Seasonal fluxes of carbonyl sulfide in a midlatitude forest

Róisín Commanea,b,1, Laura K. Meredithc,2, Ian T. Bakerd, Joseph A. Berrye, J. William Mungera,b, Joseph A. Montzka,1, Pamela H. Templera, Stephanie M. Juicegb, Mark S. Zahniserb, and Steven C. Wofsb,2

OCS have linked OCS fluxes and the photosynthetic uptake of OCS throughout an annual cycle, at a mixed temperate forest.

CA 94305; Global Monitoring Division, Earth System Research Laboratory, National Oceanic and Atmospheric Administration, Boulder, CO 80305; Department of Biology, Boston University, Boston, MA 02215; and Center for Atmospheric and Environmental Chemistry, Aerodyne Research Inc., Billerica, MA 01821

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Carbonyl sulfide (OCS), the most abundant sulfur gas in the atmosphere, has a summer minimum associated with uptake by vegetation and soils, closely correlated with CO2. We report the first direct measurements to our knowledge of the ecosystem flux of OCS throughout an annual cycle, at a mixed temperate forest. The forest took up OCS during most of the growing season with an overall uptake of 1.36 ± 0.01 mol OCS per ha (43.5 ± 0.5 g S per ha, 95% confidence intervals) for the year. Daytime fluxes accounted for 72% of total uptake. Both soils and incompletely closed stomata in the canopy contributed to nighttime fluxes. Unexpected net OCS emission occurred during the warmest weeks in summer. Many requirements necessary to use fluxes of OCS as a simple estimate of photosynthesis were not met because OCS fluxes did not have a constant relationship with photosynthesis throughout an entire day or over the entire year. However, OCS fluxes provide a direct measure of ecosystem-scale stomatal conductance and mesophyll function, without relying on measures of soil evaporation or leaf temperature, and reveal previously unseen heterogeneity of forest canopy processes. Observations of OCS flux provide powerful, independent means to test and refine land surface and carbon cycle models at the ecosystem scale.

carbonyl sulfide | carbon cycle | sulfur cycle | stomatal conductance

Carbonyl sulfide (OCS) is the most abundant sulfur gas in the atmosphere (1), and biogeochemical cycling of OCS affects both the stratosphere and the troposphere. The tropospheric OCS mixing ratio is between 300 and 550 parts per trillion (ppt) (1) (10^{-12} mol OCS per mol dry air), decreasing sharply with altitude in the stratosphere (2). In times of low volcanic activity, the sulfur budget and aerosol loading of the stratosphere are largely controlled by transport and photooxidation of OCS from the troposphere (3). The processes regulating emission and uptake of OCS are thus important factors in determining how changes in climate and land cover may affect the stratospheric sulfate layer.

Oceans are the dominant source of atmospheric OCS (4), with smaller emissions from anthropogenic and terrestrial sources, such as wetlands and anoxic soils (e.g., refs. 5 and 6) and oxic soils during times of heat or drought stress (e.g., refs. 7 and 8). The terrestrial biosphere is the largest sink for OCS (1, 4, 9, 10) with uptake by both oxic soils (e.g., ref. 11) and vegetation (e.g., ref. 9). Once OCS molecules pass through the stomata of leaves, the uptake rate of OCS is controlled by reaction with carbonic anhydrase (CA) within the mesophyll, to produce H2S and CO2. CA is the same enzyme that hydrolyzes carbon dioxide (CO2) in the first chemical step of photosynthesis (12).

Studies considering the large-scale atmospheric variability of OCS have linked OCS fluxes and the photosynthetic uptake of CO2 for regional and global scales (1, 4, 13). Leaf-scale studies have confirmed the OCS link to photosynthesis (14, 15). Initial OCS ecosystem flux estimations were made using flask sampling followed by analysis via gas chromatography–mass spectrometry (GC-MS) (13, 16), but these studies did not have sufficient resolution to examine daily or hourly controls on the OCS flux. Laser spectrometers have been developed (17, 18) to enable direct, in situ measurement of OCS fluxes by eddy covariance, and measurements of OCS ecosystem fluxes have been reported, for periods of up to a few weeks, above arid forests (19) and an agricultural field (8, 20).

