

Lecture 23

Exemplary Inverse Problems
including
Earthquake Location

Syllabus

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Lecture 03	Probability and Measurement Error, Part 2
Lecture 04	The L_2 Norm and Simple Least Squares
Lecture 05	A Priori Information and Weighted Least Squared
Lecture 06	Resolution and Generalized Inverses
Lecture 07	Backus-Gilbert Inverse and the Trade Off of Resolution and Variance
Lecture 08	The Principle of Maximum Likelihood
Lecture 09	Inexact Theories
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Lecture 20	Linear Operators and Their Adjoint
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Lecture 23	Exemplary Inverse Problems, incl. Earthquake Location
Lecture 24	Exemplary Inverse Problems, incl. Vibrational Problems

Purpose of the Lecture

solve a few exemplary inverse problems

thermal diffusion

earthquake location

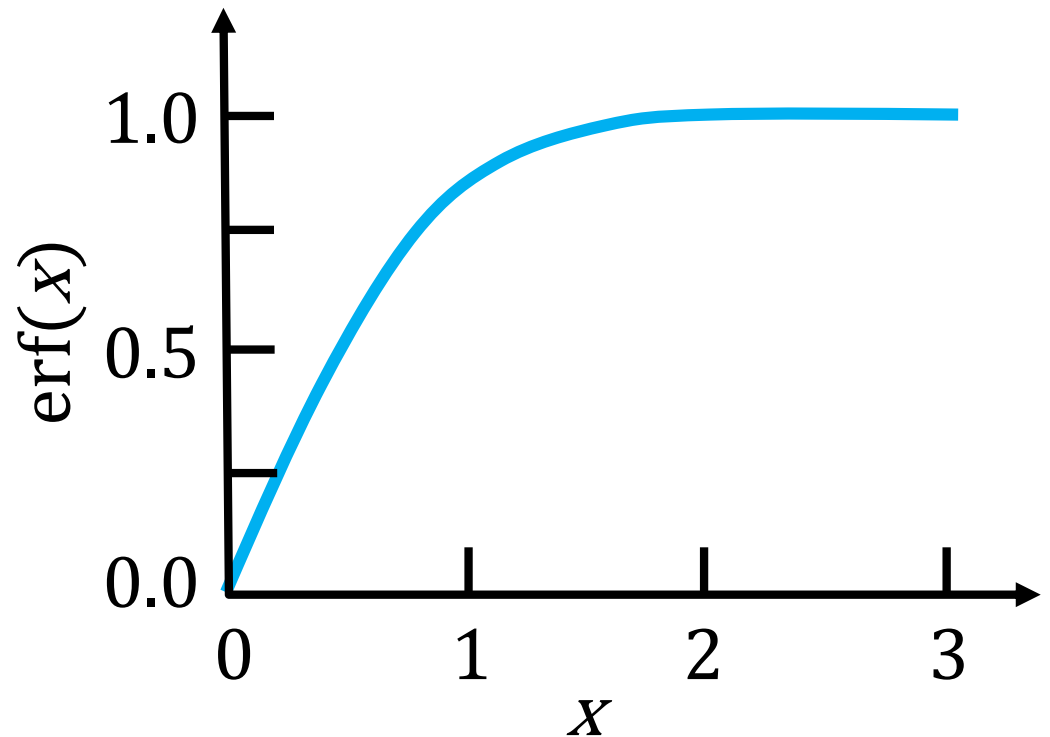
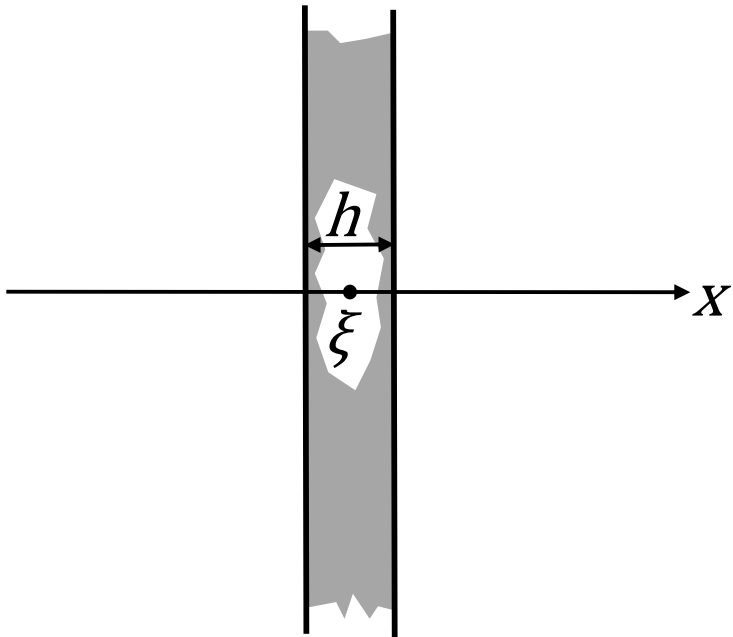
fitting of spectral peaks

Part 1

thermal diffusion

temperature in a cooling slab

$$T(x, t) = \frac{1}{2}T_0 \left\{ \operatorname{erf} \left[\frac{x - (\xi - \frac{1}{2}h)}{\sqrt{t}} \right] - \operatorname{erf} \left[\frac{x - (\xi + \frac{1}{2}h)}{\sqrt{t}} \right] \right\} = T_0 g(x, t, \xi)$$



temperature due to M cooling slabs
(use linear superposition)

$$T(x_i, t) = \sum_{j=1}^M g(x_i, t, \xi_j) m_j$$

temperature due to M slabs
each with initial temperature m_j


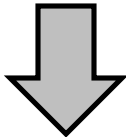
$$T(x_i, t) = \sum_{j=1}^M g(x_i, t, \xi_j) m_j$$

