SOURCE LOCATION ALGORITHM FOR INFRASONIC MONITORING

Milton Garces, ¹ Claus Hetzer, ¹ Kent Lindquist, ² and Douglas Drob³

University of Hawaii Manoa, Lindquist Consulting, and Naval Research Laboratory

Sponsored by National Nuclear Security Administration Office of Nonproliferation Research and Engineering Office of Defense Nuclear Nonproliferation

Contract No. DE-FC04-98AL79801

ABSTRACT

Two large bolides have been recorded by International Monitoring System (IMS) infrasound stations in Hawaii and Alaska. On 25 August 2000 at 01:12:25 UTC, Department of Defense and Department of Energy satellites observed an object at 14.45 North and 106.13 West, with a total visible estimated energy of 1.4 X 10¹² joules. This object, known as the Acapulco bolide, was observed by IMS stations in Hawaii, Alaska, Bolivia, Canada, and French Guiana, as well as by DLIAR in New Mexico. On 23 April 2001, at 06:12:35 UTC, satellites observed an object at an altitude of 28.5 km at 27.9 North and 133.89 West, with a total visible energy estimate of 4.6 X 10¹² joules. This bolide explosion was observed by IMS stations in Hawaii, Alaska, California, Canada, and Germany. We use these two events to study the capabilities and limitations of source location procedures based on travel times and azimuth deviations that are derived from ray tracing formulations. A software algorithm has been developed to ingest accurate atmospheric profiles, which may be provided in near-real-time, use the tau-p method to compute the effective speed, or celerity, of specified infrasonic phases, and export these model results into any standard location algorithm, such as the Generic Locator (genloc) module within the Antelope software platform. This implementation is flexible as well as computationally efficient, and allows the exploitation of Center for Seismic Studies (CSS) (now Center for Monitoring Research) database structures and analysis tools.

OBJECTIVE

The aim of this work is to describe a ray-tracing algorithm for locating infrasonic sources, and to apply this algorithm to the determination of the position and origin time of two bolides detected by infrasound arrays in Hawaii and Alaska.

RESEARCH ACCOMPLISHED

Introduction

The tau-p method of Garcés *et al.* (1998, 2001, 2002) has been refined to clearly identify propagating infrasonic phases. Broadband infrasonic array measurements can be used to extract the amplitude, arrival time, apparent horizontal phase velocity, and azimuth of an arrival. All parameters can be estimated as a function of frequency using the PMCC method (Garcés and Hetzer, 2002). These detection parameters can be used to identify a phase, which is a prerequisite for estimating a source location. Given an atmospheric model, travel-time curves and azimuth deviations can be computed for each phase at any specified arrival azimuth. The apparent speed, or celerity, of a guided arrival is defined as the ratio of the range to the travel time for one or more multi-path bounces. Infrasonic phases identified to date are defined in Table 1, and are compatible with Brown's (1999) nomenclature and the IAVCEI list of propagating seismic phases.

Initial studies concentrated on an exact specification of the propagation path (Le Pichon *et al.*, 2002a,b; Liszka and Garcés, 2002), and these works were successful in cases when the range from the source to the receiver was not too large or along the dominant stratospheric wind direction. However, the detection of signals with high celerity propagating against the dominant stratospheric wind direction and the observation of apparent horizontal phase velocities lower than the speed of sound at the ground suggest that the interaction of infrasonic waves with atmospheric or topographical structures may scatter and diffract energy from elevated wave guides into the ground. We illustrate the critical issue of phase identification and the application of the location algorithm by using the bolide explosions of April 23, 2001, and August 25, 2000, as case studies. Because of its greater simplicity, we begin with an analysis of the 2001 bolide.

Event Selection and Array Detections of the April 23, 2001, Bolide

Infrared sensors aboard US Department of Defense (DOD) satellites detected the impact of a bolide on 23 April 2001 at 06:12:35 UTC (988006355 Epoch time). The bolide appeared to explode at an altitude of 28.5 km above the coordinates of 27.9 North and 133.89 West. The impact was simultaneously detected by space-based visible wavelength sensors operated by the US Department of Energy (DOE). The total energy in the visible band was 4.6 X 10¹² joules. The location for the April 23 event determined by the optical systems is referred to as LO1. Since bolides often propagate at speeds of a few tens of kilometers per second, and there is no guarantee that the infrared detection coincided with the infrasonic signal generation, we postulate that the time of the LOI is more precise than the location of the detection.