Net carbon exchange in terrestrial ecosystems [net ecosystem exchange (NEE)] can be measured by eddy flux methods. NEE may be regarded as the sum of two gross fluxes: gross ecosystem productivity (GEP) and ecosystem respiration (Reco). GEP is the light-dependent part of NEE, estimated by subtracting daytime ecosystem respiration (Reco), computed by extrapolation of the temperature dependence of nighttime NEE (NEE − Reco = GEP) (e.g., refs. 21–24). At night, NEE includes all autotrophic and heterotrophic respiration processes. During the day, GEP approximates the carboxylation rate minus photorespiration at the ecosystem scale (25). Extrapolation of nighttime Reco introduces major uncertainty in the interpretation of GEP, which could be reduced, and the ecological significance of GEP increased, by developing independent methods of measuring rates of photosynthetic processes. As shown below, fluxes of OCS give more

Significance

The flux of carbonyl sulfide (OCS) provides a quantitative, independent measure of biospheric activity, especially stomatal conductance and carbon uptake, at the ecosystem scale. We describe the factors controlling the hourly, daily, and seasonal fluxes of OCS based on 1 year of observations in a forest ecosystem. Vegetation dominated uptake of OCS, with daytime fluxes accounting for 72% of the total uptake for the year. Nighttime fluxes had contributions from both incompletely closed stomata and soils. Net OCS emission was observed at high temperature in summer. Diurnal and seasonal variations in OCS flux show variable stoichiometry relative to photosynthetic uptake of CO2. An effective model framework is shown, using an explicit representation of ecosystem processing of OCS.


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Data deposition: The data have been deposited in the Harvard Forest Data Archive, harvardforest.fas.harvard.edu:8080/eisweb/app/datasets/showData.html?id=hf214.

1To whom correspondence should be addressed. Email: rcommane@seas.harvard.edu.

2Present address: Stanford School of Earth, Energy & Environmental Sciences, Stanford University, Stanford, CA 94305.

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direct information on one of the major controls on GEP, stomatal conductance, rather than GEP itself, providing a powerful means for testing and improving ecosystem models and for scaling up leaf-level processes to the whole ecosystem.

Here we describe the factors controlling the hourly, daily, seasonal, and total fluxes of OCS in a forest ecosystem, using a year (2011) of high-frequency, direct measurements at Harvard Forest, MA. We report the seasonal cycle, the response to environmental conditions, and the total deposition flux of OCS throughout the year 2011. We compare these fluxes to corresponding measurements of CO₂ flux and to simulations using the Simple Biosphere model (SiB3).

### Results and Discussion

Details of the measurement method and deployment at the Environmental Measurement Site (EMS) flux tower at Harvard Forest are presented in Methods and Supporting Information.

#### Seasonal Fluxes of OCS Show Strong Vegetative Uptake

Ecosystem fluxes of OCS (F_{OCS}) varied with air and surface soil temperature throughout the year and showed complex behavior (Fig. 1). The observed time series of OCS mixing ratios in 2011 followed the typical seasonal cycle measured previously at Harvard Forest (Fig. S1) (1). Total net OCS flux for 2011 was $-1.36 \pm 0.01 \text{ mol OCS per ha per y}$ ($-43.5 \pm 0.5 \text{ g S per ha per y}$, uptake from the atmosphere). The nighttime fluxes averaged $-0.38 \pm 0.01 \text{ mol OCS per ha per y}$ ($-12.3 \pm 0.4 \text{ g S per ha per y}$), $-28\%$ of total uptake, peaking in spring and autumn (Fig. 1B and Supporting Information).