temperature
measured at
time $t > 0$

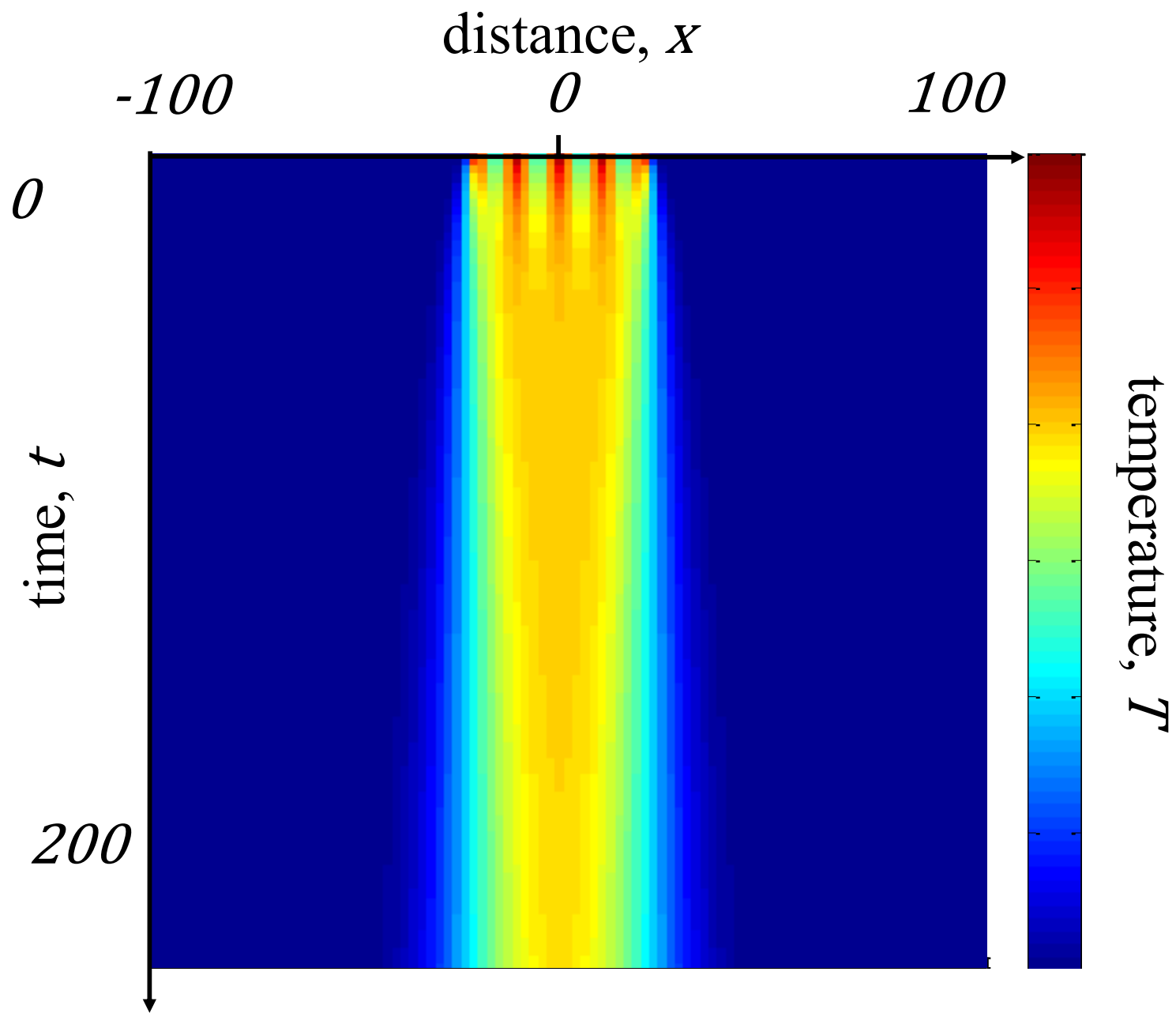
initial
temperature

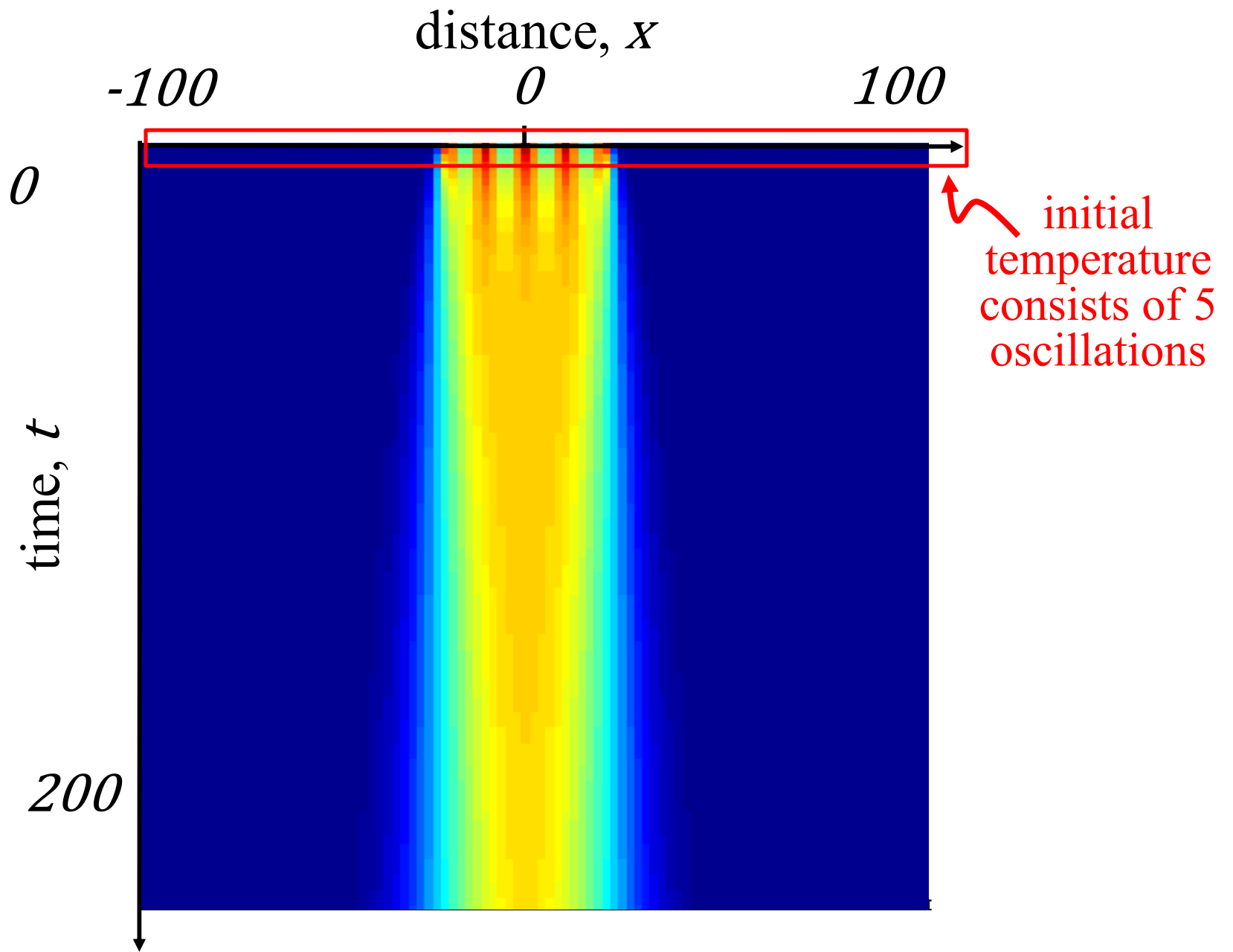
inverse problem
infer initial temperature \mathbf{m}
using temperatures measures at a suite of x s
at some fixed later time t

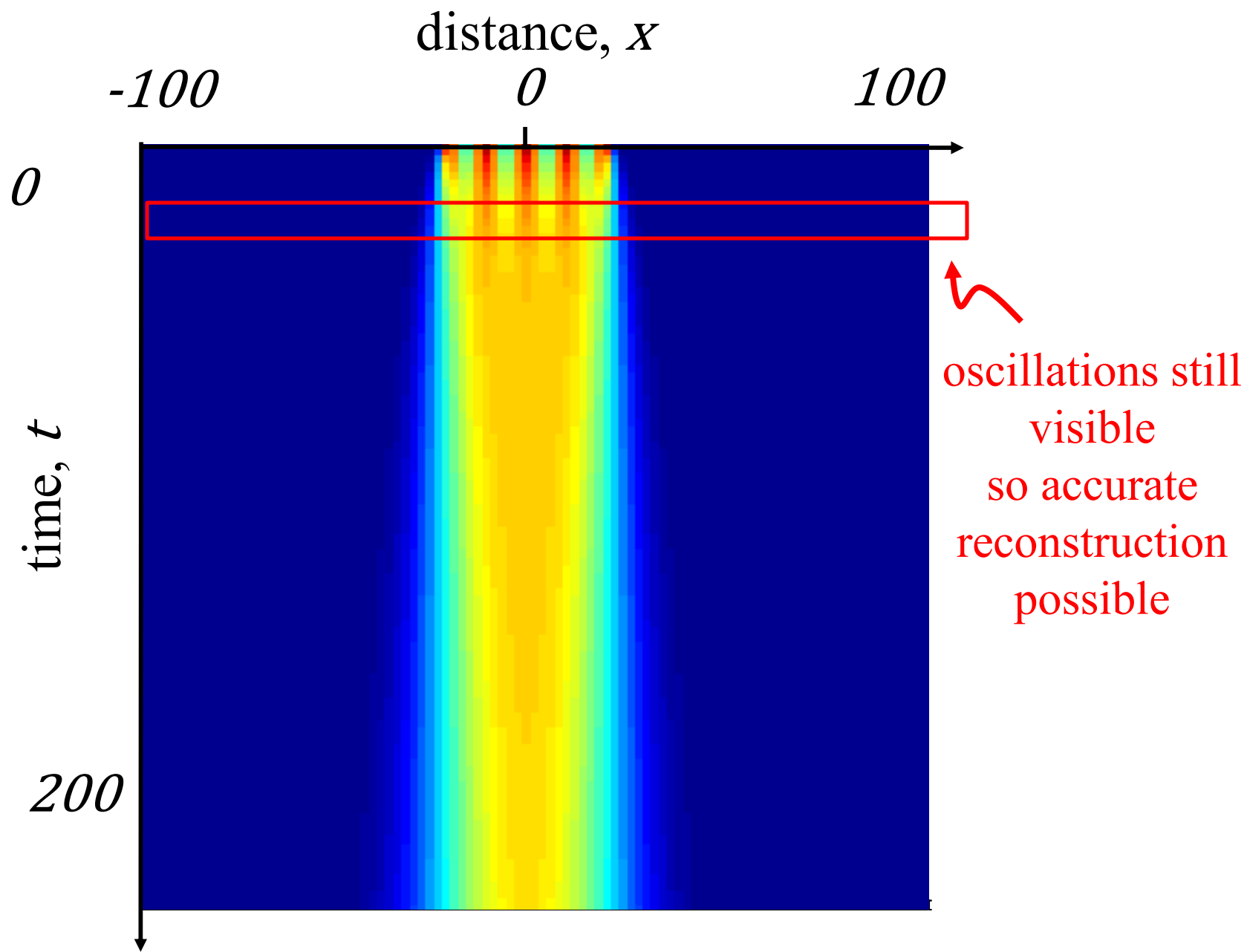
$$T(x_i, t) = \sum_{j=1}^M g(x_i, t, \xi_j) m_j$$

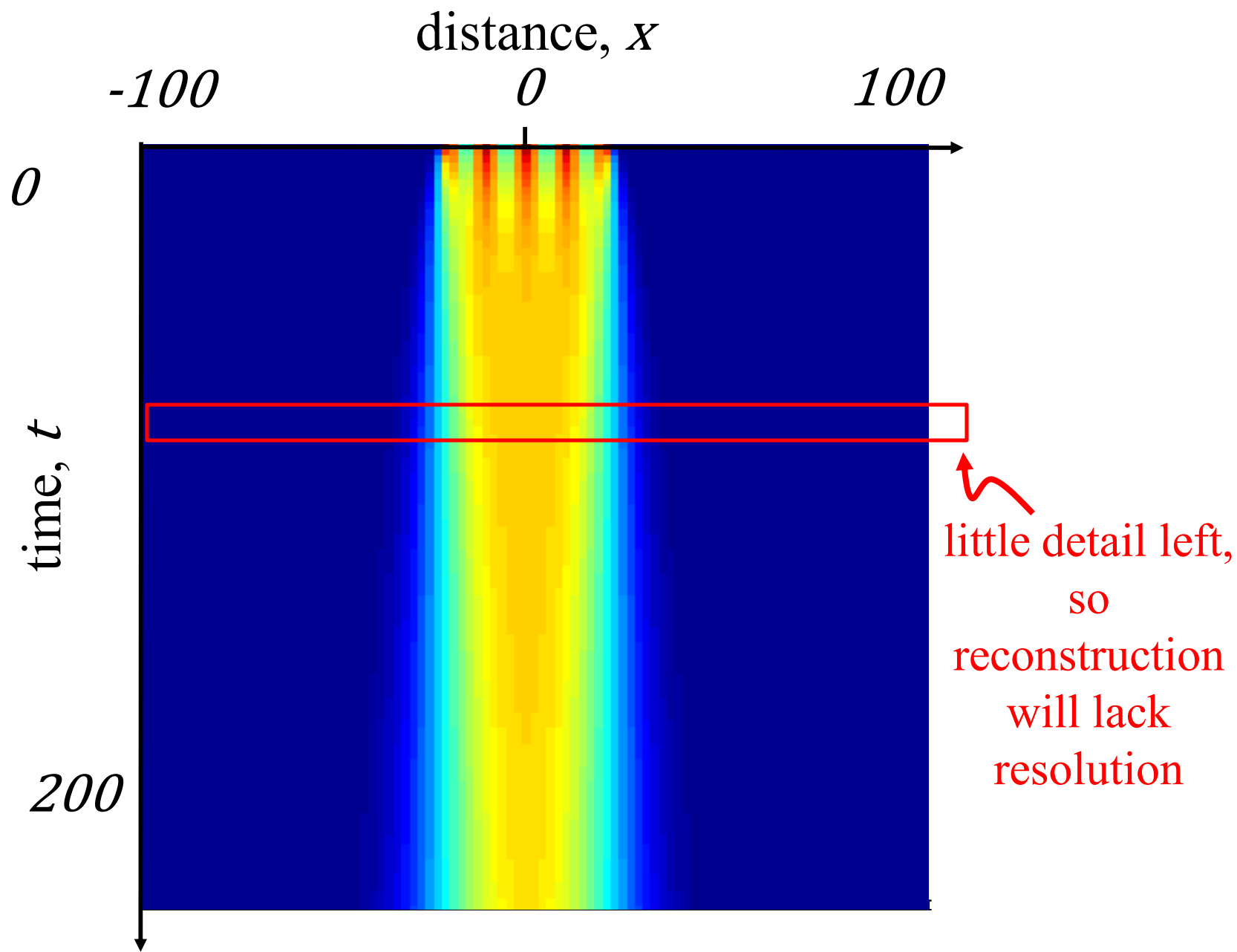
data   model parameters

$$\mathbf{d} = \mathbf{G} \mathbf{m}$$









What Method ?

The resolution is likely to be rather poor, especially when data are collected at later times

damped least squares

$$\mathbf{G}^{-g} = [\mathbf{G}^T \mathbf{G} + \varepsilon^2 \mathbf{I}]^{-1} \mathbf{G}^T$$

damped minimum length

$$\mathbf{G}^{-g} = \mathbf{G}^T [\mathbf{G} \mathbf{G}^T + \varepsilon^2 \mathbf{I}]^{-1}$$

Backus-Gilbert

What Method ?


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actually, these
generalized
inverses are
equal

Backus-Gilbert

What Method ?