The locations of the three nearest IMS arrays that detected the April 23 event are given in Table 2.1, and are shown in the upper panel of Figure 1. The upper panels of Figures 2-5 shows the PMCC (Garcés and Hetzer, 2002) detections for all three arrays. Comparison with the InfraTool detections for IS59 (Garcés et al., 2001) shows that the first arrivals for PMCC are 226 seconds (4.4 minutes) earlier than the InfraTool first arrivals. This is an important result, as the first arrival time and the phase identification of that first arrival are essential to the location of infrasonic sources. Table 2.2 shows the maximum celerity values for select propagating phases. The list of Table 1 outlines all candidate phases, but for propagating ranges greater than a few thousand kilometers, it is difficult to sustain the iw phases and the Iw, Is, and It phases are not applicable. For long ranges, is phases are also found to be unstable, as the stratospheric winds can vary significantly along a meridian. However, the it, itd, and isd phases are found to exist for almost all azimuths and geographic locations. Only when an is phase exists does the itd phase disappear. For long ranges, the time contribution from the source height to the ground or the source height to the upper waveguide boundary is negligible, multiple bounces produce overlapping travel-time curves that begin to appear as a continuous curve, and shadow zones disappear. The itd and isd phases (Table 1) correspond to leaky wave guides that are suspended above the ground, but either scatter or diffract acoustic energy to the ground. Sources that explode at heights of ~30 km, such as the April 23 bolide, place acoustic energy in the middle of the low-velocity zone of the stratosphere, and thus efficiently trap energy in the stratospheric duct. This waveguide would be able to duct energy with minimal attenuation, and most first arrivals with high celerity may be attributed to isd phases. Due to the continuous appearance of travel-time curves and the further degradation of shadow zones by

scattering and diffraction, we opt to use the celerity as the key propagation model output for the source location iteration.

Infrasonic Location Procedure for April 23 Event

From intersecting back azimuths, it is possible to produce a seed location for the source inversion procedure. However, as can be seen in Figure 1, the back azimuth location may be expected to be in error. An initial origin time can be obtained from the seed location by assuming a constant celerity of 0.3 km/s and computing the great circle paths to the source. Alternatively, if ground truth is known on an event, such as the epicentral location issued by the DoD release, the range of each station to the source and the expected azimuth of the incoming signal can be readily computed. Once an initial azimuth from source to receiver is determined, the tau-p method of Garcés et al. (1998, 2002) can be used to compute the celerity of each propagating phase. The lower panels of Figures 2-4 show the celerity computed at each array for the April 23 bolide using the Naval Research Laboratory's (NRL) SAGE atmospheric profiles that include accurate specifications of the troposphere and stratosphere (Garcés et al., 2001). The estimated times of arrival of the first or second arrival at the station were used to compute the residuals from a grid search around the seed source location and origin time. Various locations were made assuming different phase identifications for both first and second arrivals (when present), but only two solutions (with the minimal residuals) are shown in Table 2.3. The first solution, LA S1, assumes all first arrivals correspond to isd phases, and produced a very good match to the origin time provided by the satellite observation (LO1). Although in seismic location an 8-s time differential is unacceptable, when scaled to the total travel time of the signal, the percent error is small and comparable to the best seismic location accuracies (Table 2.4). The difference in the source location between LO1 and LA S1 may be attributed to the high speed of the bolide and could correspond to a difference in where the peak sound and infrared energy are radiated. This result suggests that isd phases can be used to explain the arrival of signals in the upstream stratospheric direction, and the relatively higher frequency of the first arrivals observed in Figure 2 and 3 suggest that scattering is an important factor. A second solution, LA S2, assumed that the first arrival at IS53 was an isd phase, the second arrival at IS57 was an itd phase, and the second arrival at IS59 was an isd phase. These phases were selected in an attempt to match the source location at the expense of the origin time. However, we favor the first solution because we believe the uncertainty in time ($\sim 10 \text{ s}$) may be less than the uncertainty in position for the satellite location LO1.

