Receiven Function times. Time along ray 13 just x.p Htang PP ps ray Px=sud VA H 0 H Pz= coso H = Htowd Px+HPz = Hsmit + H coso Vp coso + H coso = H (Sind + cosid)  $T_P = \frac{A}{V_P} = X_A R_X + H R_2^P$ - H Vpcost  $T_{S} = \frac{B}{V_{S}} + (XA - YB)'PX$ correction = XBPx + HP2 + XAPx - XBP4 for different rags = HPz + XAPX & just difference of vertical scources. TS-Tp = AP2 - HP2 proof That "vertical slowness method" is equivalent to a ray calculation That includes the effect of the PP and PS rays being different.

Derivation of formulae for predicting times of P-to-S converted phases from a single interface. Author – Bill Menke. Typed up by Vadim Levin

Travel time along the ray is a vector product of the slowness vector **p** and distance vector **x**. Distance vector for the PP ray is **x**=(XA,H), slowness vector for the PP wave is  $\mathbf{p}=(px, pz)$ . Horizontal slowness of a PP wave  $\theta_{\mathsf{P}}$ in the upper layer:  $p_x = \frac{\sin\theta_P}{V_P}$ Ηc Vertical slowness of a PP wave in the upper layer:  $X_B$  $p_{z}^{p} = \frac{\cos\theta_{P}}{V_{P}}$ Length of the PP ray in the upper layer is

$$A = H/cos\theta_P$$

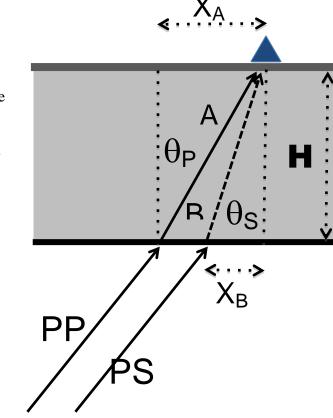
Travel time  $T_P$  along ray PP that is a vector product **p**·**x** evaluates to

$$T_P = X_A p_x + H p_z^p = H tan \theta_P p_x + H p_z^p$$

For PS ray the horizontal slowness  $p_x = \frac{\sin\theta_S}{V_S} = \frac{\sin\theta_P}{V_P}$  (Snell's Law), while the vertical slowness is  $p_z^S = \frac{\cos\theta_S}{V_S}$ .

Travel time for the PS phase will include the travel along ray B in the crust, and a fraction of travel along the ray below the boundary that propagates there while PP wave has already crossed into the crust:

$$T_S = \frac{B}{V_S} + (X_A - X_B)p_x$$



$$= X_B p_x + H p_z^S + X_A p_x - X_B p_x$$
$$= H p_z^S + X_A p_x$$

Consequently, the difference in travel time (the *delay* of the phase) in a receiver function will be

 $T_P - T_S = Hp^S_{\ z} - Hp^P_{\ z}$