The resolution is likely to be rather poor, especially when data are collected at later times

damped least squares

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damped minimum length

$$\mathbf{G}^{-g} = \mathbf{G}^T [\mathbf{G} \mathbf{G}^T + \varepsilon^2 \mathbf{I}]^{-1}$$

Backus-Gilbert



might produce solutions
with fewer artifacts

Try both

damped least squares

Backus-Gilbert

Solution Possibilities

1. Damped Least Squares:

Matrix \mathbf{G} is not sparse

no analytic version of $\mathbf{G}^T\mathbf{G}$ is available

$M=100$ is rather small

experiment with values of ε^2

```
mest=(G' *G+e2*eye (M,M) ) \ (G' *d)
```

2. Backus-Gilbert

use standard formulation, with damping α

experiment with values of α

```
GMG = zeros (M,N) ;
```

```
u = G*ones (M,1) ;
```

```
for k = [1:M]
```

```
    S = G * diag( ([1:M]-k).^2 ) * G' ;
```

```
    Sp = alpha*S + (1-alpha)*eye (N,N) ;
```

```
    uSpinV = u' / Sp ;
```

```
    GMG (k,:) = uSpinV / (uSpinV*u) ;
```

```
end
```

Solution Possibilities

1. Damped Least Squares:

Matrix \mathbf{G} is not sparse

no analytic version of $\mathbf{G}^T\mathbf{G}$ is available

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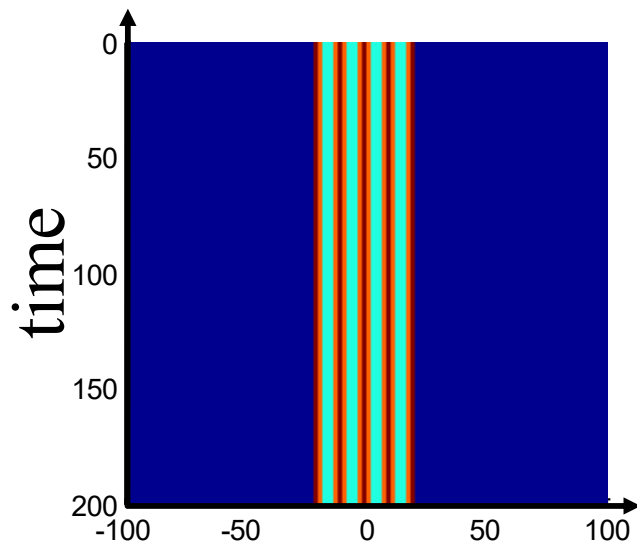
```
end
```



try both

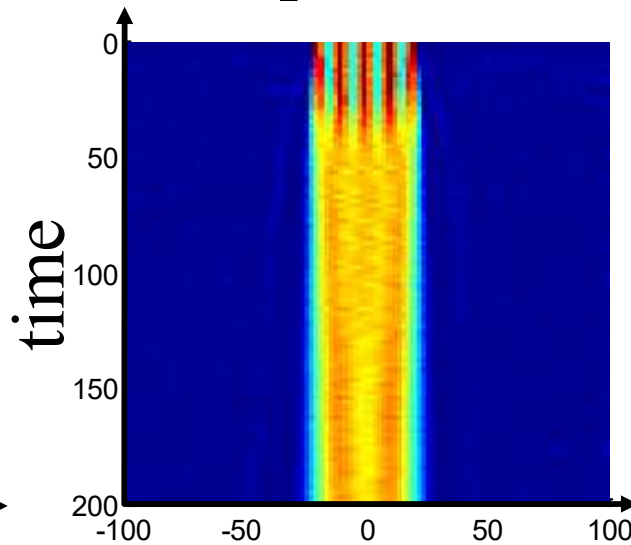
estimated initial temperature distribution as a function of the time of observation

True



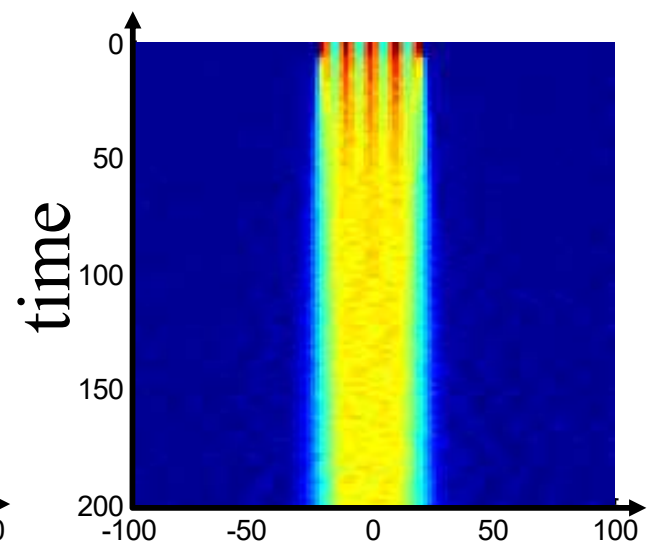
distance

Damped LS



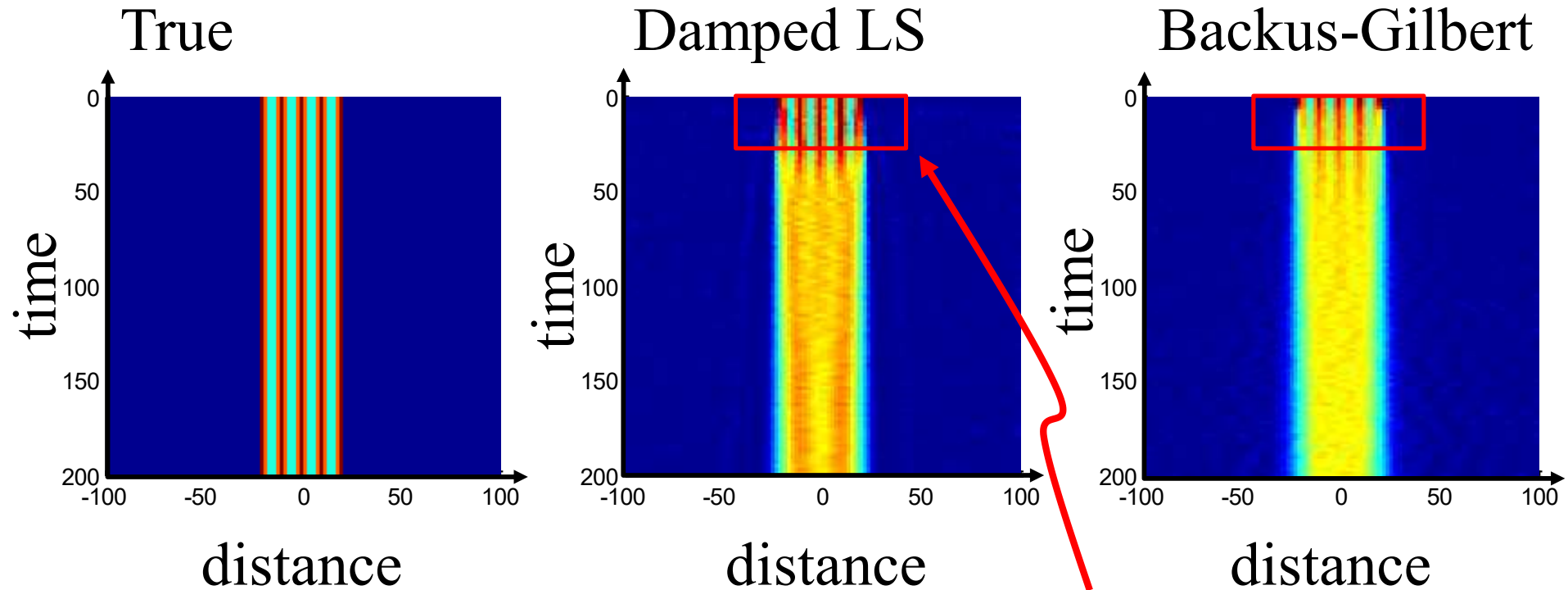
distance

Backus-Gilbert



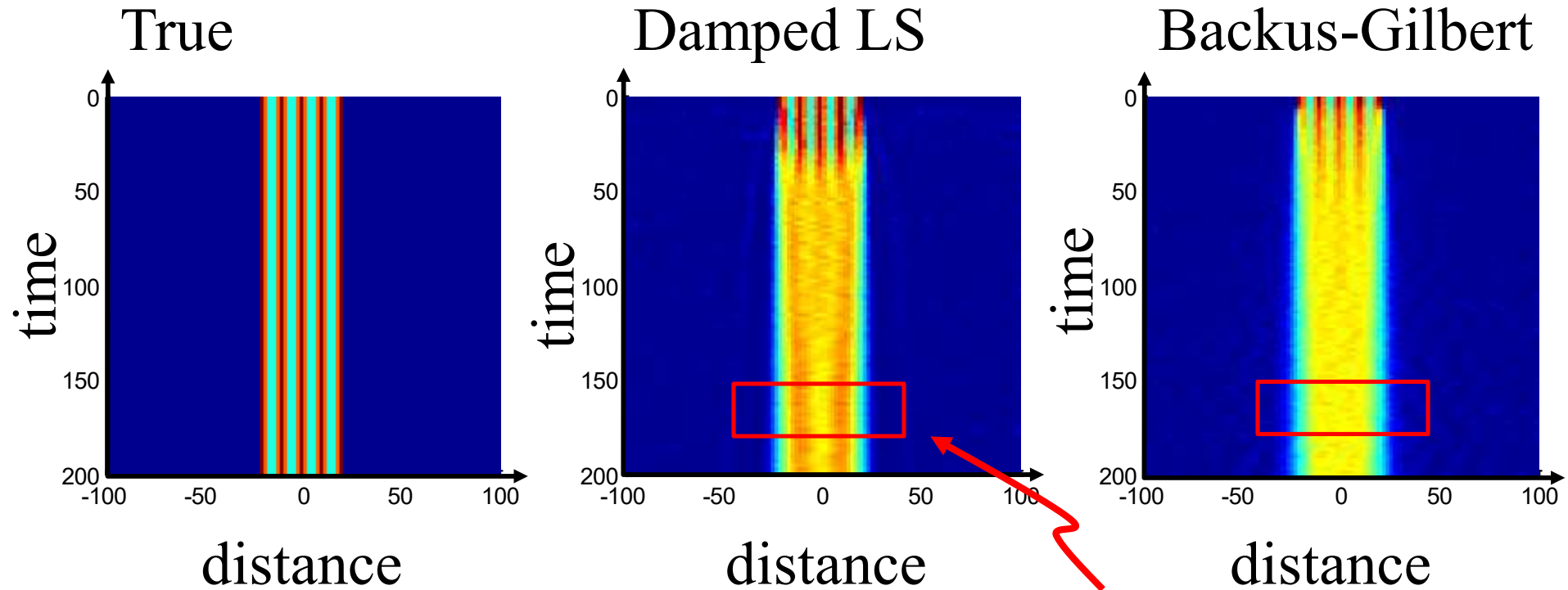
distance

estimated initial temperature distribution
as a function of the time of observation



