

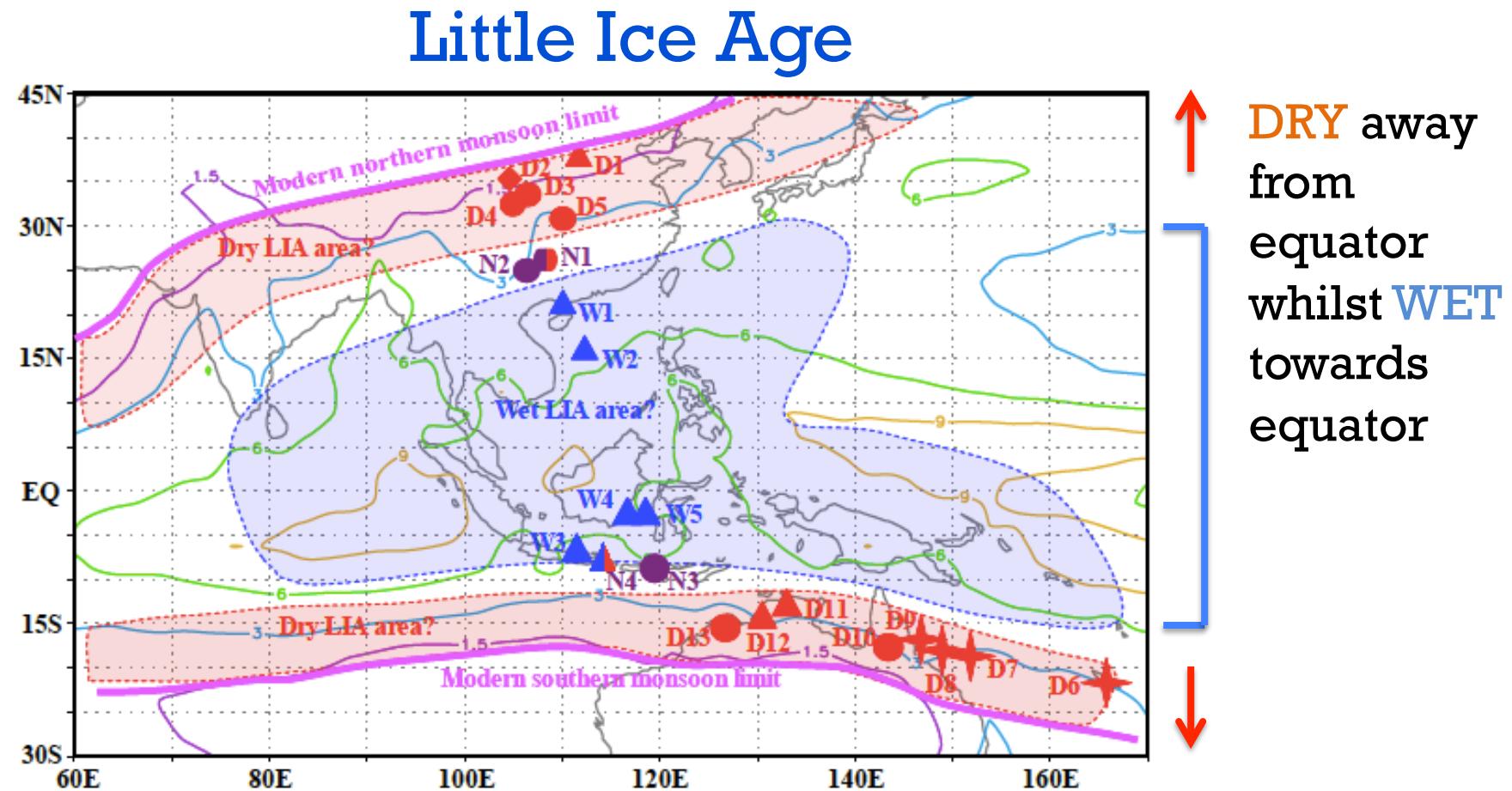
Australasian Monsoon Variability During the Common Era Inferred from Indo-Pacific Speleothem Records

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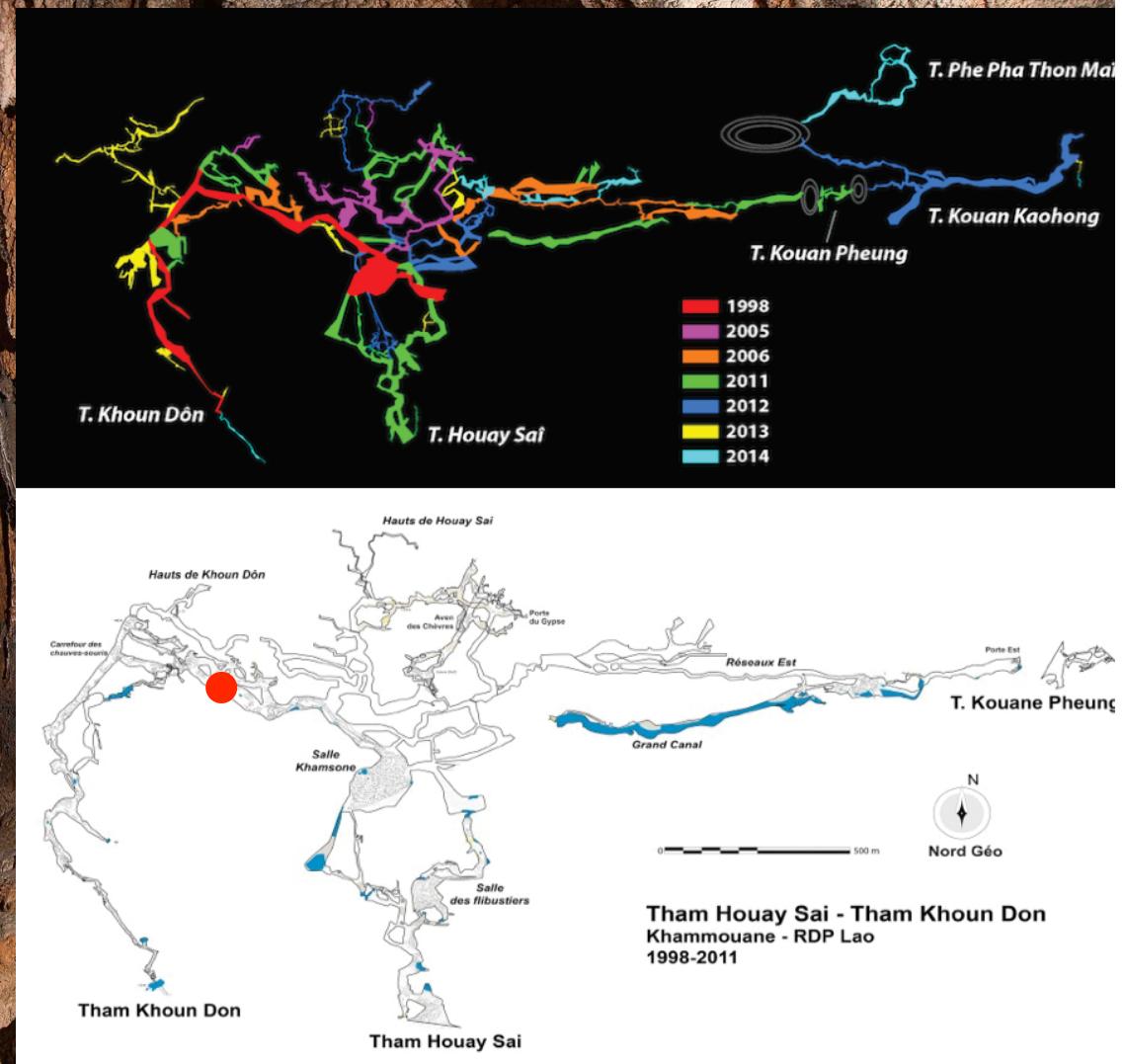
WILLIAM
PATERSON
UNIVERSITY



Hydroclimate records from the Australasian monsoon region for the Common Era



Tham Khoun Don-Houay Sai, Laos



Map: Explor-Laos

Photo: Serge (speleomag.com)

EXPLO-LAOS, Khammouane Province, 2015



Speleothems: “Fossilized Groundwater”

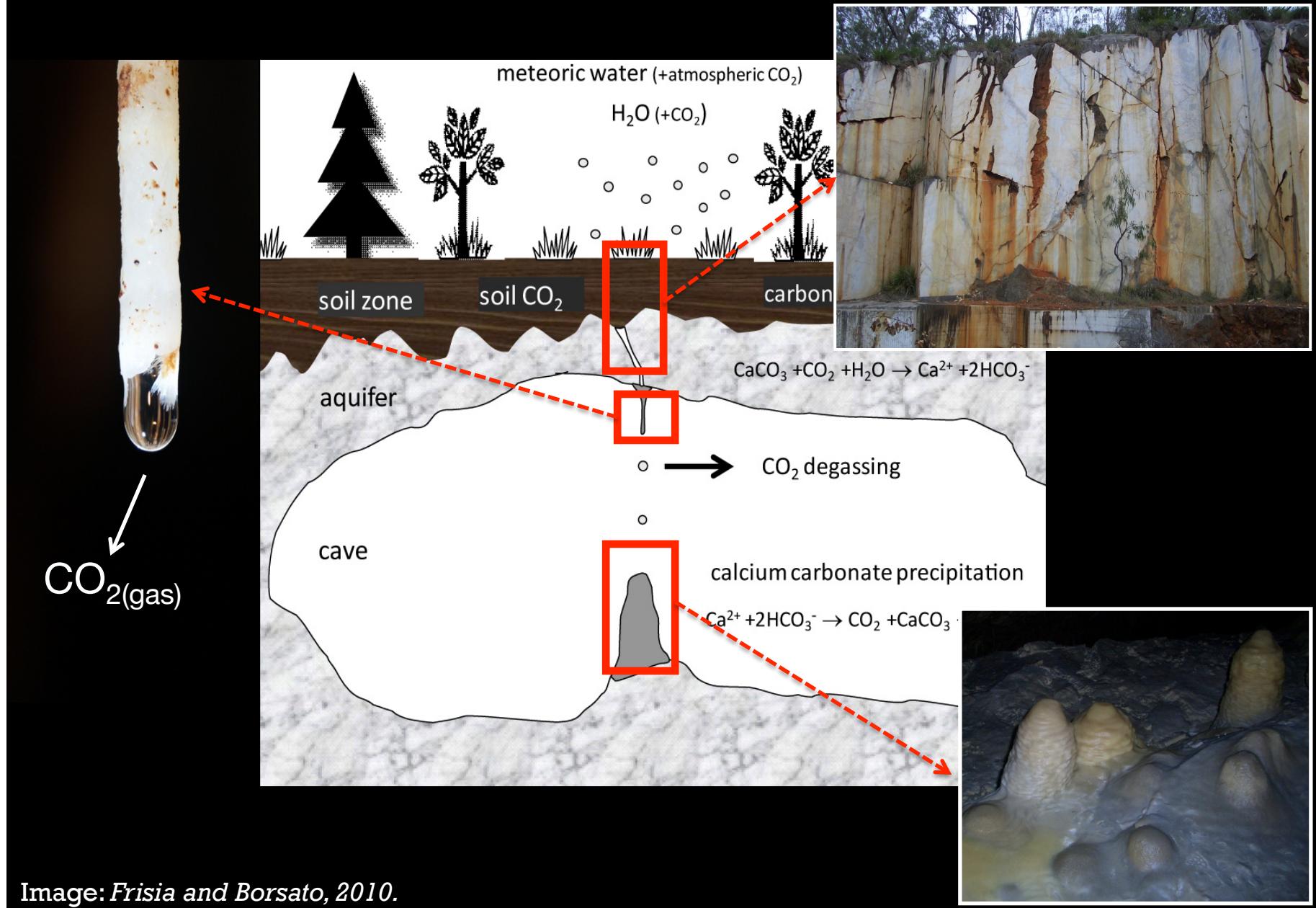
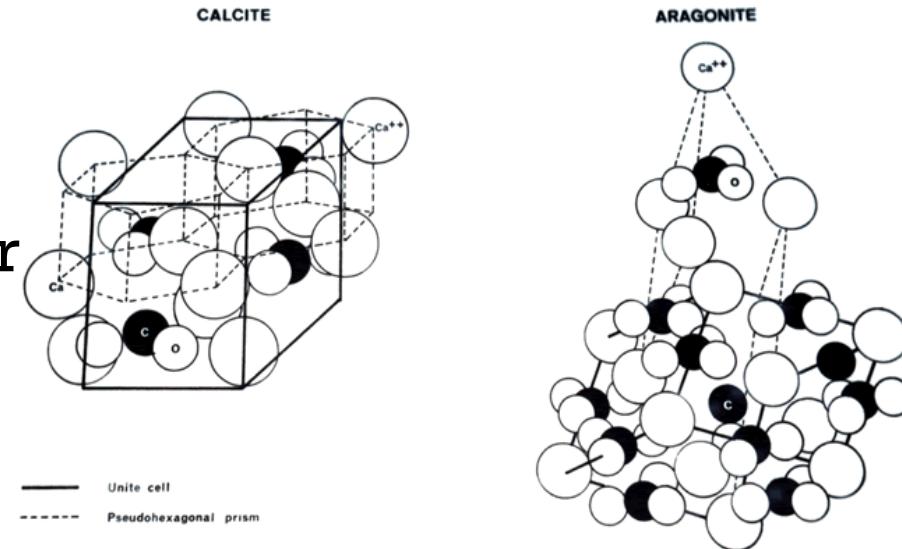


Image: Frisia and Borsato, 2010.