Infrasonic Location Procedure for the August 25, 2000, Event

The August 25, 2000, bolide, known as the Acapulco bolide, is more difficult to unravel. Table 3.1 shows the infrasonic observations for the Acapulco bolide. The arrival information for the stations in South America was derived from PMCC results (Le Pichon, personal communication, 2000). The lower panel of Figure 1 shows the station locations and back azimuths, again showing a poor azimuth fit of some of the stations to the actual source location. In general, it appears that the worst azimuth deviations occur in stations with extreme topography along the propagation path, specifically IS53 (Alaska) and IS08 (Bolivia). From the seed locations, azimuths from the stations to the source were estimated and the celerity computed at each station. Table 3.2 shows the maximum celerity for each phase. Figure 6 shows the PMCC detection at DLIAR, the closest station, and the change in detection azimuth with time. Following the same procedure as for the April 23 event, we computed locations for various permutations of phase identifications for the first and second arrival times. An attempt was also made to separate the detections into two events, one recorded by the US stations and another by the South American stations. However, two-station locations are inherently unstable, and even the solution using the three US stations was unstable because of the near parallel alignment of the DLIAR and IS53 propagation paths from the source. Three solutions are shown in Table 3.3. The first solution, LA DL1, assumes all first arrivals are isd phases, and yielded a large error on the origin time estimate as well as in the time differentials for some of the stations (Table 3.4). The second solution, LA DL2, assumed a combination of it, itd and isd phases (Table 3.5) and yielded a good fit to the origin time, although the time differentials per station are not as good as the April 23 solution. A source location west of the LO1 location is consistent with the DLIAR observations. Further work is needed to determine whether this event consists of a single bolide or more than one bolide arriving at different times, as suggested by the infrasound detections from IS25.

CONCLUSIONS AND RECOMMENDATIONS

Infrasonic estimates of the origin time and location of the April 23 event were performed using realistic atmospheric profiles and the tau-p method to estimate the celerity of propagating phases in the atmosphere. Arrivals corresponding to phases propagating in leaky stratospheric ducts produced the best fit to the origin time obtained from the satellite infrared detection. This existence of these leaky stratospheric ducts may explain how stratospheric

phases can be observed along both the upstream and downstream dominant stratospheric wind directions. Further work is needed on how topography and perturbations in the atmosphere can scatter and diffract infrasonic energy into and out of elevated wave guides. The next generation of the tau-p software will address propagation in range-dependent environments.

REFERENCES

- Brown, D.J. (1999). Summary of infrasound source location meeting: San Diego, November 9-10, 1998, *Center for Monitoring Research*, Arlington, VA.
- Garcés, M, and C. Hetzer (2002). Evaluation of infrasonic detection algorithms. This volume.
- Garcés, M., D. Drob, and M. Picone (2002). A theoretical study of the effect of geomagnetic fluctuations and solar tides on the propagation of infrasonic waves in the atmosphere. *Geophys. J. International*, **148**, 77-87.
- Garcés, M, C. Hetzer, K. Lindquist, and R. Hansen, J. Olson, C. Wilson, D. Drob, M. Picone, (2001). Infrasonic location of the April 23, 2001, Bolide event. 23nd Annual DTRA/NNSA Seismic Research Review, Jackson Hole, 1-5 October 2001.
- Garcés, M, D. Drob, M. Picone, K. Lindquist, and R. Hansen (2000). Characterization Of Infrasonic Waves Observed In Hawaii During Summer, 22nd Seismic Research Symposium: Planning for Verification of and Compliance with The Comprehensive Nuclear-Test-Ban Treaty, New Orleans.
- Garcés, M. A., R. A. Hansen, and K. Lindquist (1998). Travel times for infrasonic waves propagating in a stratified atmosphere, *Geophys. J. International*, **135**, 255-263.
- Le Pichon, A. M. Garcés, E. Blanc, M. Barthélémy, D. Drob (2002). Acoustic propagation and atmosphere characteristics derived from infrasonic waves generated by the Concorde. *J. Acoust. Soc. Am.*, **111**, 629-641.
- Le Pichon, A., J. Guilbert, A. Vega, and M. Garcés (2002), Ground-coupled air-waves and diffracted infrasound from the Arequipa earthquake of June 23, 2001, *Geophysical Research Letters* (in press).
- Liszka, L., and M. Garcés (2002), Infrasonic Observations of the Hekla Eruption of February 26, 2000. *Journal of Low Frequency Sound* (in press).
- Kulichkov, S. N. (1998). On problems of infrasonic monitoring of small energy explosions, *Proceedings of the Informal Workshop on Infrasound*, Bruyers-Le-Chatel, France.