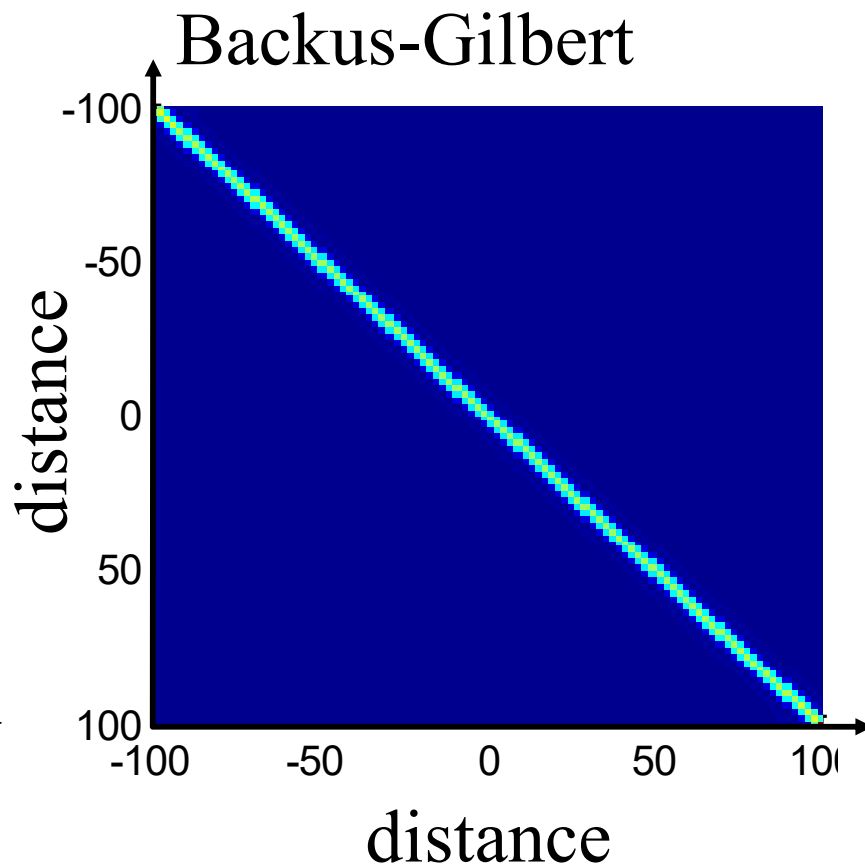
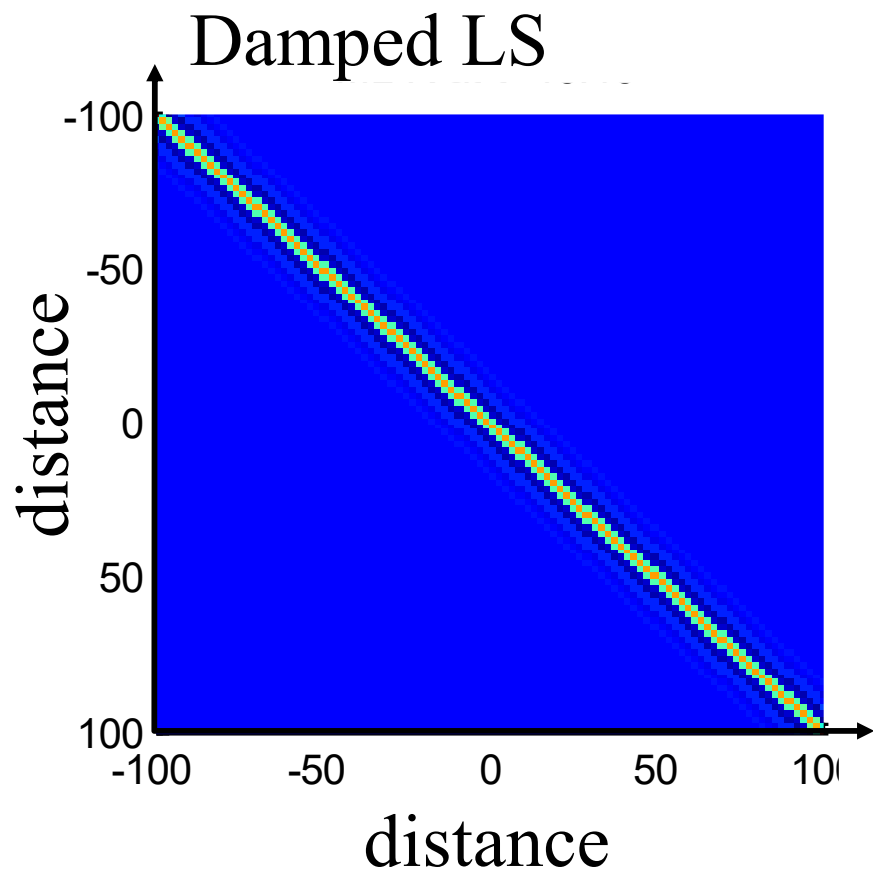
Damped LS
does better at
earlier times

estimated initial temperature distribution
as a function of the time of observation

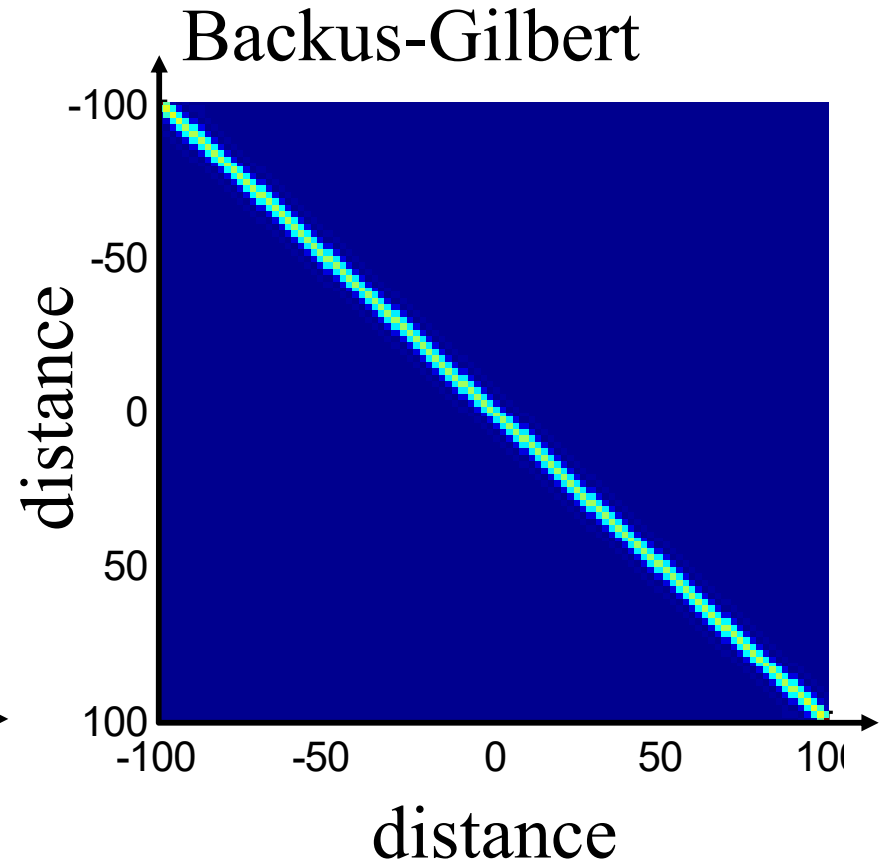
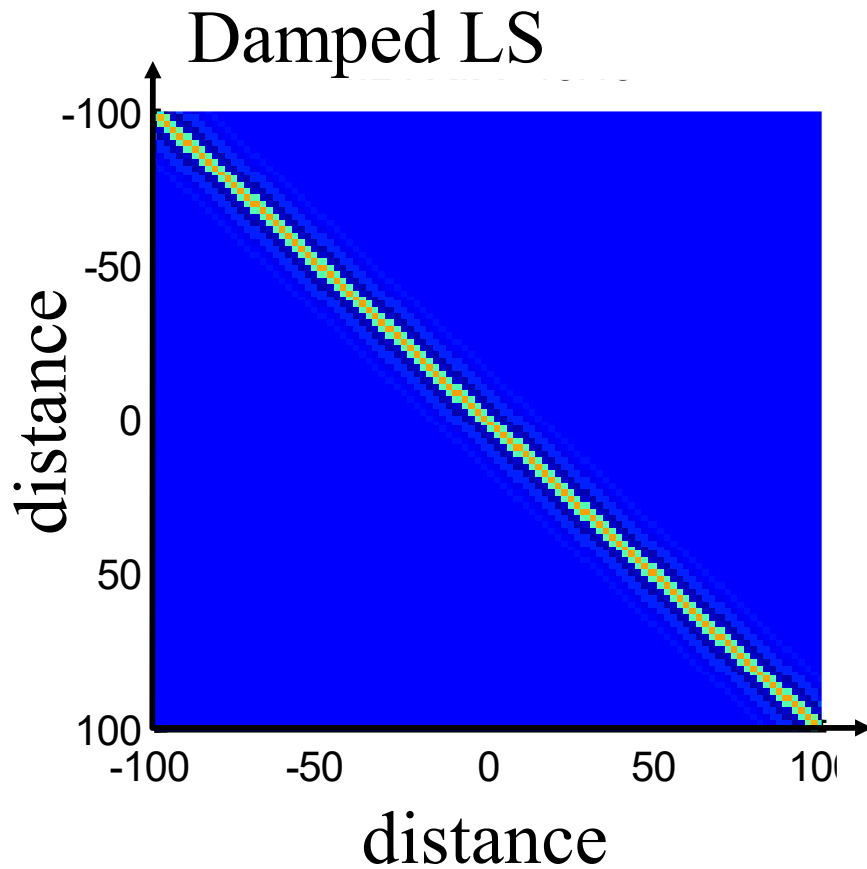