Reconstructing past climate using carbonate mineral deposits

CaCO_3 can be either calcite or aragonite



Trace amounts of other elements, e.g. Mg, Sr, U can substitute for Ca in the crystal structures

CaCO_3

The carbon atom can be stable ^{12}C or ^{13}C or radioactive ^{14}C

The oxygen atoms can be ^{16}O , ^{17}O , or ^{18}O



Environmental effects on speleothem $\delta^{18}\text{O}$



$$\delta^{18}\text{O} = \left\{ \frac{\left(^{18}\text{O}/^{16}\text{O} \right)_{\text{Sam.}} - \left(^{18}\text{O}/^{16}\text{O} \right)_{\text{Std.}}}{\left(^{18}\text{O}/^{16}\text{O} \right)_{\text{Std.}}} \right\} \times 10^3$$

Assuming equilibrium deposition:

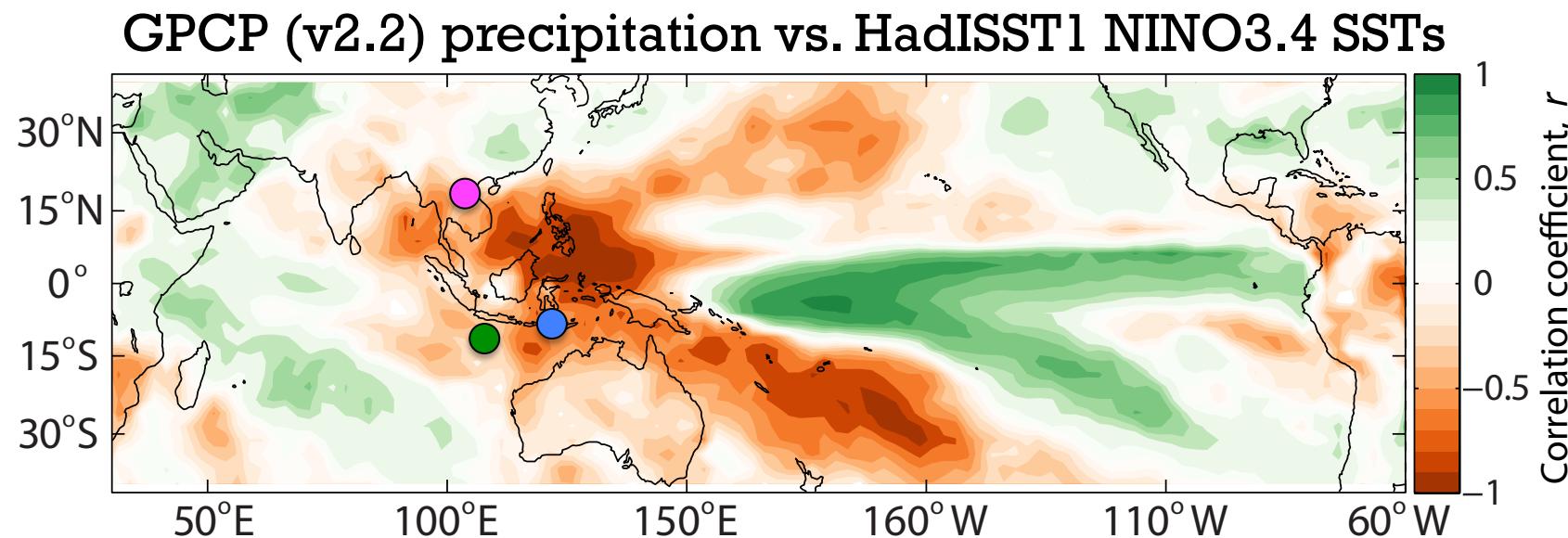
$$\delta^{18}\text{O}_{\text{calcite}} = f [\delta^{18}\text{O}_{\text{water}}, \text{Cave T (MAT)}]$$

$$\begin{array}{c} \uparrow \\ \text{d}\delta^{18}\text{O}/dT = -0.23 \text{ ‰ / } ^\circ\text{C} \\ \uparrow \end{array}$$

Accounts for most of variability

(precipitation $\delta^{18}\text{O}$ + hydrology + (fractionation/kinetics?))

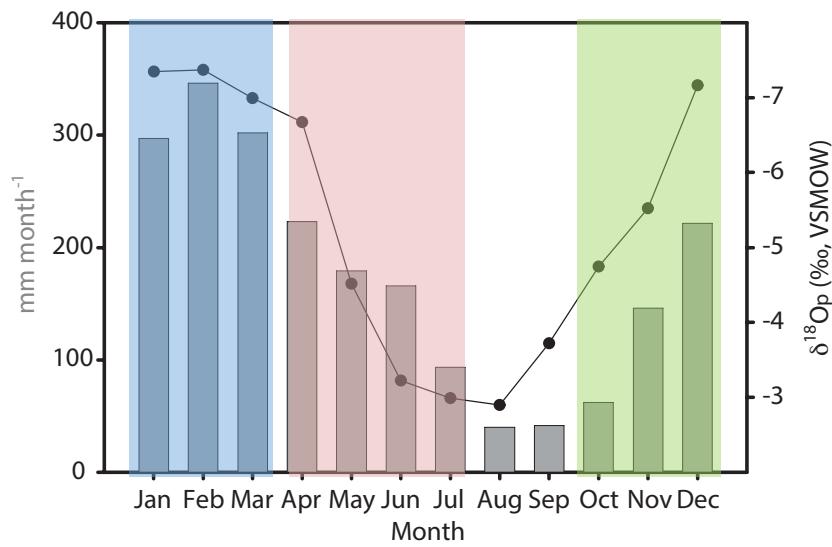
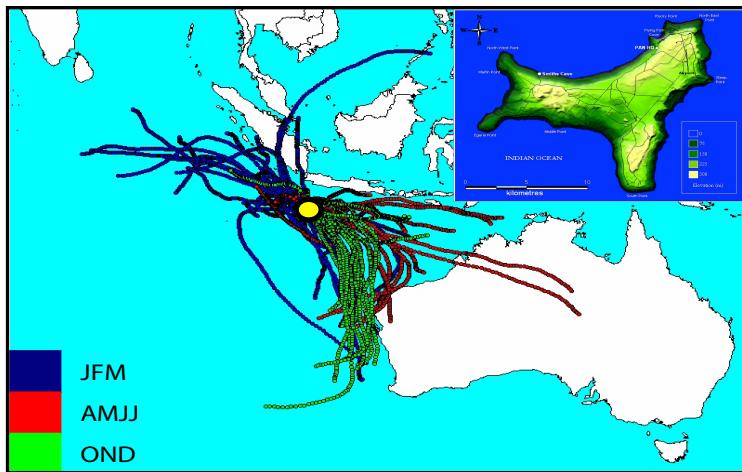
Indo-Pacific hydroclimate over the past 1-2 millennia



- Christmas Island (~1375-2003 C.E.)
- Flores (~0-2000 C.E.)
- Northern Laos (~0-2008 C.E.)

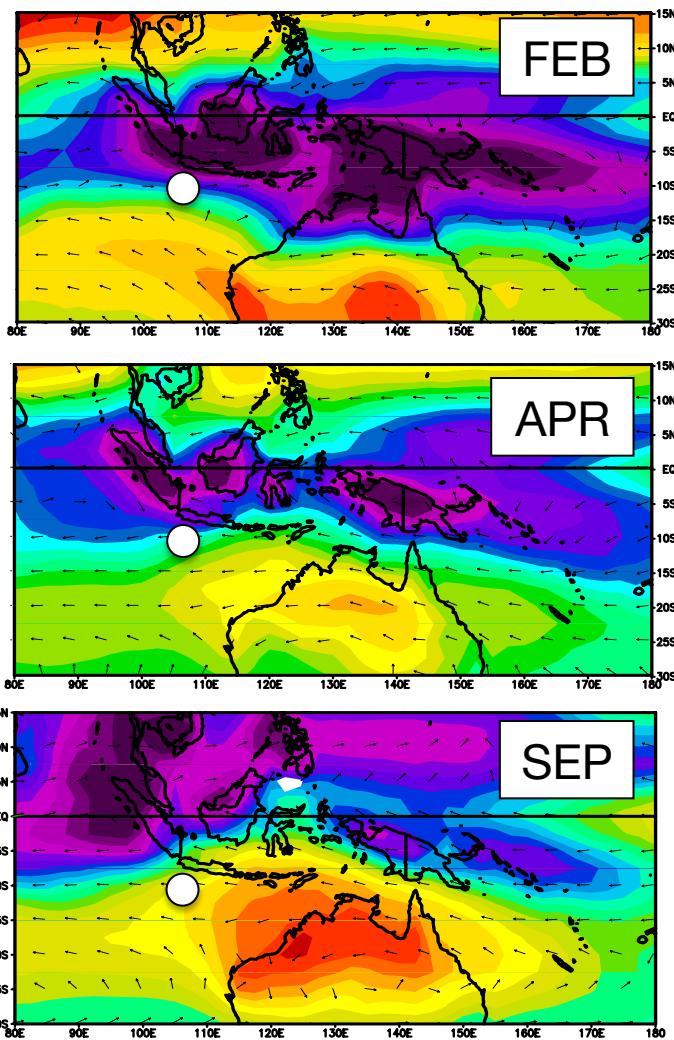
Christmas Island: climatology

HYSPLIT back trajectories (rain events > 90th percentile)



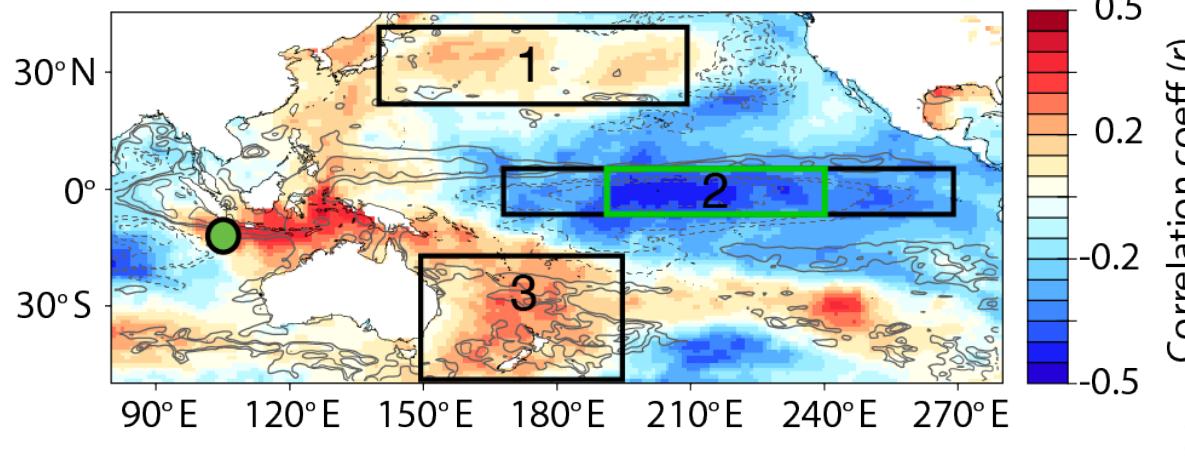
Source: BOM and IsoGSM

Outgoing Longwave Radiation (OLR)