As expected, the largest uptake fluxes were observed during the growing season (Fig. 1), starting in April when conifer trees became active and the snowpack melted to expose the forest soil. Daytime uptake of OCS (Fig. L4) increased through May and June in parallel with photosynthesis, marked by bud break of deciduous trees (May 5) and sharply increased rates of sap flow (May 19). This trend was unexpectedly interrupted by strong emission of OCS during midday hours in late July, when soil moisture was lowest and air temperature was the warmest of the year. As soil moisture gradually increased in August, net OCS uptake resumed in the daytime, but net OCS emission was observed at night (Fig. 1B). In September and October, the daily total and daytime OCS uptake flux diminished as air and soil temperature decreased, whereas nighttime OCS uptake resumed. Daytime emissions of OCS were observed yet again in early November, during the senescence of red oak (*Quercus rubra*) leaves, cancelling the nighttime uptake and resulting in a daily mean $F_{OCS} \approx 0$. In mid-December 2011, anomalously low snowfall and above-freezing air and soil temperature appeared to stimulate daytime OCS uptake, possibly reflecting uptake by conifer trees.

#### Nighttime OCS Uptake

Nighttime, light-independent uptake of OCS is likely mediated by both soils and vegetation. Nighttime transpiration through incompletely closed stomata has been observed in many tree species (26, 27), and nighttime OCS uptake has been observed in deciduous and conifer forests during the growing season (28). Soil fluxes are significant for both CO₂ and OCS but typically have opposite signs: CO₂ is respired from soils, whereas OCS is generally taken up. Carbonic anhydrase is present in the soil microorganisms typical of oxic soils found at Harvard Forest (29). OCS has been observed to be taken up by oxic soils, but the rate is notably slower on the ecosystem scale than OCS uptake by vegetation (30). Maseyk et al. (8) attributed $\sim29\%$ of total OCS flux by winter wheat to nighttime OCS uptake by foliage, with only 1–6% due to soils, at the peak of the growing season (8). The results of these studies generally agree with our results during the growing season. However, the continued strong uptake of OCS from October through December (Table S1), after the decline in activity of the deciduous canopy, implicates soil uptake as a significant portion of annual uptake, when accounting for the dormant season. We infer that uptake by soils, and potentially by conifer leaves, may contribute to the strong vertical gradient in OCS mixing ratios observed over North America from October to December (1).

### Separating Vegetative and Soil Uptake of OCS and CO₂

To separate the influence of soil and vegetative processes, we examined time periods when each process dominates: early December (soil uptake dominant), April/November (soil and conifer) and May–October (soil, conifer, and deciduous trees).

In early December, before the air temperature warmed again in mid-December, deciduous leaves were absent, and air temperature was below freezing. The soil temperature at Harvard Forest in 2011 was 2.5 °C higher than the 12-y average (2001–2012) all the way through October and November, encouraging microbial activity late in the year, even when air temperature dropped below freezing. Measurements of sap flow rate (Supporting Information) show that the red oak trees activity was sharply diminished after November 13. The early December OCS uptake ($7.2 \pm 3.4 \text{ (95\% confidence interval; CI) pmol m}^{-2} \text{s}^{-1}$; Fig. S2) was similar to the total ecosystem OCS uptake in late November, with no statistical difference between daytime ($6.0 \pm 10.9 \text{ pmol m}^{-2} \text{s}^{-1}$) and nighttime ($10.3 \pm 7.6 \text{ pmol m}^{-2} \text{s}^{-1}$) OCS uptake. We infer that the OCS uptake from cold but unfrozen soils is about 7 pmol m⁻² s⁻¹. After the soils froze, the OCS flux was not measurably different from zero throughout the winter (Fig. S2).