Damped LS
contains worse
artifacts at later
times

model resolution matrix when for data collected at $t=10$

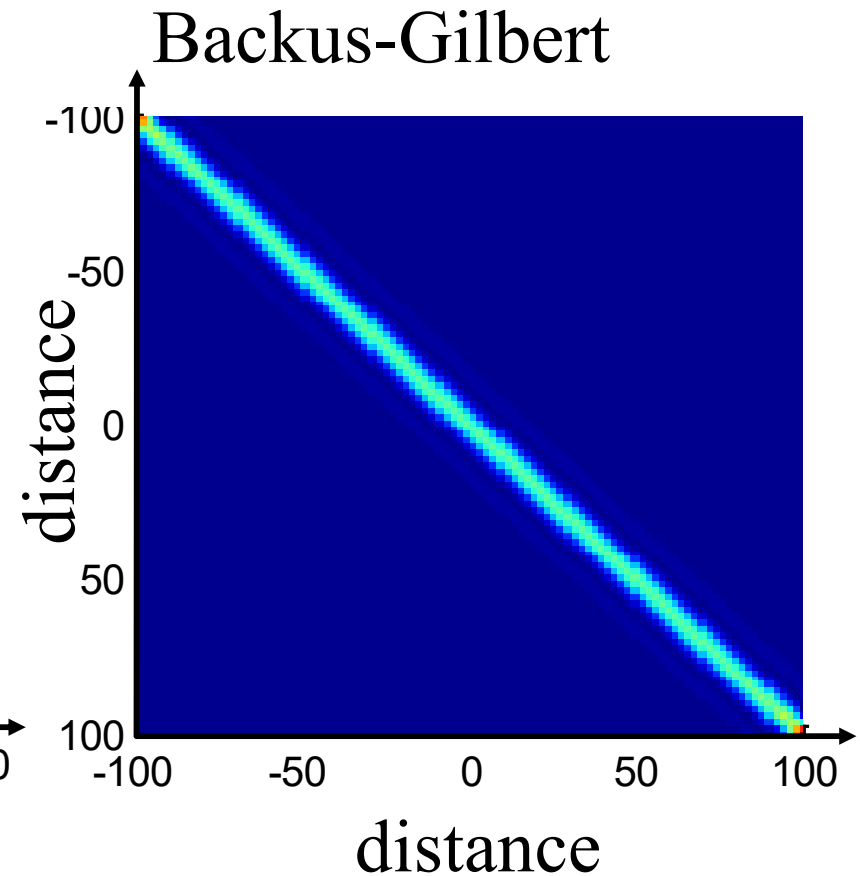
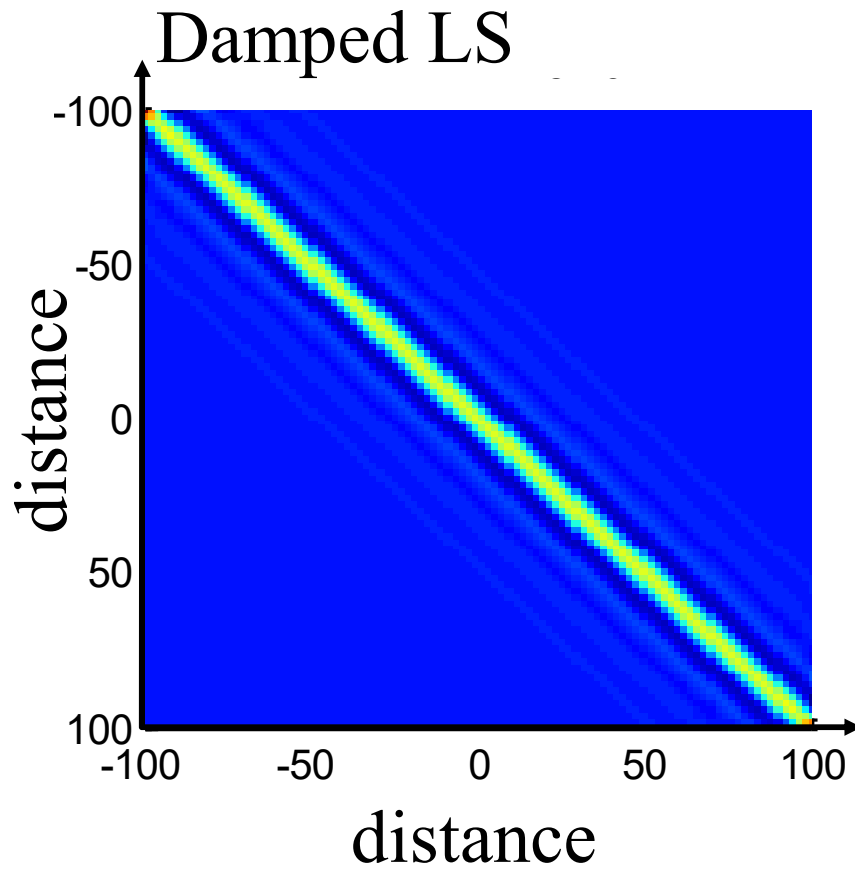