Table 1. Preliminary phase identification nomenclature for long-range infrasonic propagation

Phase ID	Description	Typical celerity of first arrival, m/s
iw	Guided wave propagating between the tropopause and the ground.	330-340
is	Guided wave propagating between the stratopause and the ground.	310-330
isd	Guided wave propagating in elevated waveguide between stratopause and the troposphere, and diffracted or scattered to the ground. May have higher frequency.	310-330
it	Guided wave propagating between the lower thermosphere and the ground.	280-300
itd	Guided wave propagating in elevated waveguide between the lower thermosphere and the troposphere, and diffracted or scattered to the ground.	280-300
It, Is, Iw	Direct arrival from the source to the receiver. May have high apparent phase velocity	N/A

Table 2.1. First arrival observations of April 23, 2001, bolide by three nearest IMS infrasound stations

Station	Lat (N)	Lon (E)	Speed (m/s)	Azimuth	ETA (epoch)
IS53	64.87	-147.84	322	150.6	988020690
IS57	33.6	-116.5	349	256.1	988012060
IS59	19.59	-155.9	348	63.6	988013790

Table 2.2. Predicted first arrival celerity (km/s) for select phases: April 23, 2001

	it	itd	isd
IS53	0.268	0.284	0.292
IS57	0.278	0.292	0.328
IS59	0.284	0.3	0.311

Table 2.3. Source location and errors relative to satellite location (LO1): April 23, 2001

Source	Lat (N)	Lon (E)	Origin Time (Epoch)	Lat error (deg)	Lon error (deg)	Time error (s)
LO1	27.9	-133.89	988006355	0	0	0
LA_S1	28.07	-135.09	988006347	0.17	-1.2	-8
LA_S2	27.79	-133.42	988006145	-0.11	0.47	-210

Table 2.4. Phase selection and time error for LA_S1 solution: April 23, 2001 PA = predicted arrival; ETA = first arrival time; OT = origin time

Station	Selected Phase	Range to LA_S1	X/T (km/s)	Predicted arrival	Time Error /Travel Time	PA-ETA	PA-OT	Ratio (%)
IS53	isd	4191	0.292	988020700	-0.06%	10	14353	0.07
IS57	isd	1877	0.328	988012070	-0.14%	10	5723	0.17
IS59	isd	2313	0.311	988013784	-0.11%	-6	7437	0.08

Table 3.1. First arrival observations of August 25, 2000, bolide by infrasound stations.

Station	Lat (N)	Lon (E)	Speed (m/s)	Azimuth	ETA (epoch)
IS08	-16.3	-68.1	340	298	967184490
IS25	5.2	-52.7	344	283	967187730
IS53	64.87	-147.84	355	144	967187230
IS59	19.59	-155.9	346	88	967182900
DLIAR	35.87	-106.33	360	180	967173900

Table 3.2. Predicted first arrival celerity (km/s) for select phases: August 25, 2000

	it	itd	isd
IS08	0.289		0.298
IS25	0.263	0.285	0.29
IS53	0.292	0.303	0.31
IS59	0.281	0.297	0.308
DLIAR	0.278		0.309

Table 3.3. Source location and errors relative to satellite location (LO1): August 25, 2000

Source	Lat (N)	Lon (E)	Origin Time (Epoch)	Lat error (deg)	Lon error (deg)	Time error (s)
LO1	14.45	-106.13	967165945	0	0	0
LA_DL1	13.68	-108.21	967166247	-0.77	-2.08	302
LA_DL2	13.37	-107.74	967165950	-1.08	-1.61	5

Table 3.4. First arrival phase selection and time error for LA_DL1 solution: August 25, 2000

	Selected	Range to	X/T	Predicted	Time Error /			
Station	Phase	LA_DL1	(km/s)	Arrival	Travel Time	PA-ETA	PA-OT	Ratio (%)
IS08	isd	5513	0.298	967184747	1.63%	257	18500	1.39
IS25	isd	6149	0.290	967187450	1.39%	-280	21203	-1.32
IS53	isd	6432	0.310	967186995	1.42%	-235	20748	-1.13
IS59	isd	5110	0.308	967182838	1.78%	-62	16591	-0.37
DLIAR	isd	2464	0.309	967174221	3.80%	321	7974	4.03

Table 3.5. First arrival phase selection and time error for LA DL2 solution: August 25, 2000

Station	Selected Phase	Range to LA_DL2	X/T (km/s)	Predicted Arrival (PA)	Time error / Travel time	PA-ETA	PA-OT	Ratio (%)
IS08	it	5452	0.289	967184815	0.03%	325	18865	1.72
IS25	itd	6096	0.285	967187339	0.02%	-391	21389	-1.83
IS53	itd	6481	0.303	967187339	0.02%	109	21389	0.51
IS59	isd	5168	0.308	967182729	0.03%	-171	16779	-1.02
DLIAR	isd	2495	0.309	967174024	0.06%	124	8074	1.54