In late April, once the soils thawed with warming air temperature and conifer activity began, daytime uptake ($18.8 \pm 18.0 \text{ pmol m}^{-2} \text{s}^{-1}$) was greater than nighttime OCS uptake ($7.7 \pm 5.4 \text{ pmol m}^{-2} \text{s}^{-1}$), suggesting daytime conifer leaf uptake of $11 \pm 18 \text{ pmol m}^{-2} \text{s}^{-1}$. The late April nighttime uptake, after soils thaw, is comparable to early December total uptake, which we have attributed to soil uptake. However, we were not able to partition the nighttime uptake of OCS into soil and vegetative contributions in late April. This estimate, for active soils in April and December is in close agreement with the average soil uptake measured in a creek area in Colorado (28) of $7 \pm 2.6 \text{ pmol m}^{-2} \text{s}^{-1}$ and is slightly greater
Dependence of OCS Flux on Wind Direction and Temperature. Fluxes of OCS reveal heterogeneity by wind direction, reflecting the tree species distribution within the flux tower footprint (Supporting Information). In June, August, and September, daytime fluxes of OCS for air from the northwest (NW; mixed conifer and deciduous, 40.9 ± 8.2 pmol·m⁻²·s⁻¹) were almost twice as large as the OCS uptake in air from the southwest (SW; deciduous dominated, 23.5 ± 8.2 pmol·m⁻²·s⁻¹). By contrast, the net daytime CO₂ flux was roughly the same in both wind directions \[ F_{\text{CO₂}}(\text{NW}) = 23.0 ± 1.0 \text{pmol}·\text{m}^{-2}·\text{s}^{-1}, \quad F_{\text{CO₂}}(\text{SW}) = 22.1 ± 0.9 \text{pmol}·\text{m}^{-2}·\text{s}^{-1} \]. The higher daytime OCS uptake in air from the NW sector, combined with the deciduous-dominated SW sector, suggests that the magnitudes of daytime \( R_{\text{eco}} \) and GEP are both greater in this conifer-dominated sector, compared with the deciduous-dominated SW sector. In this example, \( F_{\text{OCS}} \), combined with \( F_{\text{CO₂}} \), provide unique information about the metabolic activity of plant leaves at the ecosystem scale, which are traceable to the controlling factors of photosynthesis. However, careful interpretation is required because \( F_{\text{OCS}} \) is not a direct measure of photosynthesis.

Fluxes of OCS, CO₂, and GEP showed strong dependence on air temperature (Fig. 2A). When the air temperature rose above 16 °C, net \( F_{\text{CO₂}} \) changed from positive (respiration dominated) to negative (photosynthesis dominated). When the canopy was fully developed and leaves in the canopy were most active, uptakes of both OCS and CO₂ were strongest, peaking at the highest temperature, except for the anomalous period in July when OCS was emitted by leaves but CO₂ uptake continued.

Ecosystem OCS Flux Dependence on Light and Stomatal Conductance. Both OCS and CO₂ diffuse from the atmosphere through stomata into leaves, where they are hydrolyzed by the light-independent enzyme carbonic anhydrase (CA). For OCS, the products are H₂S and CO₂, and the process is thought to be irreversible. In contrast, fixation of CO₂ through photosynthesis is a two-step process: diffusion into the leaves, reversible hydration by CA, then light-dependent and irreversible fixation by RuBisCo. Uptake of OCS does not require light, but OCS fluxes covary with light indirectly, via stomatal opening. The OCS flux is largely controlled by the conductance of the stomata in series with the mesophyll (cell walls and membranes), which regulate the rate of diffusion of OCS from the air to the site of the CA reaction. Gas exchange studies with leaves indicate that the mesophyll component of the effective conductance scales with the amount of RuBisCo in leaves of C₃ and C₄ species. The stomatal component is linked to the instantaneous rate of photosynthesis, humidity, and the chloroplast CO₂ concentration (4).

We infer that measurements of ecosystem OCS fluxes promise to provide new means to determine stomatal conductance on the ecosystem scale. The fluxes of OCS, CO₂, and GEP show strong dependences on photosynthetically active radiation (PAR; Fig. 2B), with important differences among them. We observed strong OCS uptake earlier in the day and persisting later in the day than net CO₂ uptake, which is offset by respiration. This behavior was predicted by Goldan et al. (9) and is observed here for the first time to our knowledge at the ecosystem scale (Fig. 2B).

We explored the link between \( F_{\text{OCS}} \) and stomatal conductance using the SiB3 model (Methods). The SiB3 model uses the Ball–Berry equation for stomatal conductance and has a simple parameterization of the mesophyll component. The mean diel cycle of the observed OCS flux and the calculated stomatal conductance are very similar, with enhanced activity at low-light conditions in August and September (Fig. 3C). The simulated OCS uptake shows good agreement at times of high light (Fig. 3D) but is underestimated compared with the observed fluxes, especially at times of low light and at night (Fig. 3B). This highlights the need for model refinements, such as using OCS fluxes to constrain stomatal conductance at night.