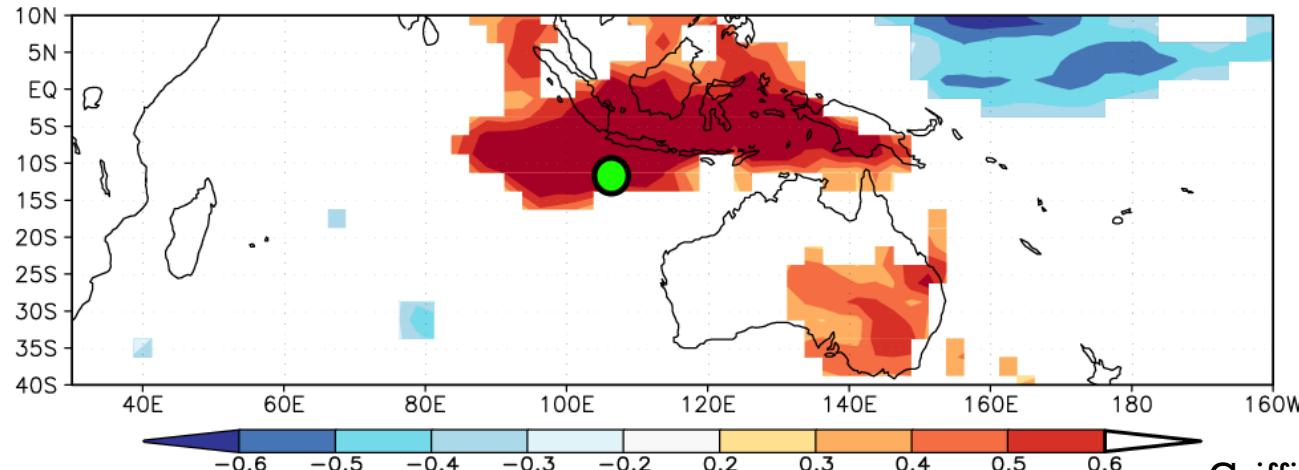
Instrumental precip: ENSO/IPO influence

Xmas rainfall (GHCN) vs. SST (shading)+sea surface height (contours)



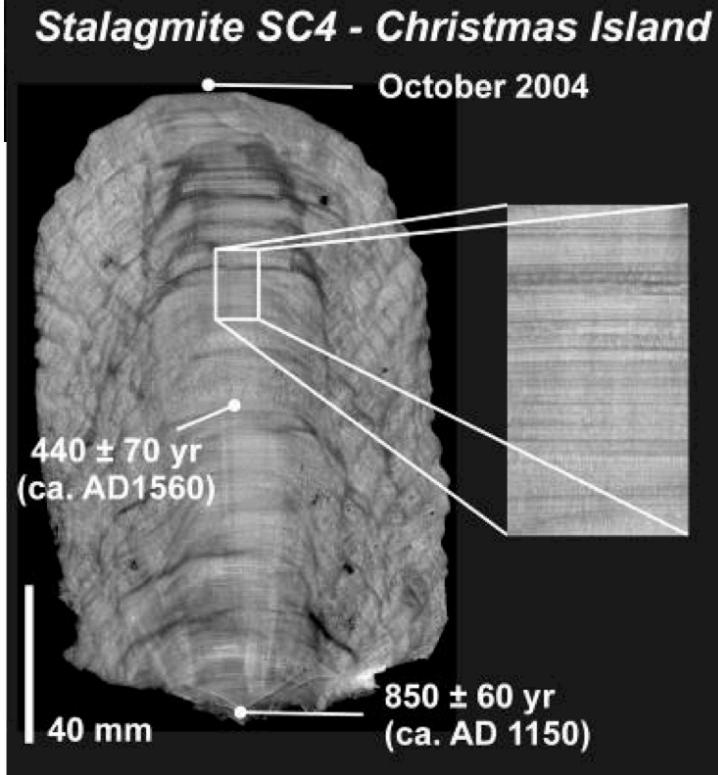
Henley et al., 2015.

Xmas rainfall (GHCN) vs. GPCP gridded rainfall (shading)+OLR (contours)
Xmas rainfall GPCP

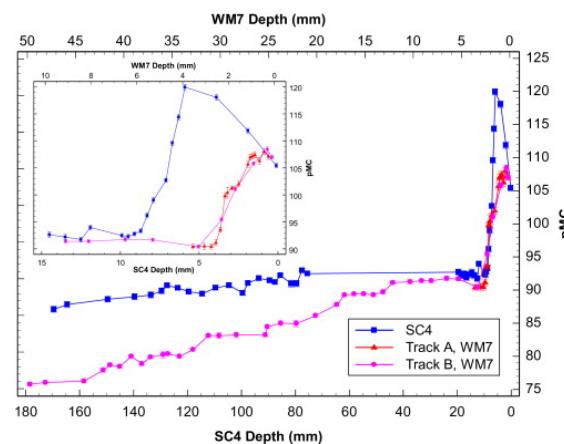


Griffiths et al., in prep.

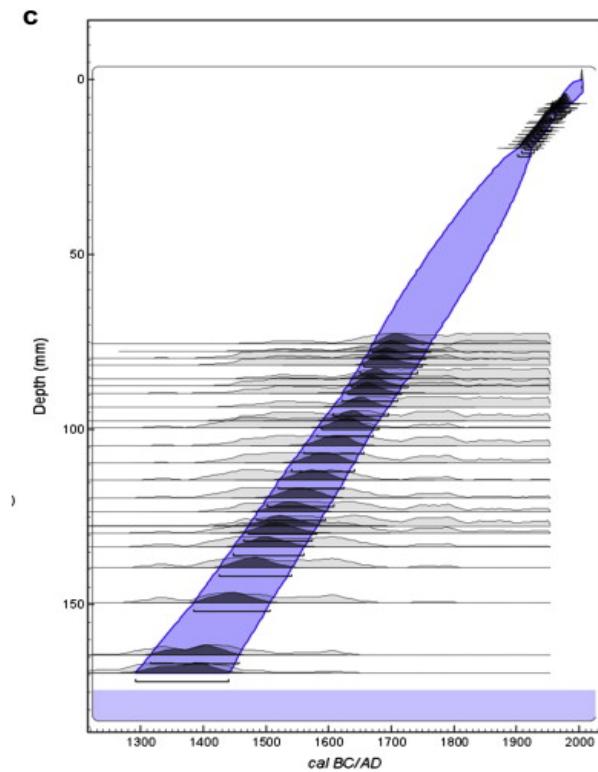
Speleothem: SC4 (Smith's Cave)



^{14}C chronology



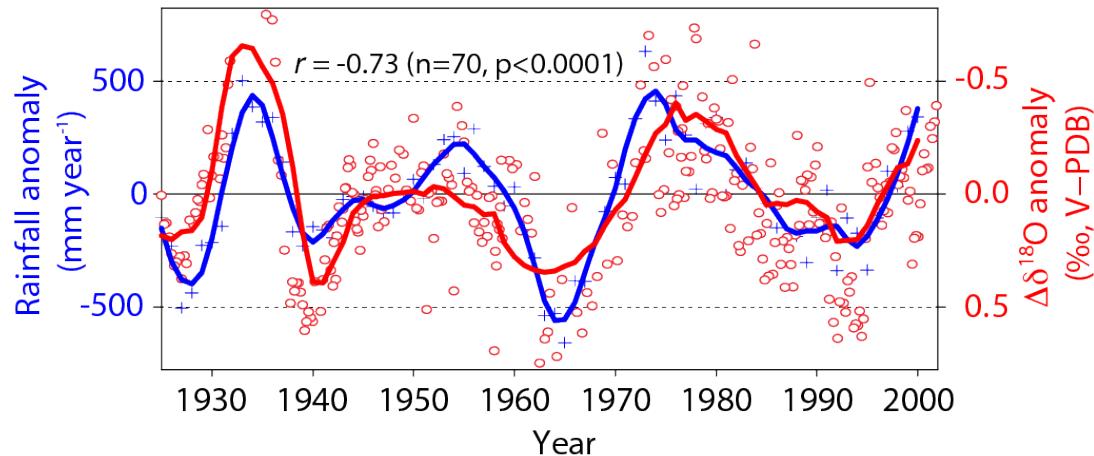
44 ^{14}C ages
used to
construct
Oxcal age
model



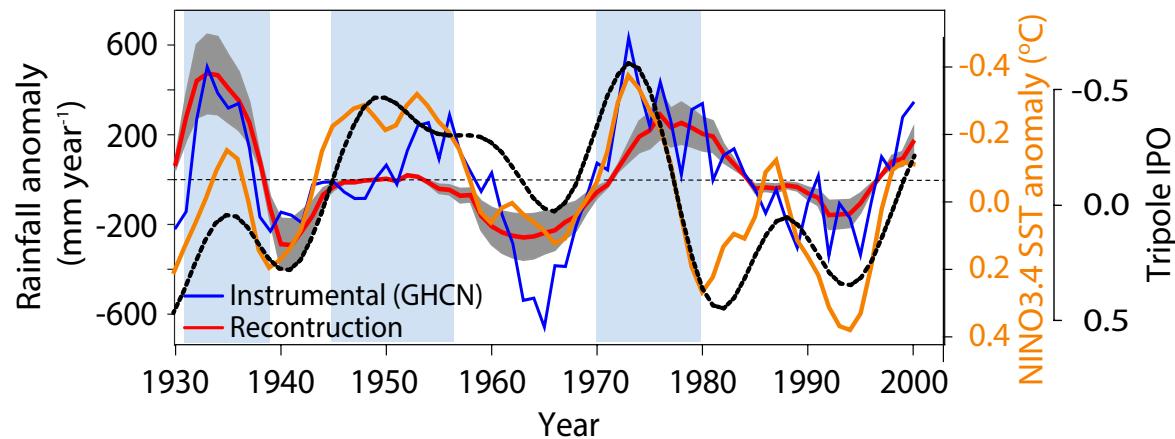
Hua et al., 2014; Quat. Geochron.

Modern SC4 $\delta^{18}\text{O}$: strong “amount effect”

Xmas rainfall (GHCN) vs. SC4 $\Delta\delta^{18}\text{O}$

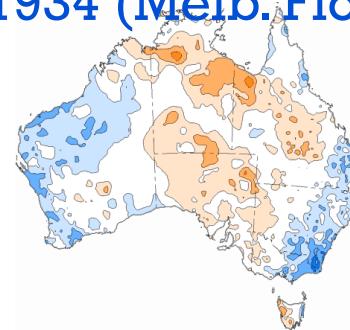


+500 mm = -0.5‰

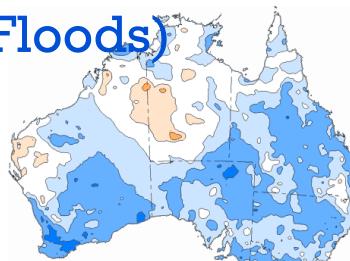


Henley *et al.*, 2015.