model resolution matrix when for data collected at $t=10$

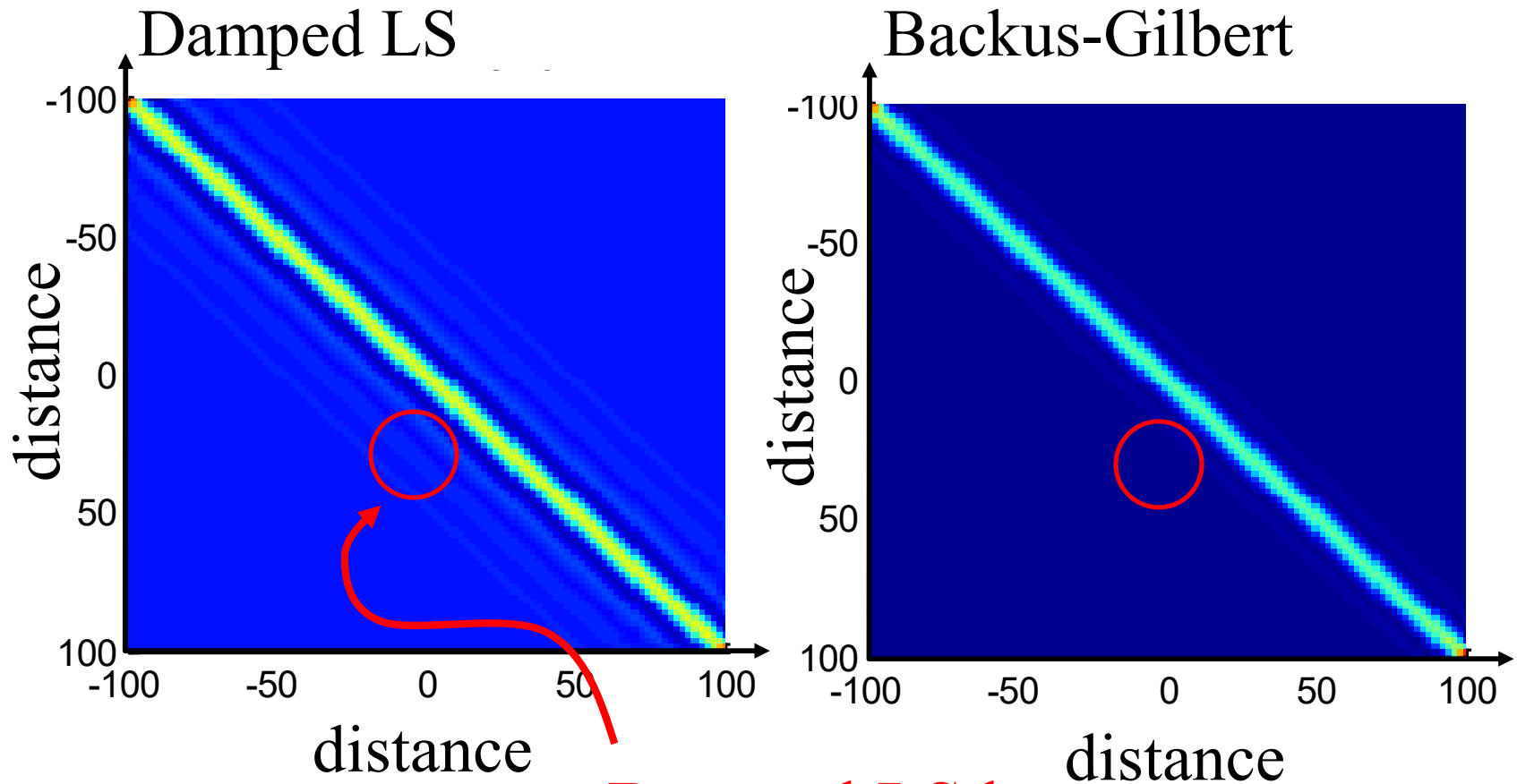


resolution
is similar

model resolution matrix when for data collected at $t=40$



model resolution matrix when for data collected at $t=40$



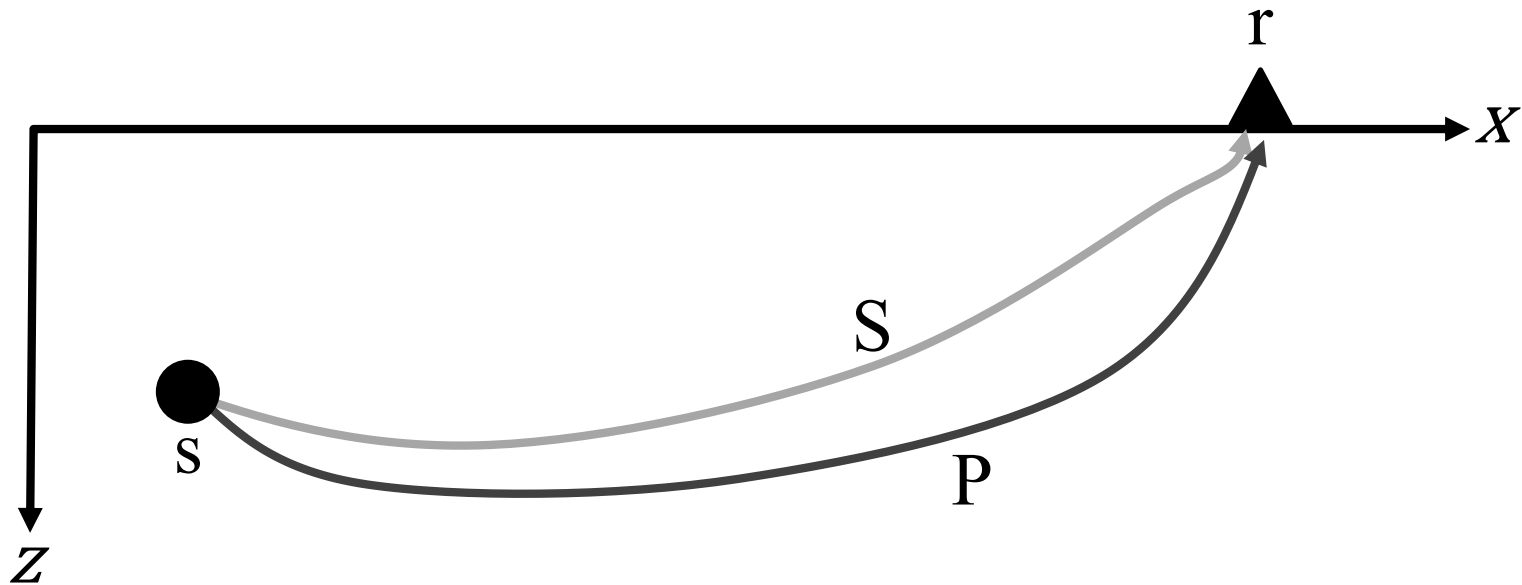
Damped LS has
much worse
sidelobes

Part 2

earthquake location

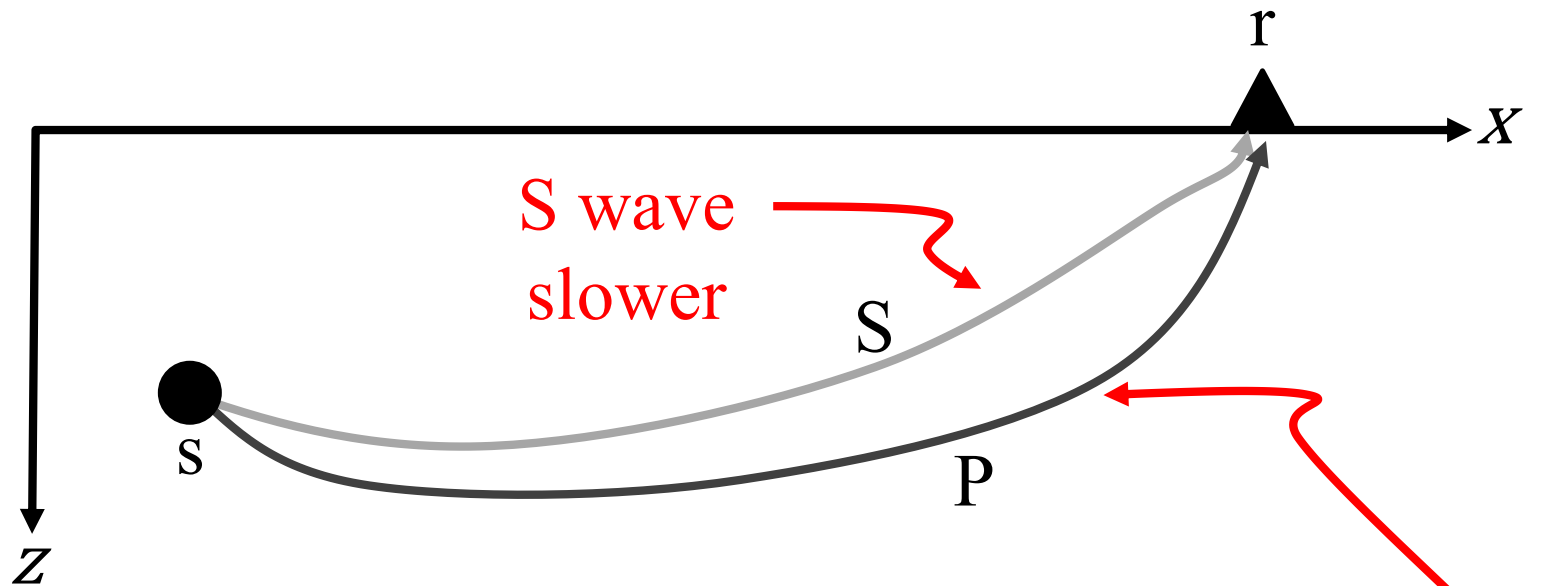
ray approximation

vibrations travel from source to receiver along
curved rays



ray approximation

vibrations travel from source to receiver along curved rays

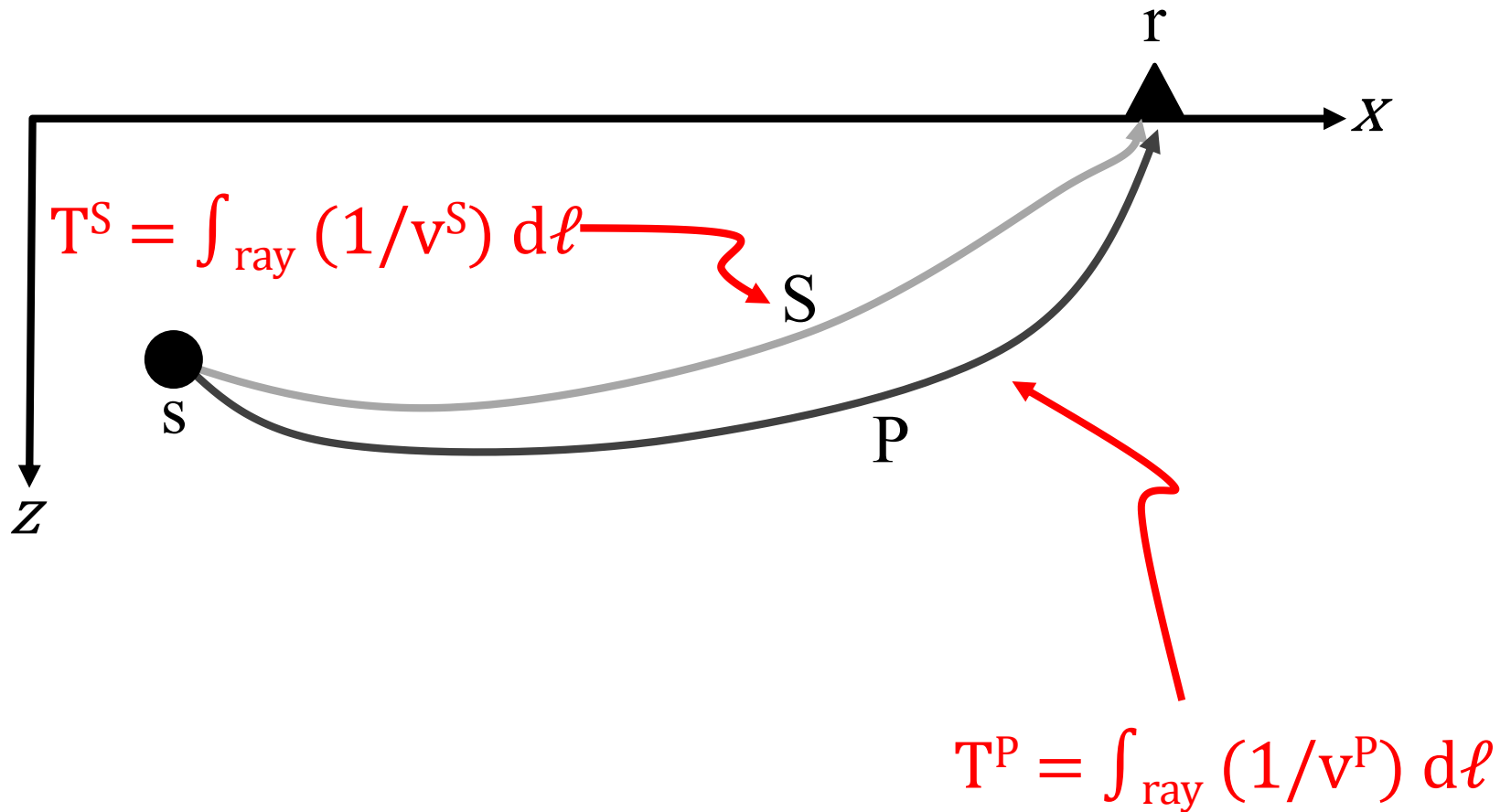


P, S ray paths not necessarily the same, but usually similar

P wave faster

travel time T

integral of slowness along ray path



arrival time = travel time along ray + origin time

$$t_i^P = T_i^P(\mathbf{x}^{(0)}, \mathbf{x}^{(i)}) + t_0 \quad \text{and} \quad t_i^S = T_i^S(\mathbf{x}^{(0)}, \mathbf{x}^{(i)}) + t_0$$

arrival time = travel time along ray + origin time

$$t_i^P = T_i^P(\mathbf{x}^{(0)}, \mathbf{x}^{(i)}) + t_0 \quad \text{and} \quad t_i^S = T_i^S(\mathbf{x}^{(0)}, \mathbf{x}^{(i)}) + t_0$$

data

data

earthquake
location
3 model
parameters

earthquake
origin time
1 model
parameter

arrival time = travel time along ray + origin time

$$t_i^P = T_i^P(\mathbf{x}^{(0)}, \mathbf{x}^{(i)}) + t_0 \quad \text{and} \quad t_i^S = T_i^S(\mathbf{x}^{(0)}, \mathbf{x}^{(i)}) + t_0$$

explicit nonlinear equation

4 model parameters
up to 2 data per station

arrival time = travel time along ray + origin time

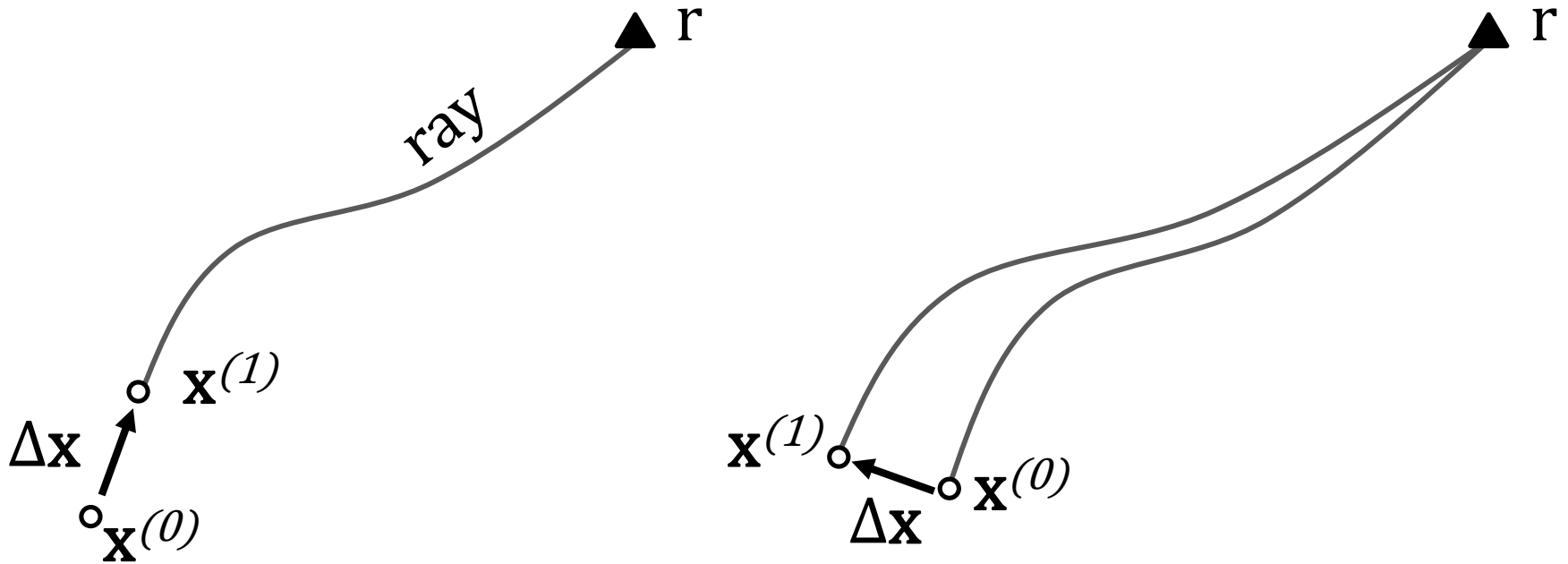
$$t_i^P = T_i^P(\mathbf{x}^{(0)}, \mathbf{x}^{(i)}) + t_0$$

linearize around trial
source location $\mathbf{x}^{(p)}$

$$t_i^P = T_i^P(\mathbf{x}^{(p)}, \mathbf{x}^{(i)}) + [\nabla T_i^P] \cdot \Delta \mathbf{x} + t_0$$