Previous laboratory studies had proposed that OCS fluxes should scale directly with stomatal conductance (32, 33); however, this is the first evidence to our knowledge demonstrating this relationship in a forest ecosystem. We find, from the observed nocturnal uptake of OCS by the canopy, strong evidence for, and potential quantification of, incomplete stomatal closure at night. The results support the view that we can use the OCS flux as a means to measure the stomatal conductance independently of the water vapor flux, providing a major advance in our capability to assess ecosystem response to environmental forcing, in a simple model framework.

Normalized Flux of OCS and CO₂. We define the flux per mole in air as \( f_x = F_X[X] \) (units: mol·m⁻²·s⁻¹), where \( F_X \) represents the observed flux of OCS, CO₂, or GEP and \( [X] \) is the ambient mole fraction of OCS or CO₂ in dry air. When adjusted for ambient temperature and pressure, this is comparable to the flux per unit molecule (units: m·s⁻¹). For periods when the canopy takes up OCS, the flux per mole represents the apparent canopy conductance of OCS that includes the stomatal, boundary layer, and

\[ \text{Normalized Flux} = \frac{F_{\text{X}}}{\text{flux}} \]

\( f_{\text{X}} / f_{\text{GEP}} \) as a function of air temperature and PAR. Values of \( f_{\text{OCS}} \) and \( f_{\text{CO₂}} \) include nighttime values. (C) The \( f_{\text{OCS}} / f_{\text{GEP}} \) ratio with air temperature, for PAR > 300 \( \mu \text{E}·\text{m}^{-2}·\text{s}^{-1} \). (D) The \( f_{\text{OCS}} / f_{\text{GEP}} \) ratio with PAR, for PAR > 300 \( \mu \text{E}·\text{m}^{-2}·\text{s}^{-1} \). Black dashed lines show 95% CI. July data are excluded. Only data with \( u^+ > 0.17 \text{ m·s}^{-1} \) are used.
compared with the vegetative uptake of OCS, or 1.3 measured using leaf chambers. Leaf chamber studies reported LRU early and late in the day (Fig. 2), not constant with PAR, with the highest values at times of low light, temperature to 24 °C before decreasing at higher temperatures (Fig. 2). If we assume that changes in soil flux are small across the day (i.e., times of full canopy) of 1.4 ± 0.3, within the range of the previous values. The variations in apparent flux ratio are substantially more complex than commonly assumed, due to the strong light dependence. Nevertheless, they can be well represented in simulations of SIB3 modified to include soil and canopy exchange of OCS (4) (Fig. 3).

Application of OCS Fluxes to Estimation of GEP. Ecosystem-scale fluxes of OCS have been proposed as a means to directly determine the photosynthetic uptake of carbon in the biosphere, independently of soil and plant respiration (1, 13, 14, 19). However, for this approach to work as proposed, a number of requirements must be met, many of which are not realized year-round at Harvard Forest. These conditions include the following: (i) \( F_{\text{OCS}} \) should be unidirectional (i.e., no OCS emission). We observed net OCS emission at times of ecosystem stress. (ii) Nighttime uptake of OCS should be negligible or relatively constant and quantifiable. We found nighttime uptake varies throughout the year and accounts for ~28% of the annual OCS uptake. (iii) The LRU of OCS/CO\(_2\) for the ecosystem type should be known. Our study shows that the ecosystem \( f_{\text{OCS}/\text{GEP}} \), as related to LRU, is not constant but may be predicted, with observed values falling within the reported range of LRU values, provided that environmental conditions are restricted to air temperature between 14 °C and 28 °C (Fig. 2B), PAR > 600 μmol·m\(^{-2}\)·s\(^{-1}\) (Fig. 2D), times of full canopy and average soil moisture.

