation) $\stackrel{?}{\Rightarrow}$ Northeast Brazil (via ITCZ).

Remaining question: How does the response time of the ocean modify the response to continental forcing at the interannual timescale?

Implications: PALEO CLIMATE STUDIES e.g. the Green Sahara

- Simulation of precipitation at edge of Sahara \Leftrightarrow correct basinwide Atlantic SST boundary conditions
- Simulation of precipitation in Sahara \Leftrightarrow correct simulation of soil albedo (soil moisture and vegetation).
- The greening of the Sahara should be visible in paleo records of the ITCZ position (e.g. from the Cariaco basin). Is it?

Caveat: what's the role of ocean dynamics?