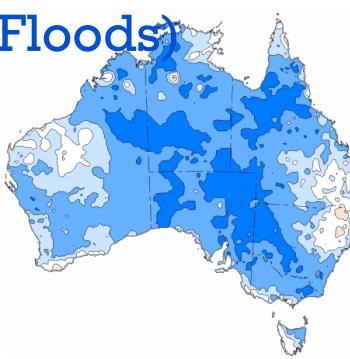
1934 (Melb. Floods)



1955 (QLD, Hunter Floods)

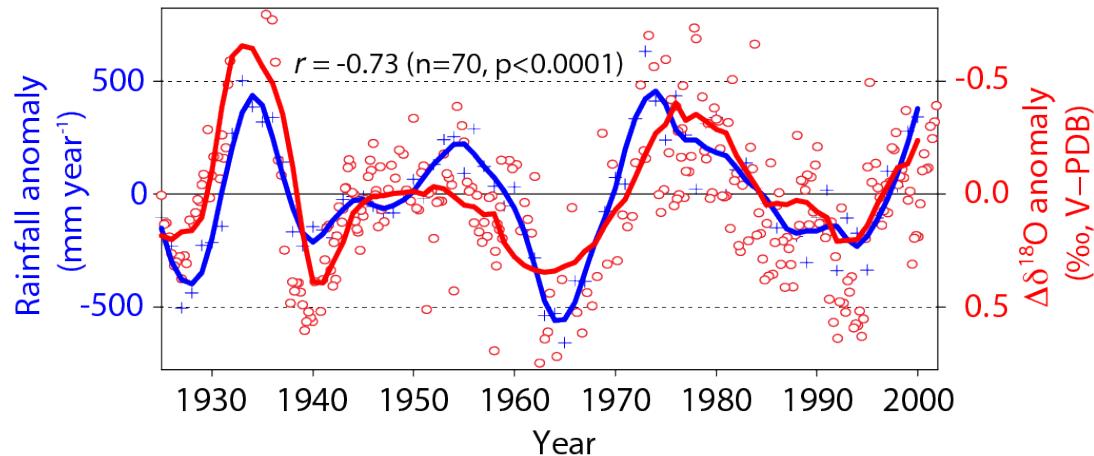


1974 (QLD, NSW Floods)

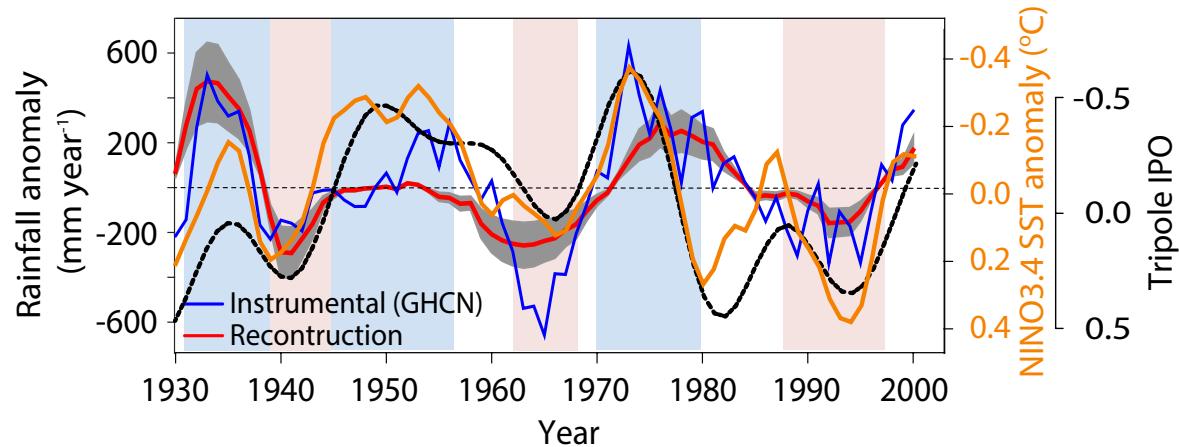


Modern SC4 $\delta^{18}\text{O}$: strong “amount effect”

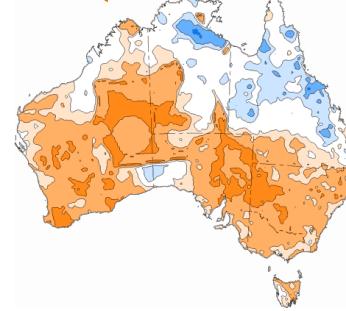
Xmas rainfall (GHCN) vs. SC4 $\Delta\delta^{18}\text{O}$



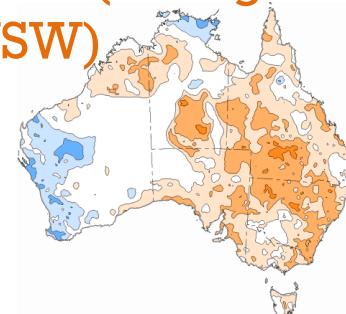
+500 mm = -0.5‰



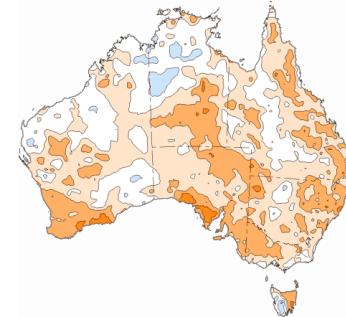
1940 (WWII Drought)



1965 (Drought begins NSW)



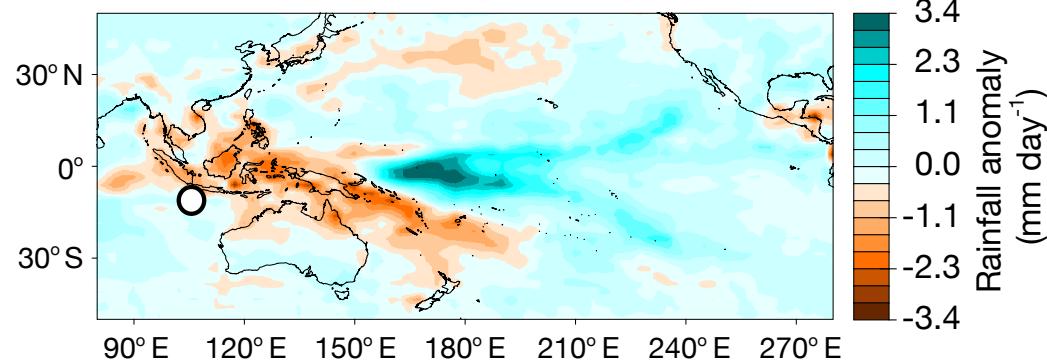
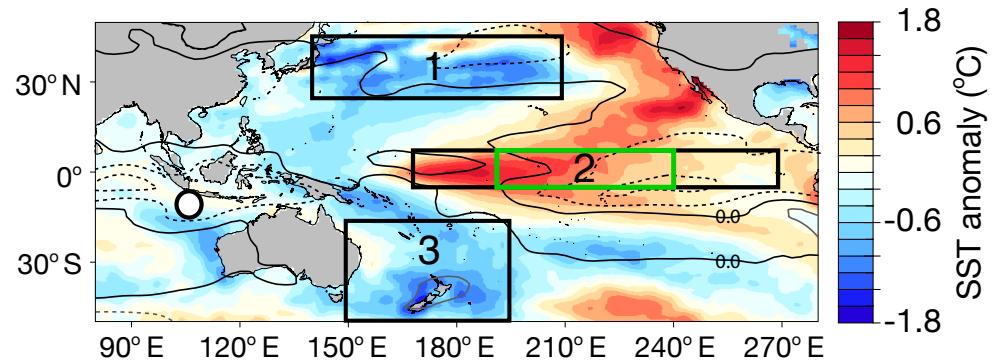
1994



Henley *et al.*, 2015.

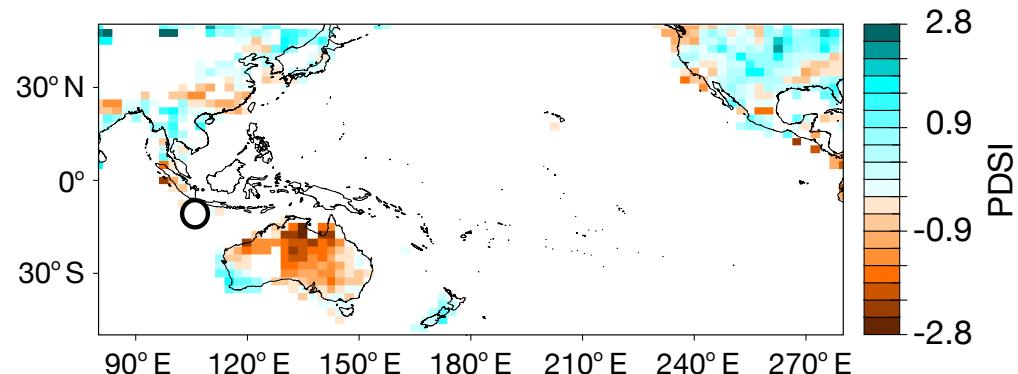
Modern SC4 $\delta^{18}\text{O}_c$: Pacific influence

Mean (NOAA OI) SST (shaded) + zonal winds (contours) for years in the lower quartile minus years in the upper quartile for the XI reconstruction.

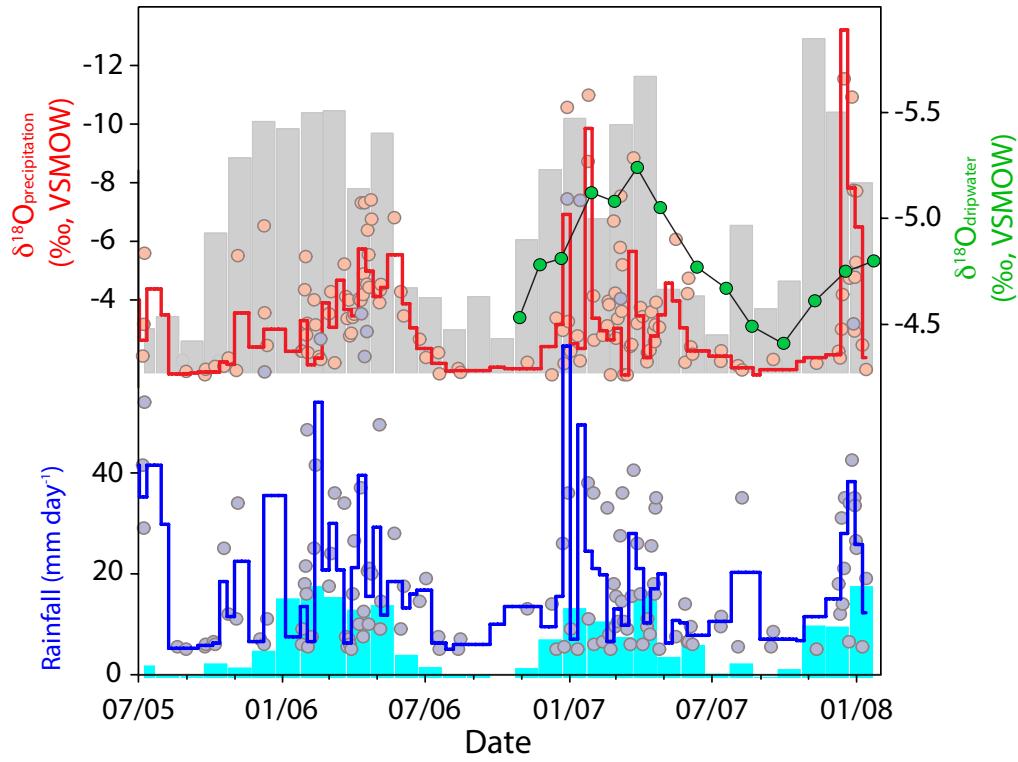


GPCP rainfall

Palmer Drought Severity Index (PDSI)

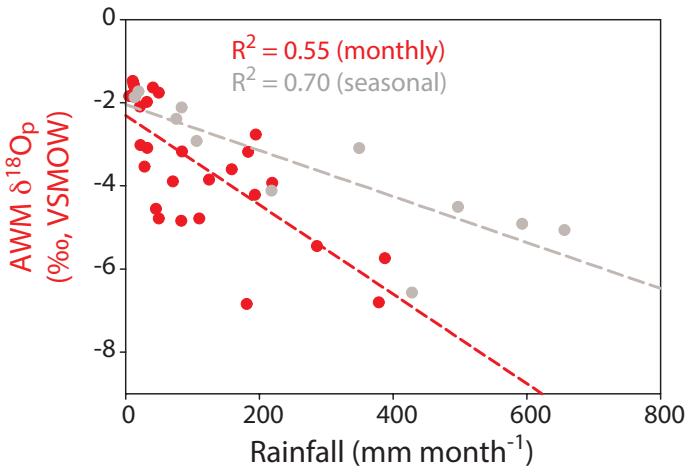


Modern rainfall and drip-water $\delta^{18}\text{O}$

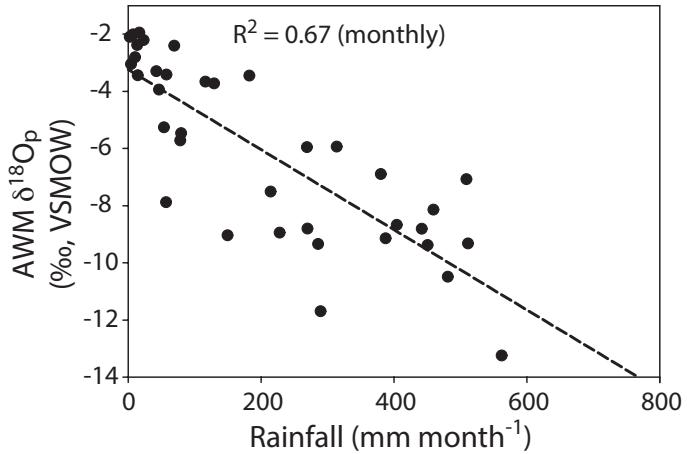


- Isotopically depleted monsoon-season rainfall \Rightarrow similar cave drip-water pattern though small seasonal range.
- Overall good correspondence with IsoGSM precipitation $\delta^{18}\text{O}$.