trick is computing this gradient

Geiger's principle



$$[\nabla T_i^P] = -\mathbf{s}/v$$

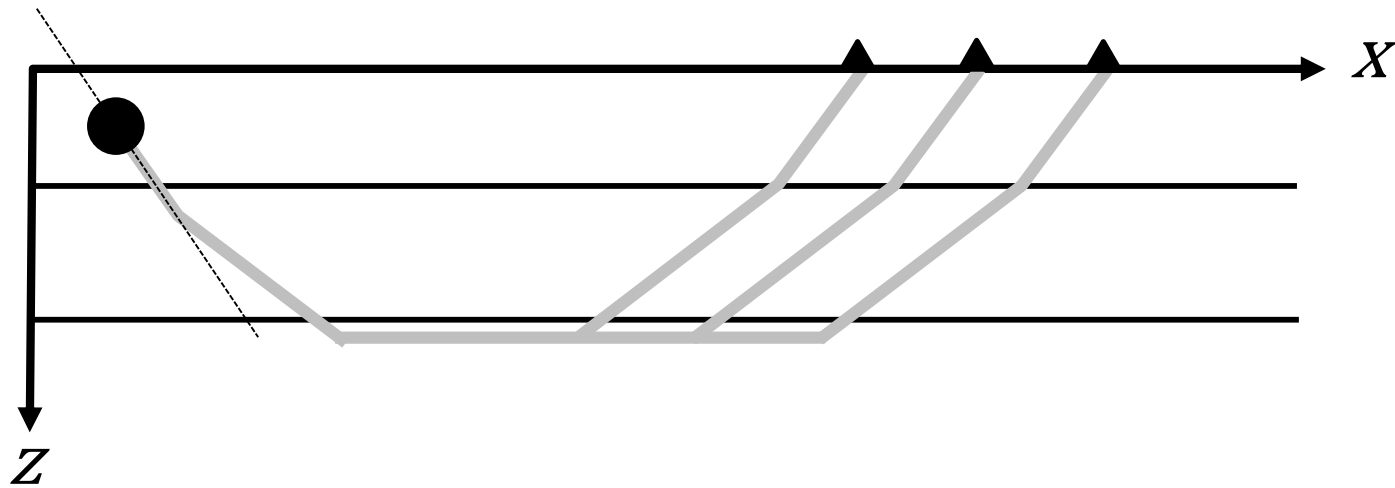
unit vector parallel to ray pointing
away from receiver

linearized equation

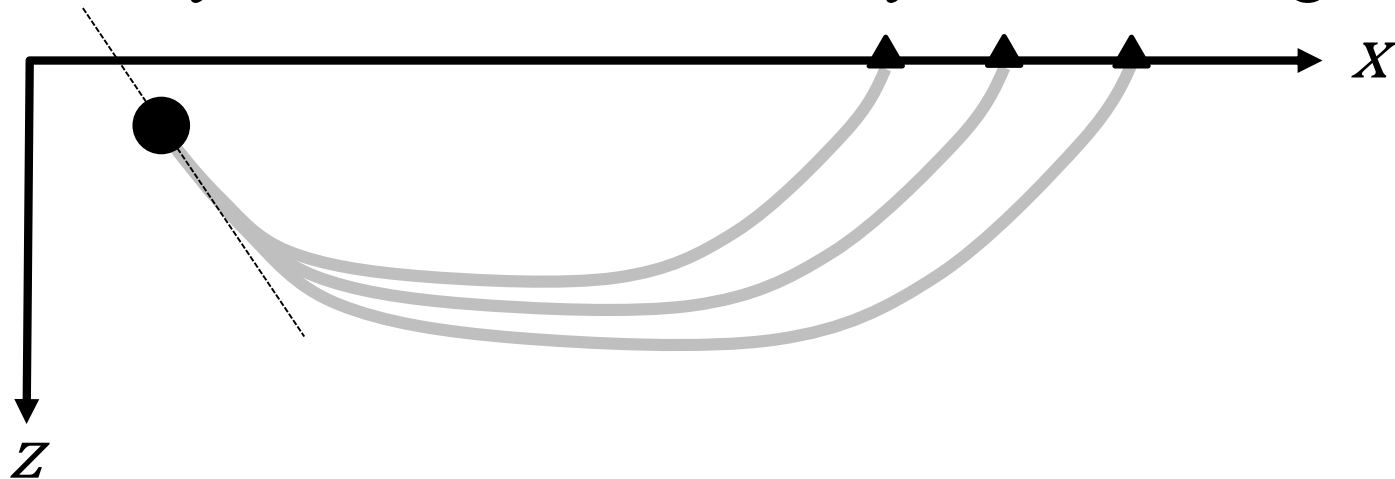
$$\begin{bmatrix} -s_1^{(1)P} / v_0^P & -s_2^{(1)P} / v_0^P & -s_3^{(1)P} / v_0^P & 1 \\ \dots & \dots & \dots & \dots \\ -s_1^{(N)P} / v_0^P & -s_2^{(N)P} / v_0^P & -s_3^{(N)P} / v_0^P & 1 \\ -s_1^{(1)S} / v_0^S & -s_2^{(1)S} / v_0^S & -s_3^{(1)S} / v_0^S & 1 \\ \dots & \dots & \dots & \dots \\ -s_1^{(N)S} / v_0^S & -s_2^{(N)S} / v_0^S & -s_2^{(N)S} / v_0^S & 1 \end{bmatrix} \begin{bmatrix} \Delta x_0 \\ \Delta y_0 \\ \Delta z_0 \\ t_0 \end{bmatrix} = \begin{bmatrix} t_1^P - T_1^P \\ \dots \\ t_N^P - T_N^P \\ t_1^S - T_1^S \\ \dots \\ t_N^S - T_N^S \end{bmatrix}$$

Common circumstances when earthquake far from stations

All rays leave source at the same angle



All rays leave source at nearly the same angle



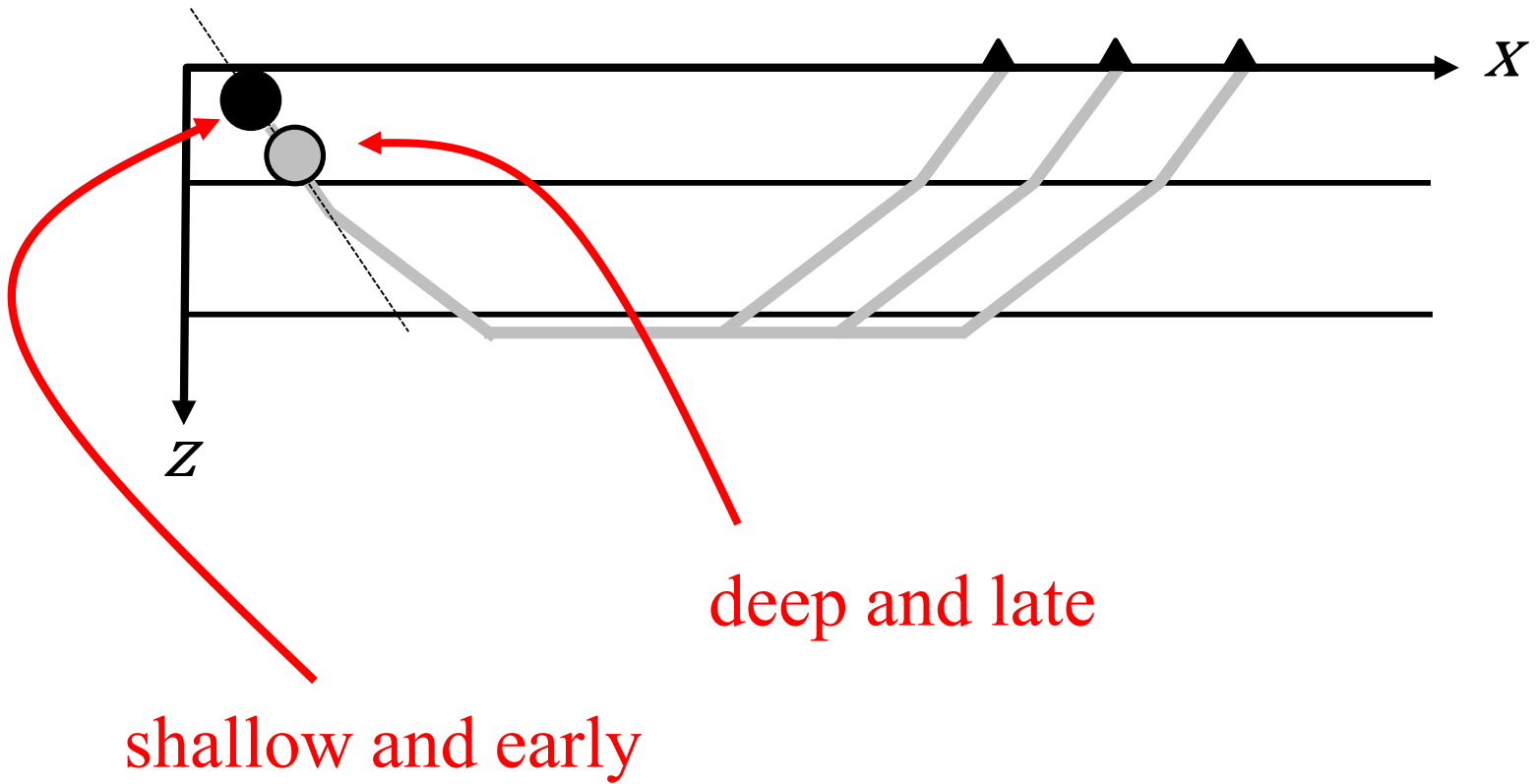
then, if only P wave data is available

$$\begin{bmatrix}
 -s_1^{(1)P} / v_0^P & -s_2^{(1)P} / v_0^P & -s_3^{(1)P} / v_0^P & 1 \\
 \dots & \dots & \dots & \dots \\
 -s_1^{(N)P} / v_0^P & -s_2^{(N)P} / v_0^P & -s_3^{(N)P} / v_0^P & 1 \\
 \vdots & \vdots & \vdots & \vdots \\
 \vdots & \vdots & \vdots & \vdots
 \end{bmatrix}
 \begin{bmatrix}
 \Delta x_0 \\
 \Delta y_0 \\
 \Delta z_0 \\
 t_0
 \end{bmatrix}
 =
 \begin{bmatrix}
 t_1^P - T_1^P \\
 \dots \\
 t_N^P - T_N^P \\
 t_1^S - T_1^S \\
 \dots \\
 t_N^S - T_N^S
 \end{bmatrix}$$