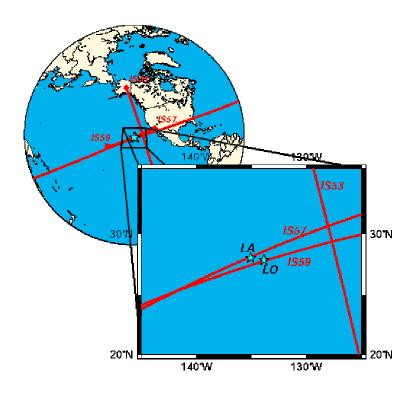


Figure 1. Satellite (LO) and infrasonic (LA_S1) location for the April 23, 2001, bolide. The observed arrival azimuths at infrasound arrays IS53, IS57, and IS59 are shown as red lines.

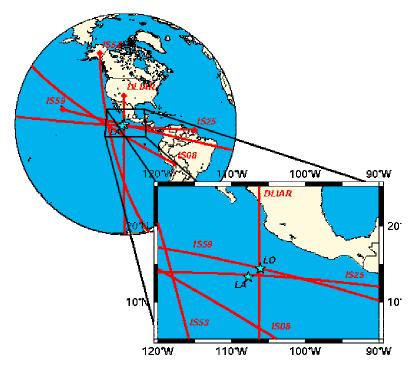
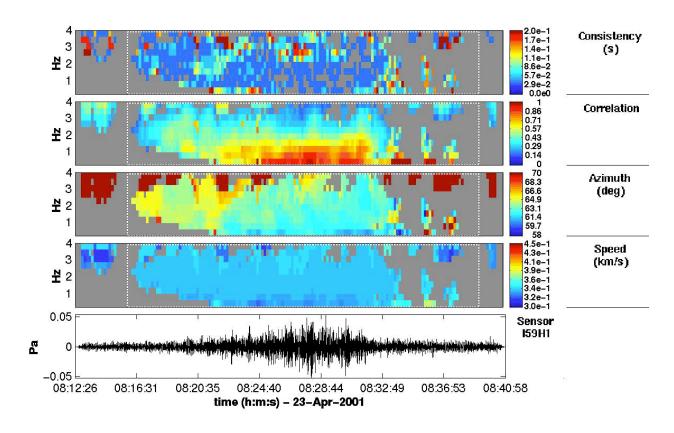


Figure 2. Satellite (LO) and infrasonic (LA_DL2) location for the Acapulco bolide. The observed arrival azimuths at infrasound arrays IS08. IS25, IS53, DLIAR, and IS59 are shown as red lines.



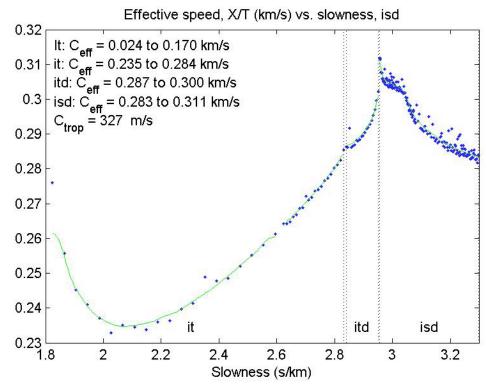
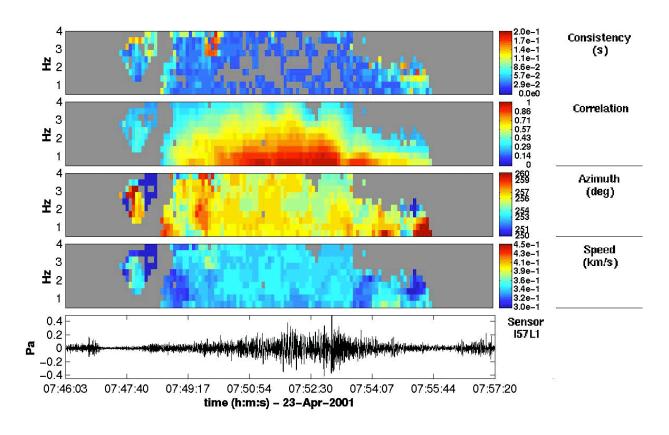


Figure 3. Detection and predicted celerity at IS59, Hawaii, for April 23, 2001, bolide.