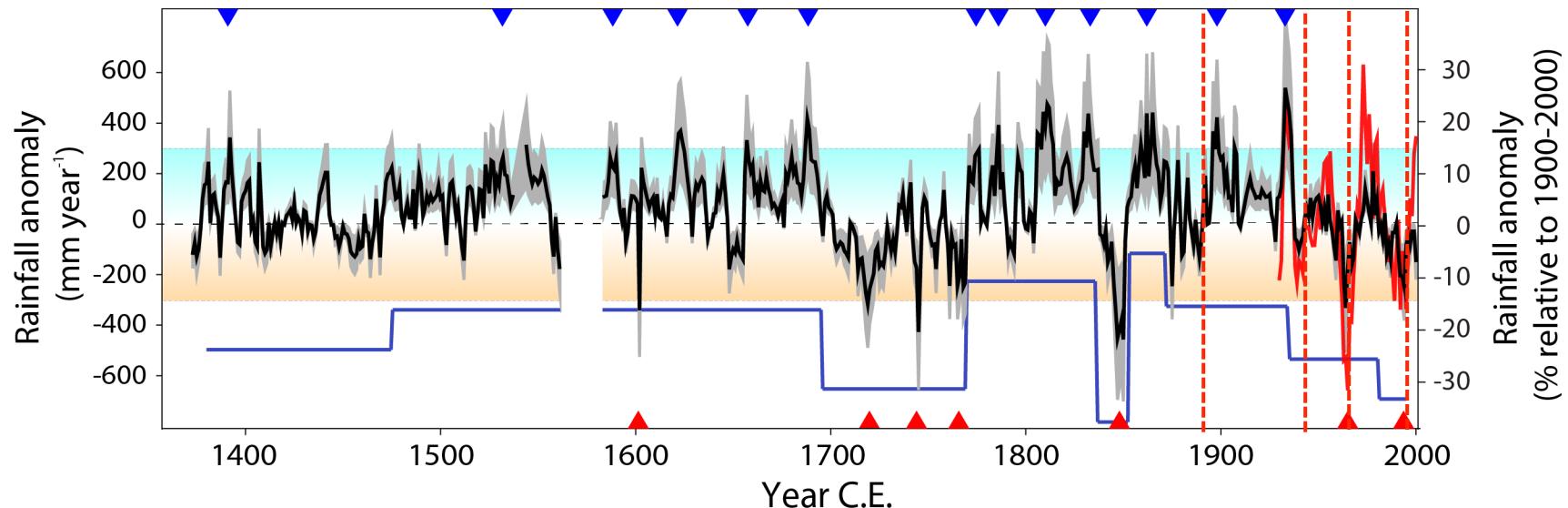
Xmas Isl. measured



Xmas Isl. IsoGSM

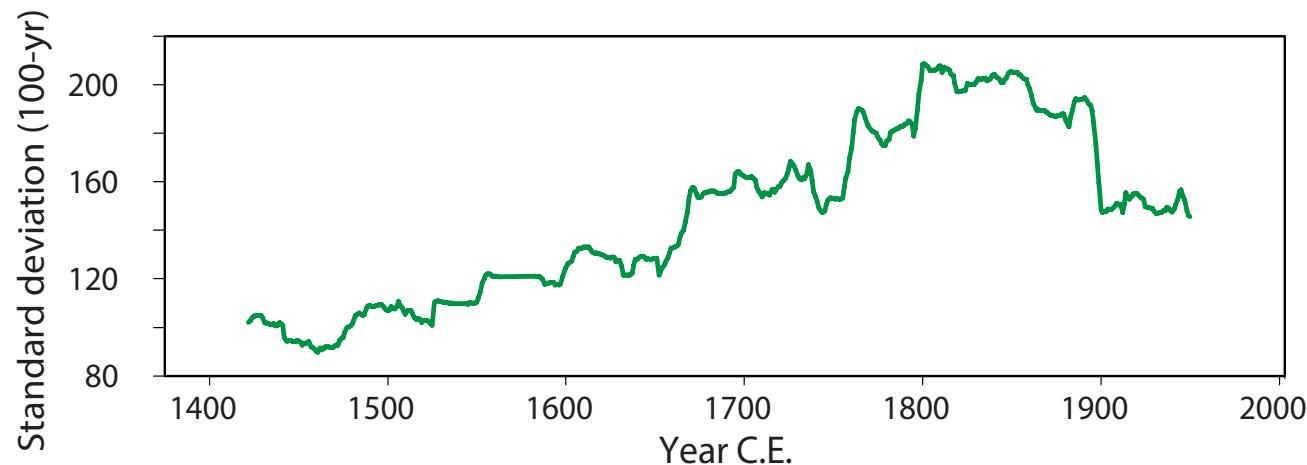


Xmas Isl. hydroclimate reconstruction

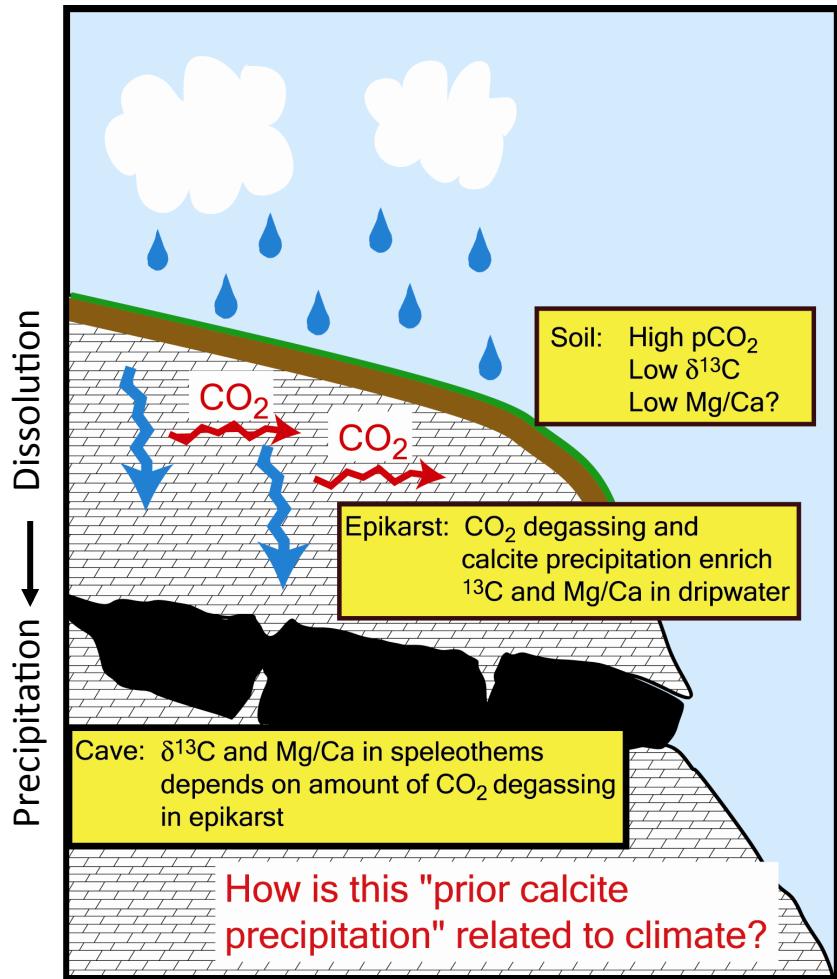


Dry: 1350-1480, ~1550s, 1700-1770, ~1850, 1890s, 20th century (Federation drought, WWII drought, 1960s, 1990s)