(no S waves)

these two columns are
proportional to one-another

depth and origin time trade off



Solution Possibilities

1. Damped Least Squares:

Matrix \mathbf{G} is not sparse

no analytic version of $\mathbf{G}^T\mathbf{G}$ is available

$M=4$ is tiny

experiment with values of ε^2

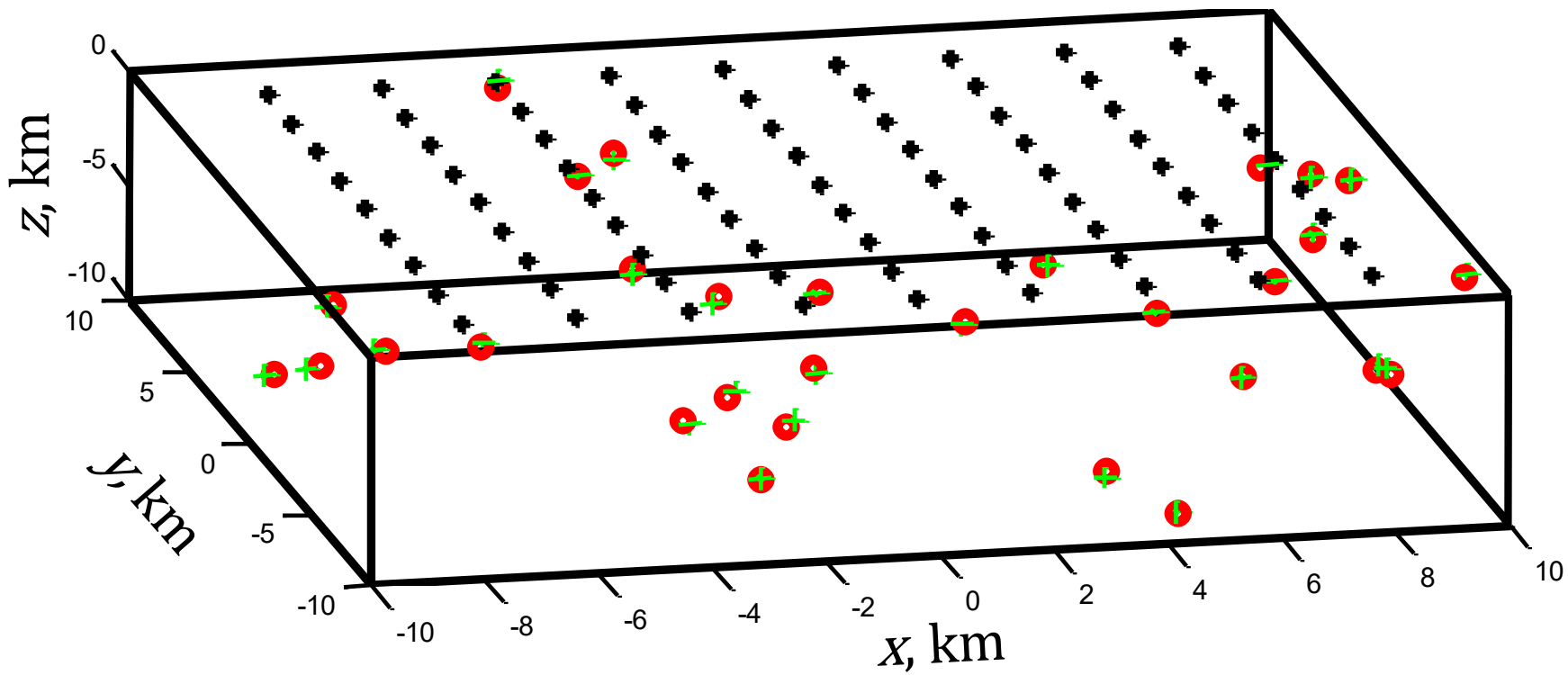
$\text{mest} = (\mathbf{G}' * \mathbf{G} + \varepsilon^2 * \text{eye}(M, M)) \setminus (\mathbf{G}' * \mathbf{d})$



test case has
earthquakes
“inside of array”

2. Singular Value Decomposition

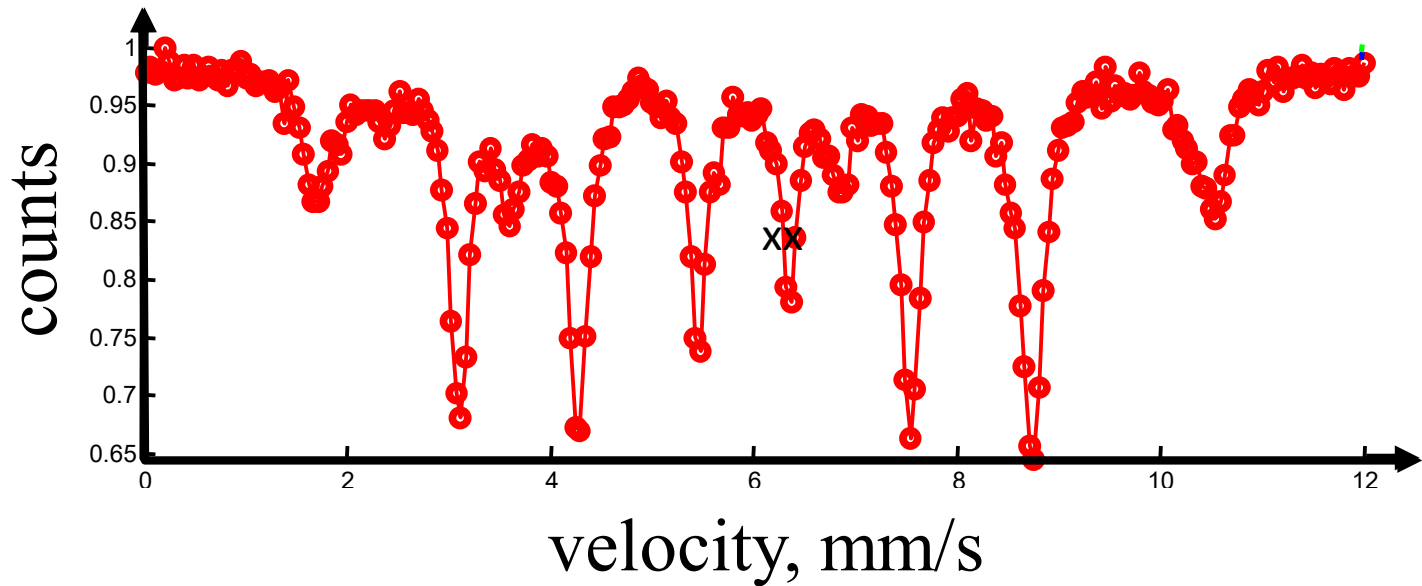
to detect case of depth and origin time trading off



Part 3

fitting of spectral peaks

typical spectrum consisting of overlapping peaks



what shape are the peaks?

Gaussian:
$$d_i = \sum_{j=1}^q \frac{A_j}{(2\pi)^{1/2} c_j} \exp \left[-\frac{(z_i - f_j)^2}{2c_j^2} \right]$$

Lorentzian:
$$d_i = \sum_{j=1}^q \frac{A_j c_j^2}{(z_i - f_j)^2 + c_j^2}$$

what shape are the peaks?

Gaussian:
$$d_i = \sum_{j=1}^q \frac{A_j}{(2\pi)^{1/2} c_j} \exp \left[-\frac{(z_i - f_j)^2}{2c_j^2} \right]$$

Lorentzian:
$$d_i = \sum_{j=1}^q \frac{A_j c_j^2}{(z_i - f_j)^2 + c_j^2}$$

try both
use F test to test whether one is better than the other

what shape are the peaks?

data \curvearrowright

Gaussian: $d_i = \sum_{j=1}^q \frac{A_j}{(2\pi)^{1/2} c_j} \exp \left[-\frac{(z_i - f_j)^2}{2c_j^2} \right]$

3 unknowns per peak

data \curvearrowright

Lorentzian: $d_i = \sum_{j=1}^q \frac{A_j c_j^2}{(z_i - f_j)^2 + c_j^2}$

3 unknowns per peak

both cases:
explicit nonlinear problem

linearize using analytic gradient

Gaussian:

$$\partial g_i / \partial A_j = [1 / (2\pi)^{1/2} c_j] \exp \left[-(z_i - f_j)^2 / 2c_j^2 \right]$$

$$\partial g_i / \partial f_j = \left[\frac{A_j}{(2\pi)^{1/2} c_j} \right] \left[(z_i - f_j) / c_j^2 \right] \exp \left[-(z_i - f_j)^2 / 2c_j^2 \right]$$

$$\partial g_i / \partial c_j = \left[\frac{A_j}{(2\pi)^{1/2} c_j^2} \right] \left[\left((z_i - f_j)^2 / c_j^2 \right) - 1 \right] \exp \left[-(z_i - f_j)^2 / 2c_j^2 \right]$$

linearize using analytic gradient

Lorentzian:

$$\partial g_i / \partial A_j = c_j^2 / [(z_i - f_j)^2 + c_j^2]$$

$$\partial g_i / \partial f_j = 2A_j c_j^2 (z_i - f_j) / [(z_i - f_j)^2 + c_j^2]^2$$

$$\partial g_i / \partial c_j = 2A_j c_j / [(z_i - f_j)^2 + c_j^2] - 2A_j c_j^3 / [(z_i - f_j)^2 + c_j^2]^2$$

issues

how to determine

number q of peaks

parameters A_i c_i f_i of each peak

our solution

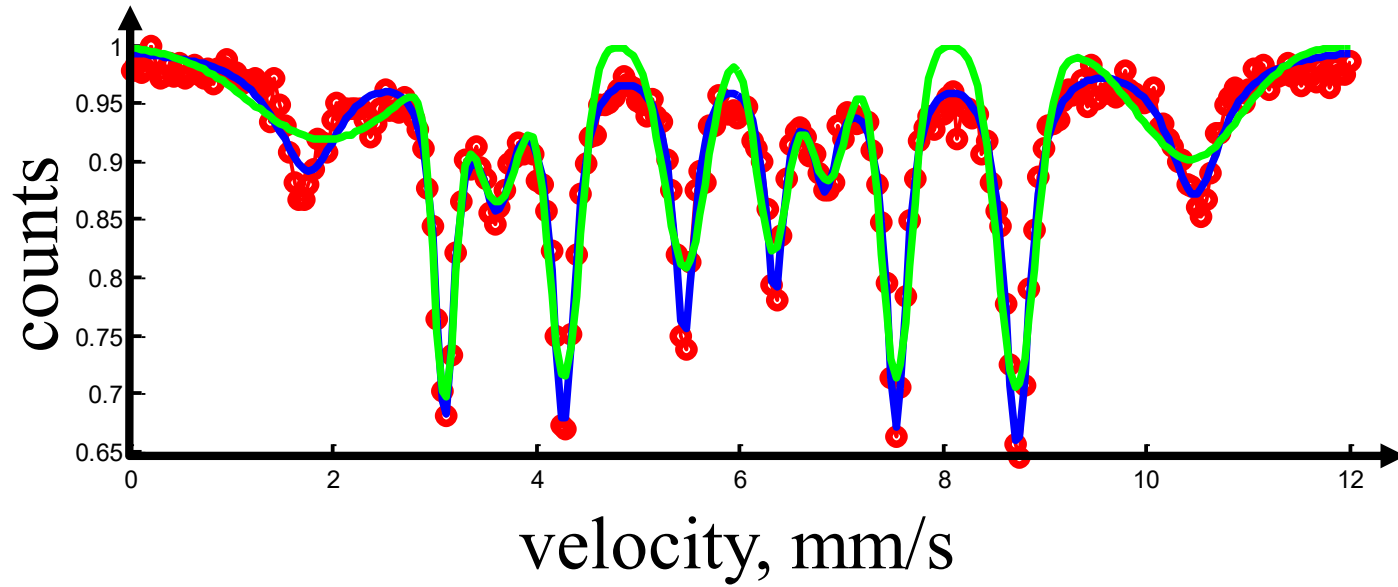
have operator click mouse

computer screen

to indicate position of each peak

MatLab code for graphical input

```
K=0;  
for k = [1:20]  
    p = ginput(1);  
    if( p(1) < 0 )  
        break;  
    end  
    K=K+1;  
    a(K) = p(2)-A;  
    v0(K)=p(1);  
    c(K)=0.1;  
end
```



— Lorentzian
— Gaussian

Results of F test

Fest = E_normal/E_lorentzian: 4.230859
P(F<=1/Fest || F>=Fest) = 0.000000

Lorentzian better fit
to 99.9999% certainty