In view of these limitations, we tested the applicability of OCS for the approximation of GEP (GEP\(_{\text{OCS}}\)) during ideal conditions (high illumination with moderate temperature and soil moisture) in September 2011 (LRU = \( f_{\text{OCS}/\text{GEP}} = 1.5 \pm 0.5 \)). Using the mean LRU of 1.5 calculated for September, the total daily sum of GEP\(_{\text{OCS}}\) and the GEP\(_{\text{CO}_2}\) agrees to within 3.5%, a good agreement in August and yet again during senescence in November. Heat stress may have been a determining factor in the observed OCS emission in summer, which was strongly enhanced at air temperature above 21 °C. During July 19–31, OCS emission increased with rising vapor pressure deficit (VPD) and bulk sap flow rate for maple trees (Fig. S3). The peak \( F_{\text{OCS}} \) (+263 ± 17.3 pmol·m\(^{-2}\)·s\(^{-1}\)) coincided with a slight depression in stomatal conductance in the afternoon. In the absence of OCS emission from the ecosystem, the expected daytime net OCS flux due to hydrolysis by CA (based on June and August peak OCS ecosystem uptake) should have been around ~30 pmol·m\(^{-2}\)·s\(^{-1}\), and hence, the observed net flux of +20 pmol·m\(^{-2}\)·s\(^{-1}\) in late July could correspond to a maximum gross emission by the responsible mechanisms of 50 pmol·m\(^{-2}\)·s\(^{-1}\) at midday. A recent study reported OCS emissions from temperature-stressed soils and senescent wheat at harvest time (8, 20).

Emission of OCS. Both light-dependent and light-independent mechanisms appear to contribute to the net OCS emissions from the ecosystem observed during an anomalous period in July. Net emissions were observed forest-wide (all wind directions), both day and night, under the conditions of high air temperature (30 °C) in late July and early August. Net OCS emission was also observed in the deciduous-dominated wind sector, a good agreement in August and yet again during senescence in November. Heat stress may have been a determining factor in the observed OCS emission in summer, which was strongly enhanced at air temperature above 21 °C. During July 19–31, OCS emission increased with rising vapor pressure deficit (VPD) and bulk sap flow rate for maple trees (Fig. S3). The peak \( F_{\text{OCS}} \) (+263 ± 17.3 pmol·m\(^{-2}\)·s\(^{-1}\)) coincided with a slight depression in stomatal conductance in the afternoon. In the absence of OCS emission from the ecosystem, the expected daytime net OCS flux due to hydrolysis by CA (based on June and August peak OCS ecosystem uptake) should have been around ~30 pmol·m\(^{-2}\)·s\(^{-1}\), and hence, the observed net flux of +20 pmol·m\(^{-2}\)·s\(^{-1}\) in late July could correspond to a maximum gross emission by the responsible mechanisms of 50 pmol·m\(^{-2}\)·s\(^{-1}\) at midday. A recent study reported OCS emissions from temperature-stressed soils and senescent wheat at harvest time (8, 20). The
metabolism of sulfur containing amino acids, which increases with temperature and plant stress, may lead to OCS production (8) in a similar manner to CO (34) and CH production (35) from thermal degradation. However, the emission observed here occurred at temperatures much lower than in the wheat field study. Nighttime OCS emission peaked in August (Fig. 1B), when CO , respiration was greatest, indicating that there is also a different, light-independent emission mechanism, possibly associated with decomposition.

In early November, OCS emissions of ~5 nmol·m -2·s -1 were observed briefly during the leaf senescence of the red oak trees. It is possible this emission occurred through a process similar to that observed during wheat senescence in Oklahoma (8). High surface soil temperature was also implicated as a source of OCS in that study. However, the high soil temperature observed in Oklahoma (45 °C) was through the canopy at Harvard Forest because the canopy shielded the forest floor from direct light, and there is no correlation of OCS emission with soil temperature in November. Therefore, we suspect that the source of OCS may have been within the senescent canopy or from freshly fallen leaves in the litter layer on the forest floor.

Because the air temperature at Harvard Forest has warmed 1.5 °C over the past 50 years (36, 37) with increasingly large interannual variability, drought and heat stress events may increase in frequency (38). Our results suggest that climate change may shift the balance between OCS uptake and emission processes at Harvard Forest and in similar terrestrial ecosystems, leading to changes in the global OCS budget.