Wet: ~1500-1900, 1930s, 1950s, 1970s

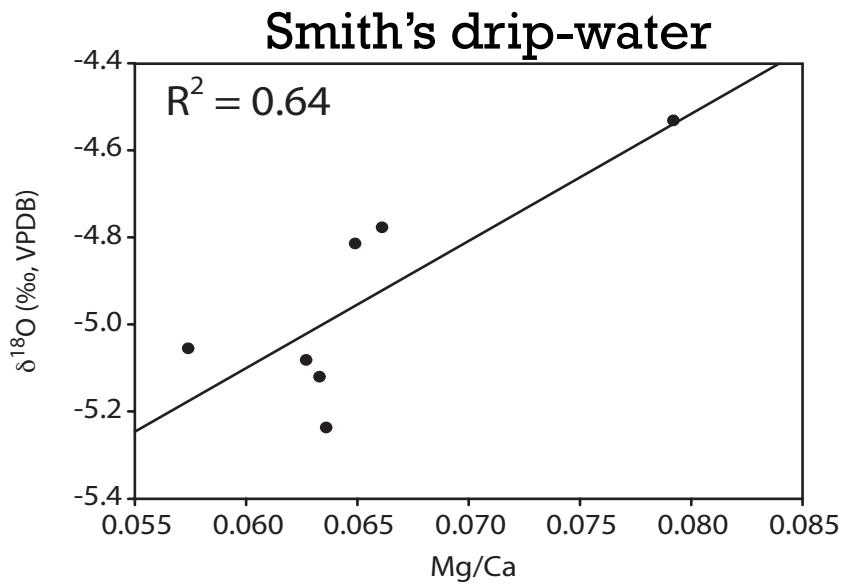


Karst hydrology: $\delta^{13}\text{C}$ and Mg/Ca

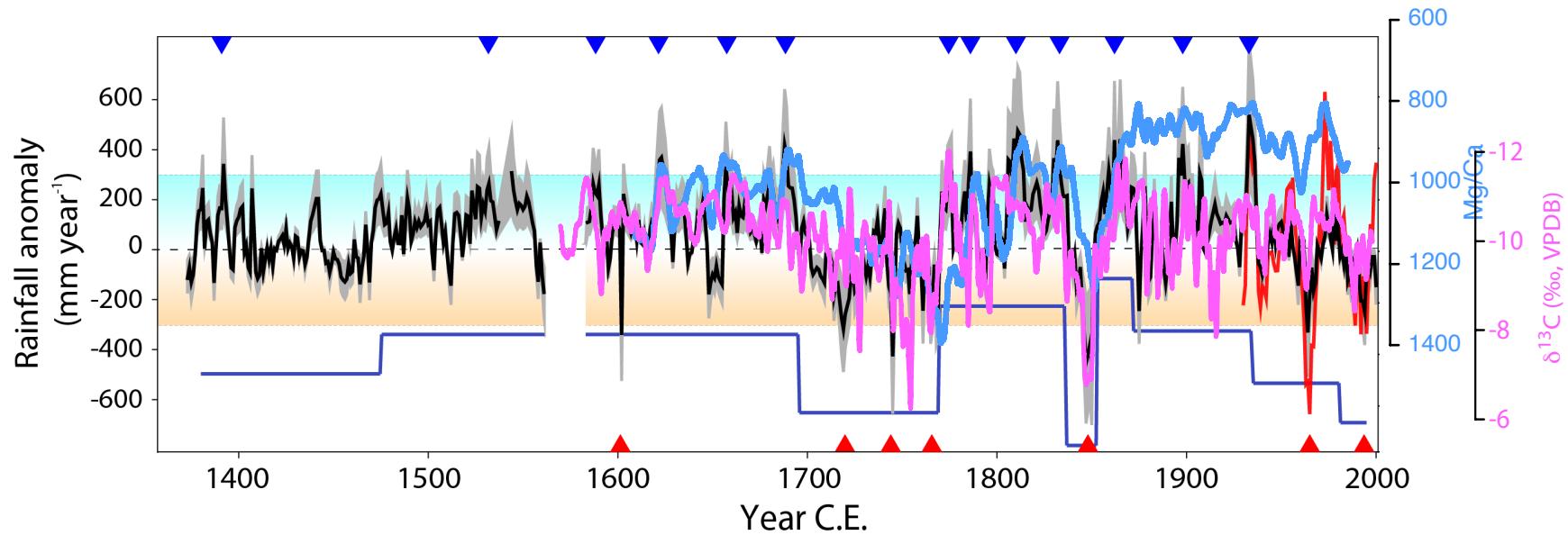


Drier periods: higher Mg/Ca and $^{13}\text{C}/^{12}\text{C}$

Wetter periods: lower Mg/Ca and $^{13}\text{C}/^{12}\text{C}$



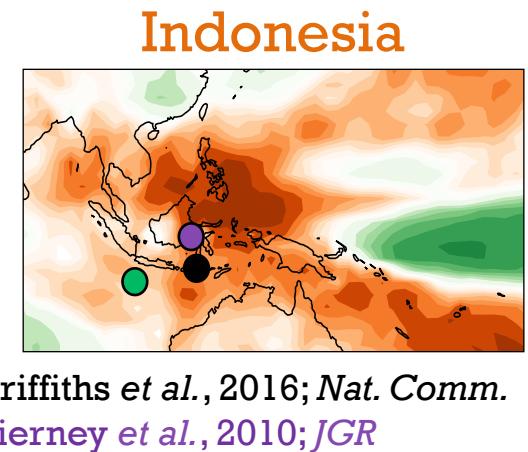
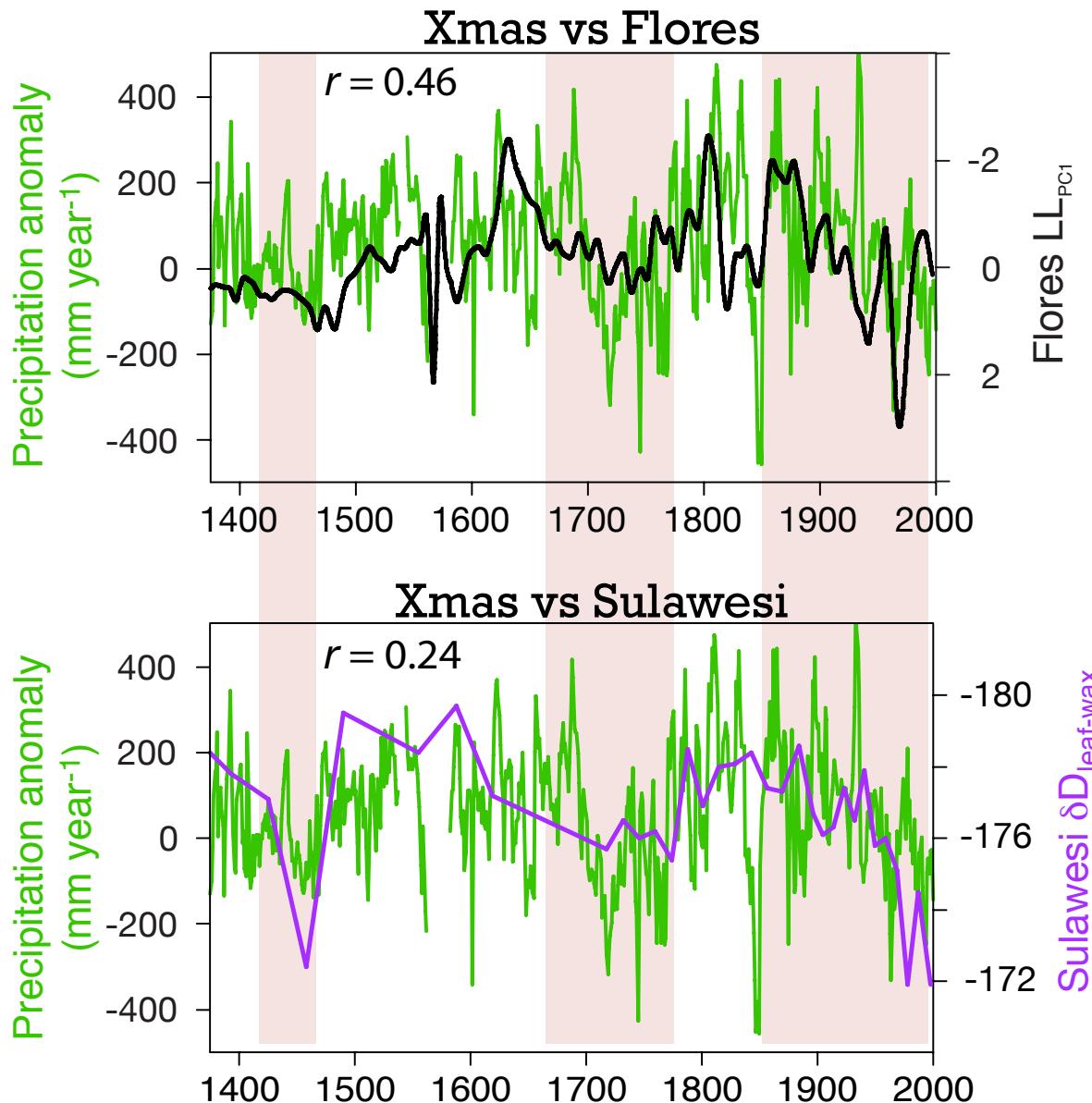
Karst hydrology: $\delta^{13}\text{C}$ and Mg/Ca



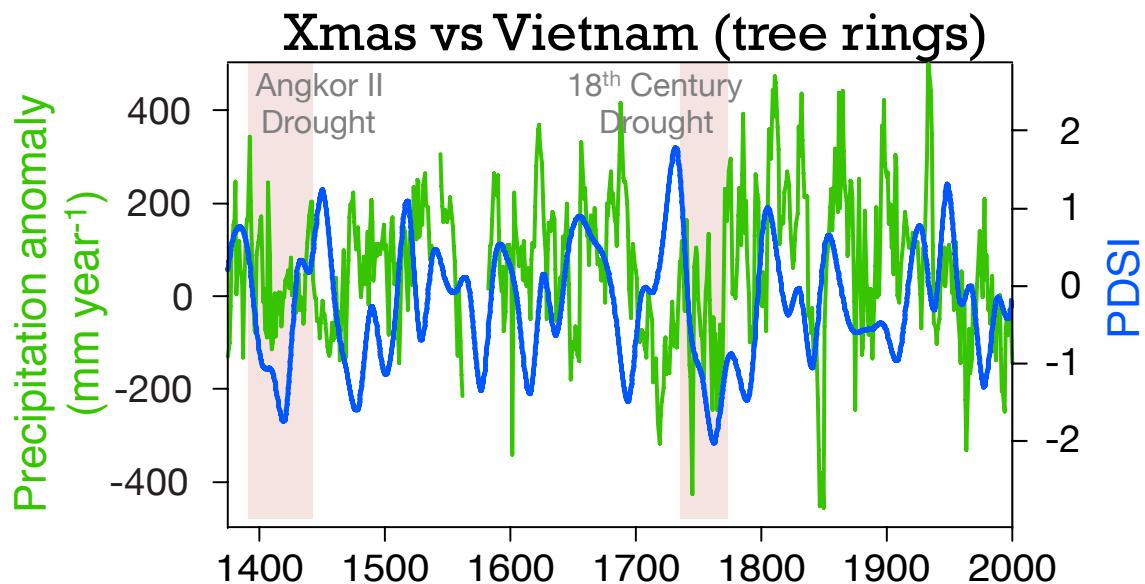
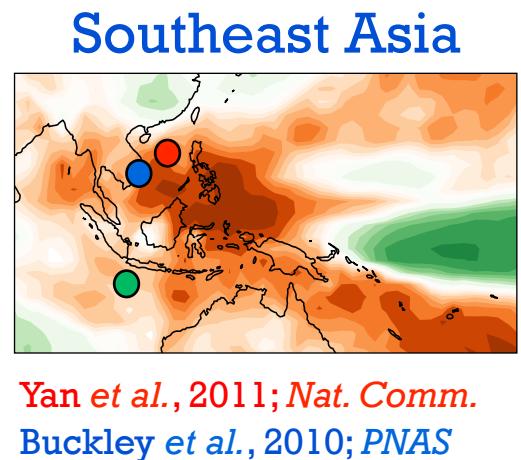
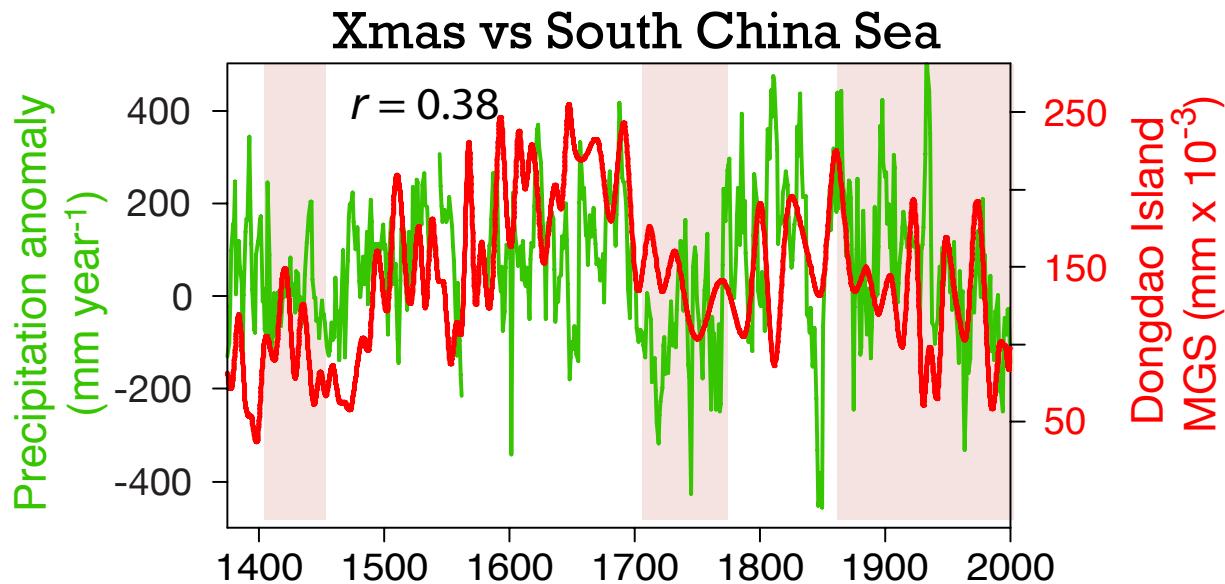
Drier periods: higher Mg/
Ca and $^{13}\text{C}/^{12}\text{C}$