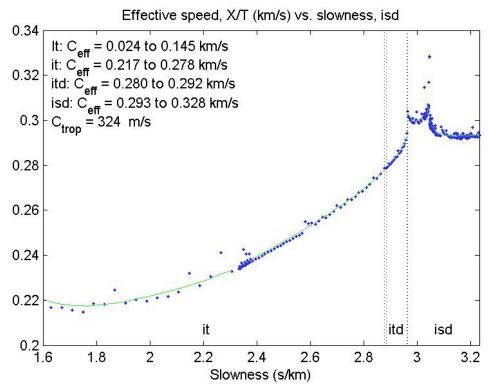
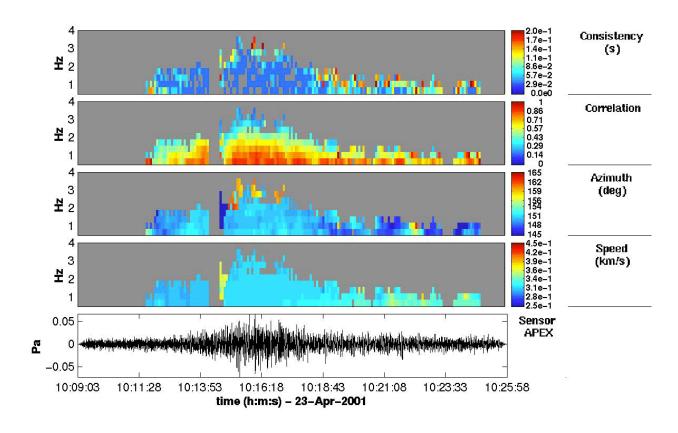


Figure 4. Detection and predicted celerity at IS57, Piñon Flats, for April 23, 2001, bolide.



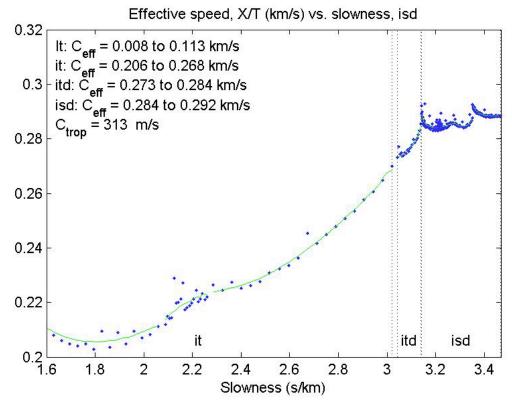


Figure 5. Detection and predicted celerity at IS53, Fairbanks, Alaska, for April 23, 2001, bolide.

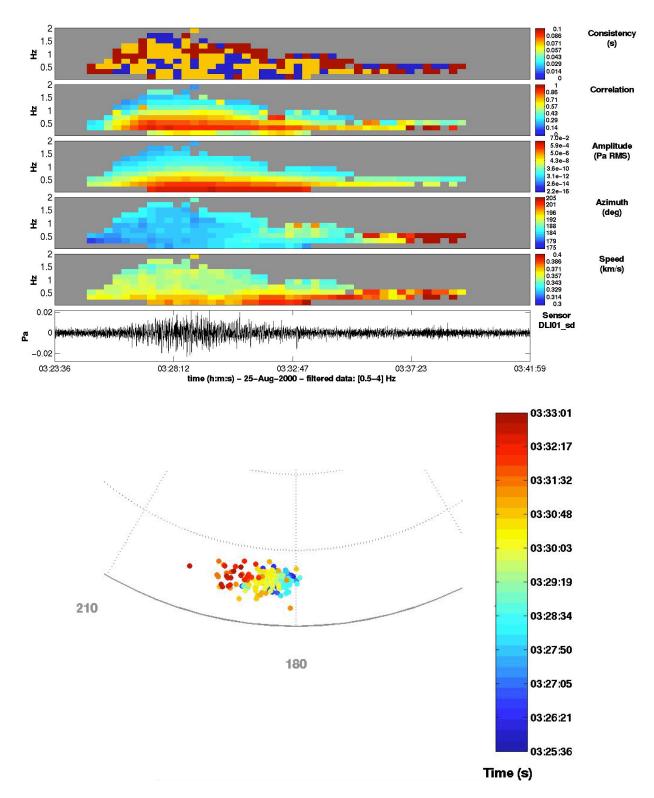


Figure 6. Detections and azimuth changes as a function of time observed at DLIAR, New Mexico, for the Acapulco bolide.