Wetter periods: lower Mg/
Ca and $^{13}\text{C}/^{12}\text{C}$

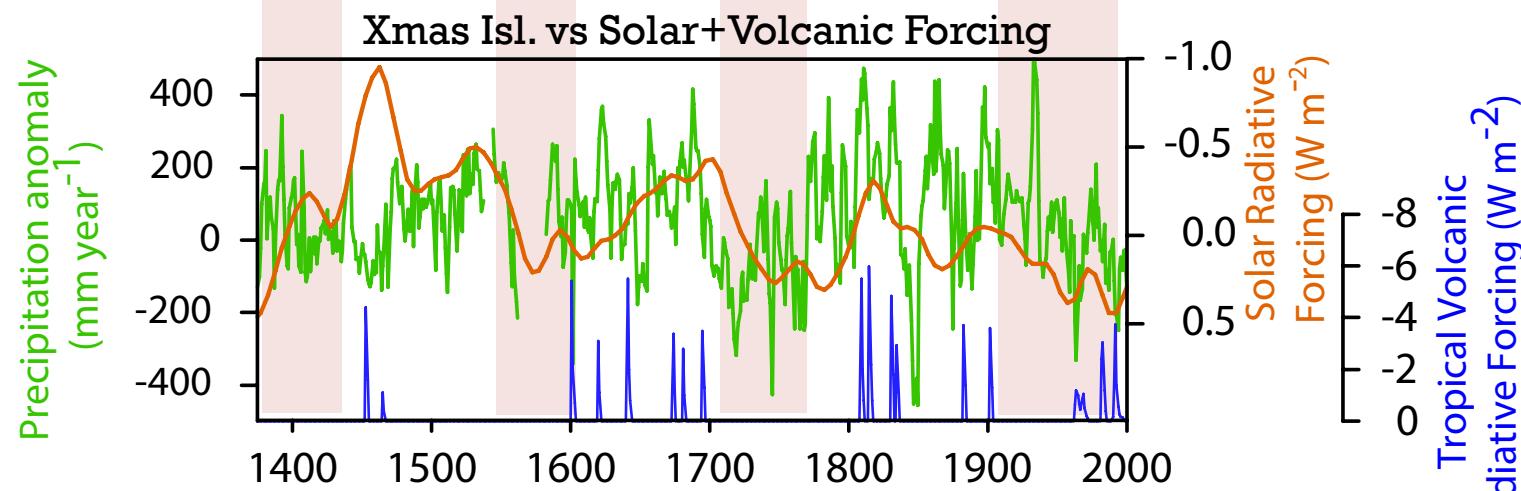
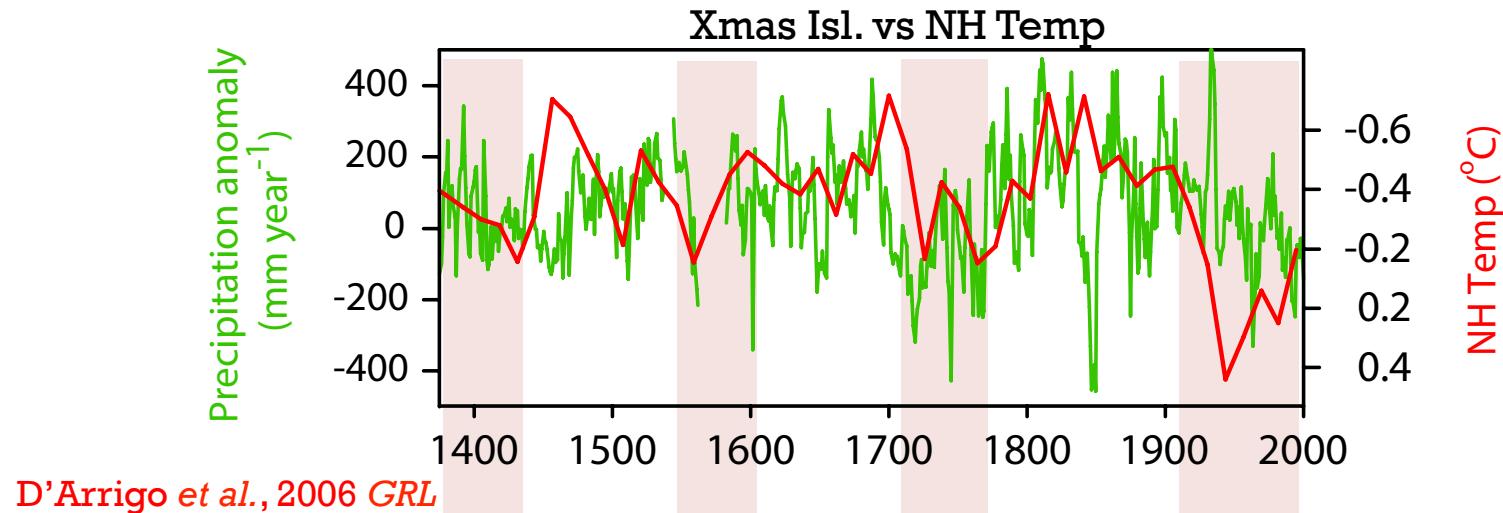
Western Pacific hydroclimate: last ~750 years



Western Pacific hydroclimate: last ~750 years



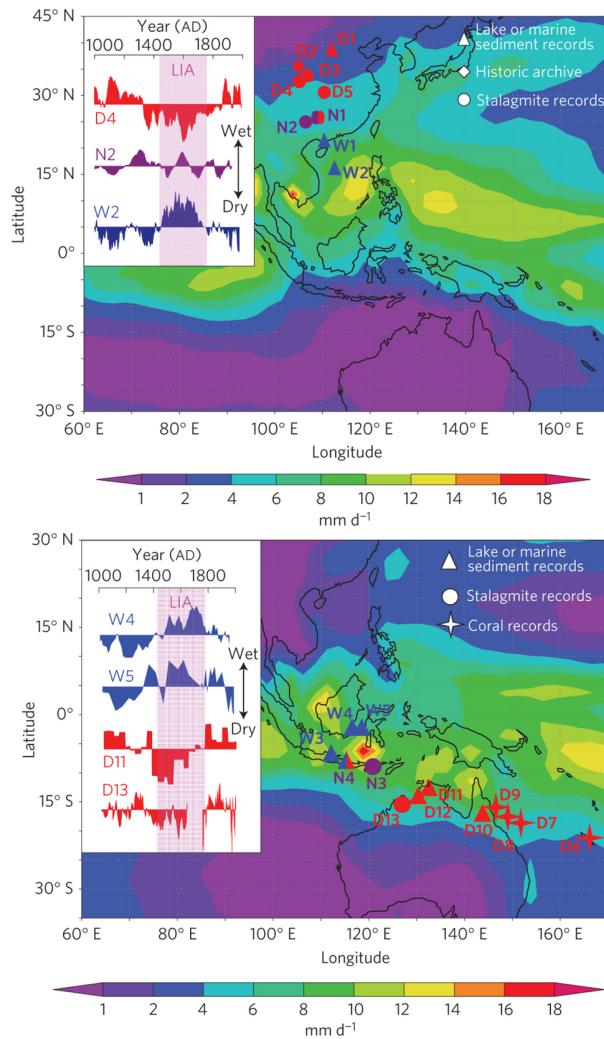
Western Pacific hydroclimate and NH Temp



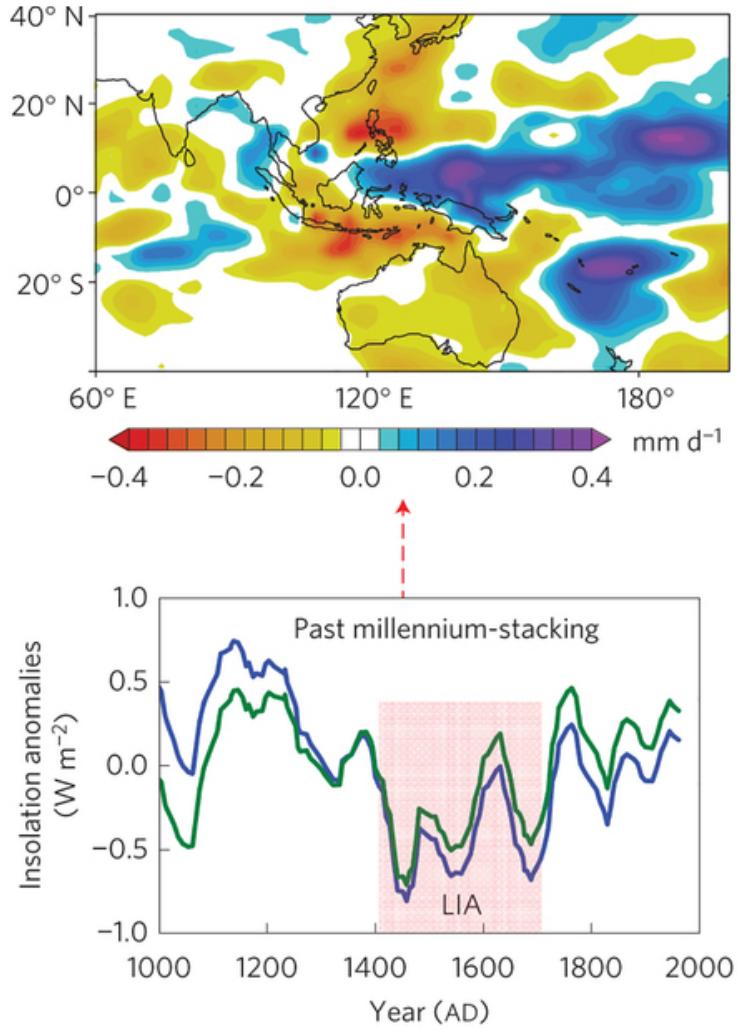
Steinhiber et al., 2009; GRL

Mann et al., 2005; J. Climate

Equatorward contraction of the ITCZ during the LIA



MPI-ESM simulated precipitation anomaly
for 1690-1740 CE



Poleward expansion of the ITCZ during the MCA

North Atlantic Drift Ice
(Bond et al., 2001)

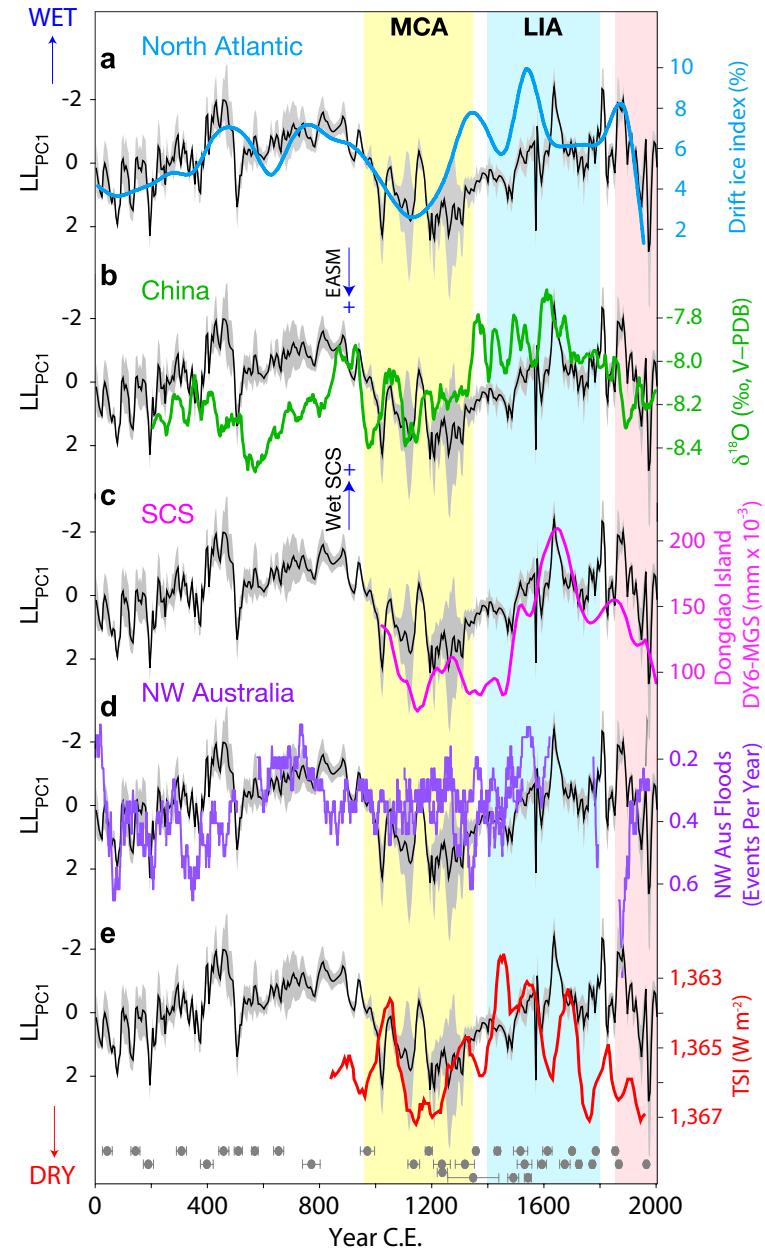
Chinese speleothem
(Zhang et al., 2008)

South China Sea
(Yan et al., 2011)

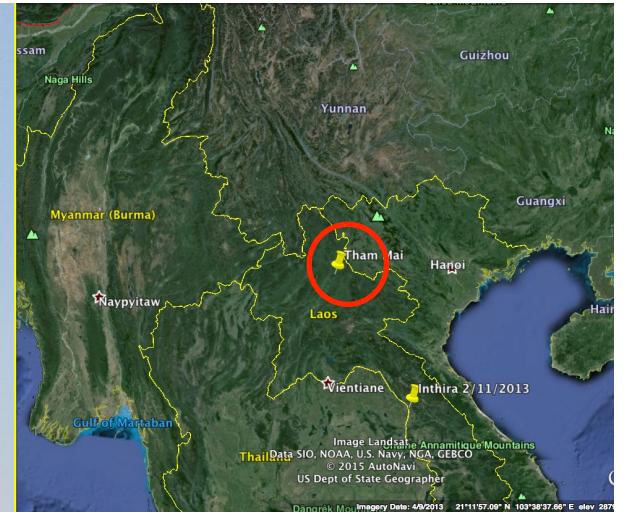
NW Australia
(Denniston et al., 2015)

TSI
(Bard et al., 2000)

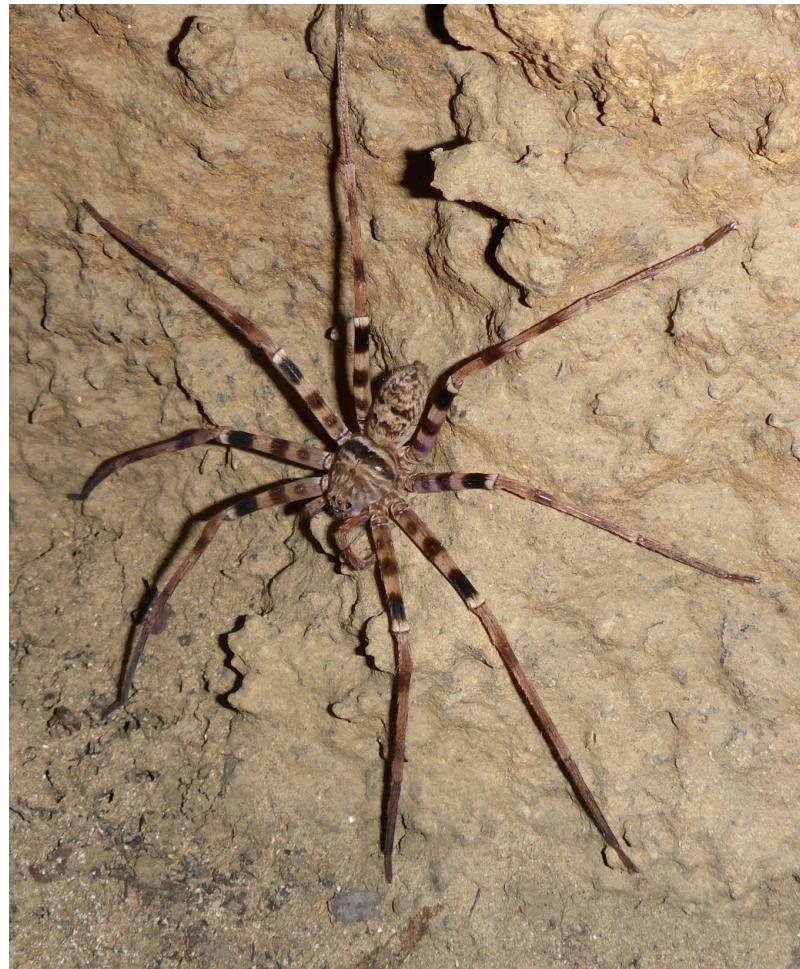
Flores multiproxy reconstruction

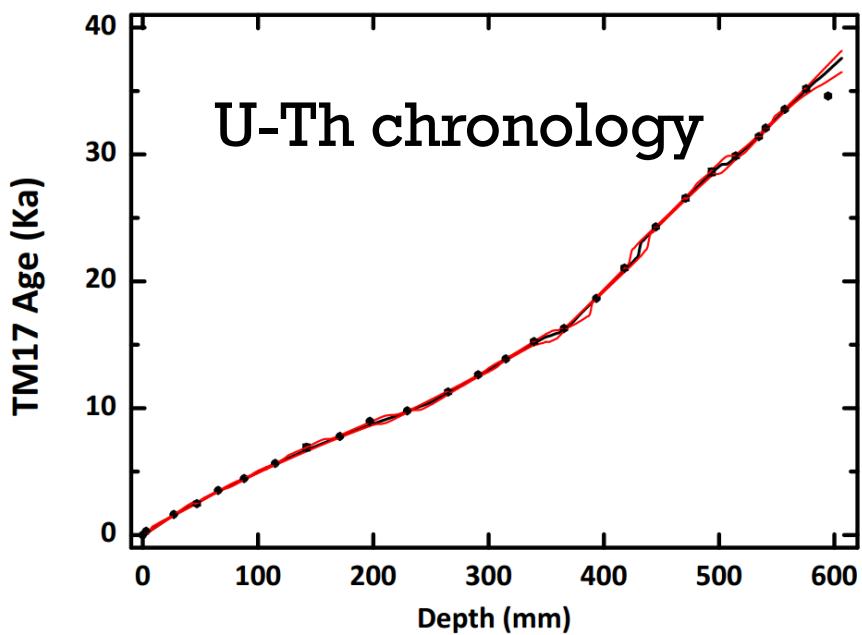


Tham Mai Cave, Luang Prabang Province, Laos

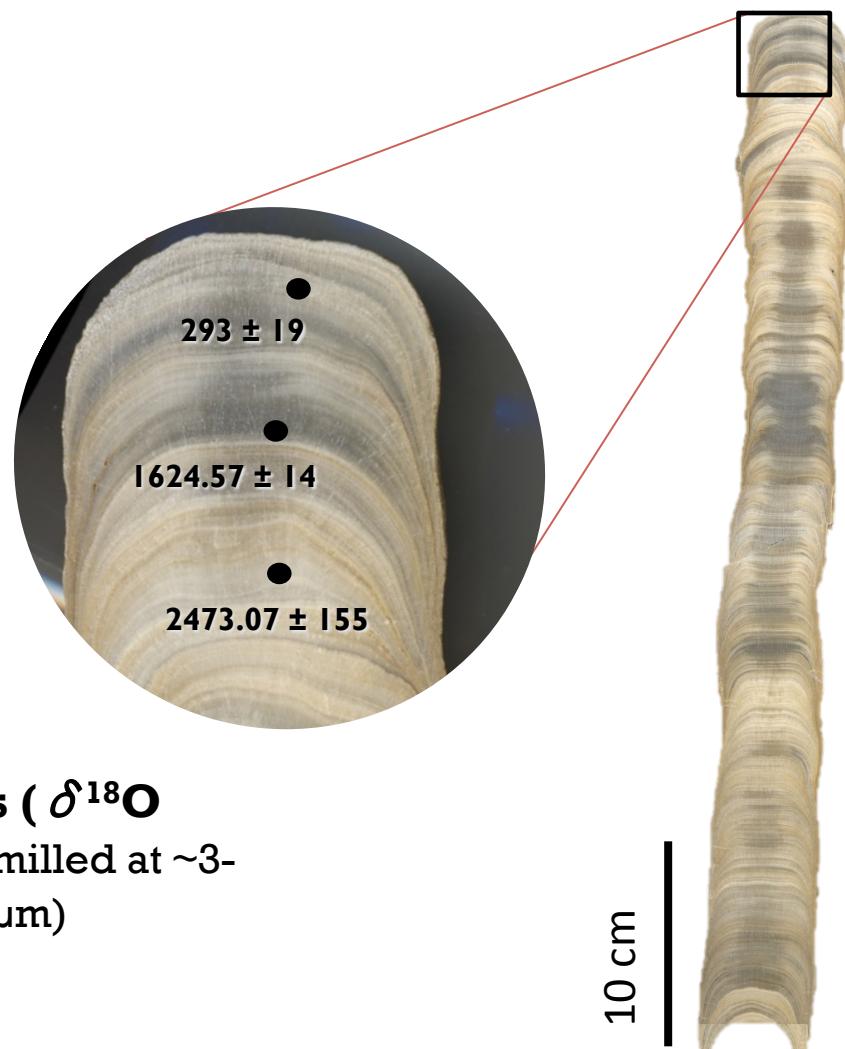


Giant Huntsman spider (*Heteropoda maxima*)

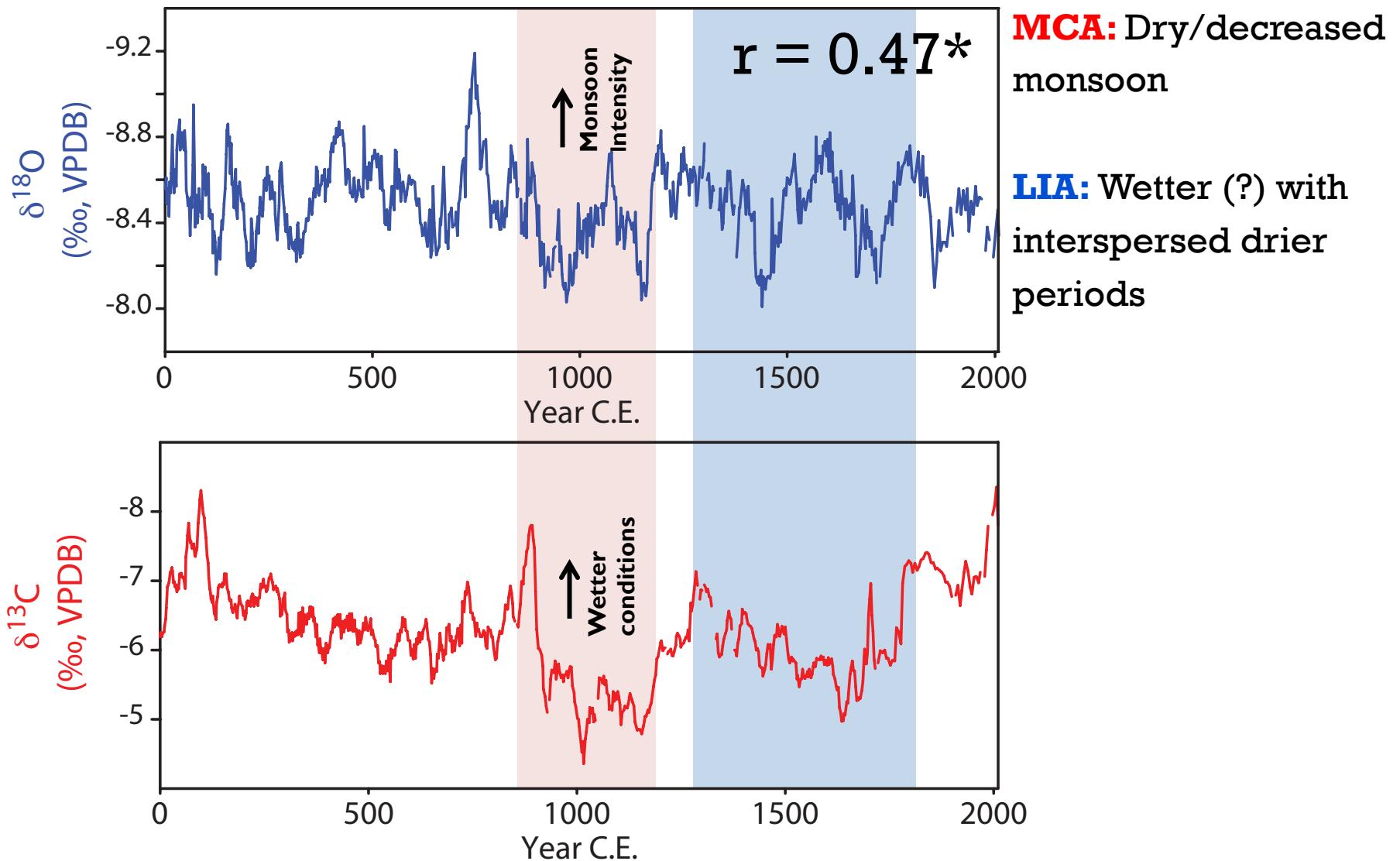




Speleothem TM-17



SE Asian monsoon variability during the Common Era



Conclusions

- Similar to other proxies (e.g. tree rings, corals), speleothems from the western Pacific have the potential to provide multi-proxy and well calibrated hydroclimate reconstructions that will be useful in model-proxy syntheses.
- New speleothem hydroclimate records from Indonesia suggest a strong link between decadal-scale rainfall variability and NH temperatures. Recent 20th century drying trend may hint at what lies ahead as the planet continues to warm.
- In light of recent findings (e.g. Yan et al., 2015), along with the new record from Laos, it is likely that the Australasian ITCZ contracted/expanded in response to radiative-forced shifts in temperature.

Acknowledgements

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