Conclusions

Our year-long measurements at Harvard Forest demonstrate that OCS flux observations provide quantitative, independent measures of metabolic activity and biophysical properties of the forest canopy at the ecosystem scale. We observed net uptake of OCS totaling 1.36 ± 0.01 mol OCS per ha per yr (43.5 ± 0.5 g S per ha per yr), predominantly in the daytime (72%), with the balance at night attributed to soil consumption and to vegetative uptake through incompletely closed stomata. The daytime OCS flux was found to be bidirectional, with net emission during very hot, dry conditions and when vegetation senesced in autumn.

Uptake of OCS by a forest canopy is regulated by stomatal conductance, mesophyll conductance, and the activity of carbonic anhydrase, acting in series. At times of peak carbon uptake (full canopy, high illumination, and adequate soil moisture), OCS fluxes are directly proportional to photosynthetic carbon flux, with minor contributions from soils and other processes. However, because OCS uptake does not depend on light levels and on CO , activity directly as for CO , the leaf scale relative uptake ( ) has systematically higher values at dawn and dusk than at midday and, likewise, spring and fall versus summer.

We found that daytime OCS uptake was well simulated by the Simple Biosphere model (SiB3), using a basic, low-dimensional representation of OCS metabolism by plants. Thus, the observations can quantitatively constrain the aggregated functioning of the photosynthetic apparatus, at ecosystem scale, in the model framework. SiB3 underestimated uptake of OCS at times of low light and at night and did not account for production processes observed under stress conditions and during senescence. Refinement of the model is needed to account for these features, but these influences on total fluxes are relatively modest.

We conclude that OCS fluxes provide a powerful means for quantitatively measuring the large-scale photosynthetic activity of the terrestrial biosphere. By using a proper model formulation, OCS flux measurements over a forest allow us to directly observe and quantify the mechanisms that mediate temporal changes and spatial heterogeneity of canopy gas exchange of CO and H O at the ecosystem scale.

Methods

A tunable infra-red laser direct absorption spectrometer (TILDAS; Aerodyne Research Inc.) was used to measure atmospheric mixing ratios and derive gradients and fluxes of carbonyl sulfide and water vapor at 2,048,495 cm -1 and 2,048,649 cm -1, respectively. Mixing ratios of OCS and H O at a frequency of 4 Hz for eddy covariance flux (eF; August 2011 to December 2011) or 1 Hz for gradient flux (gF; January 2011 to August 2011) were calculated using TDL Winfird software (Aerodyne Research Inc.). The 15 minute averaging was typically 14ppt at 4 Hz, averaging down to <1ppt at 60 s. The sensor is a further development of earlier instruments (17, 18). More details about the measurement technique and associated instrumental tests and the theory behind the flux calculations are provided in Supporting Information and Figs. S4–S7. Tests were conducted to ensure continuity of measurement techniques. A comparison of the OCS mixing ratios (TILDAS) observed at the same time as National Oceanic and Atmospheric Administration (NOAA) flask samples is shown in Fig. S1.

Measurements were made at the Environmental Measurement Site (EMS) at Harvard Forest, Petersham, MA (42.54°N, 72.17°W, elevation 340 m). The CO flux has been measured at this Long-Term Ecological Research (LTER) site since 1990 (24). Details about the site, environmental conditions, and ancillary measurements during the study period are described in Supporting Information. Environmental conditions for the study were typical of New England. Up to 75 cm of snow accumulated between January and April in 2011. The air temperature ranged from −28 °C in January to 35 °C in July. At Harvard Forest, conifer trees are generally not active when the air temperature is consistently below 0 °C (39). The CO flux from soil respiration depends mainly on microbial activity and CO diffused through the snowpack, with increased exchange from wind pumping. Microbial activity continued through the winter as the soil temperature was partially shielded from the low air temperature by the insulating snowpack (40) before the snow cover was removed as the soil in early March. Bud break was observed for deciduous species around May 5, and senescence began late in October. Prolonged power loss resulted from damage to power lines and damage to electronic equipment due to lightning on May 28. Because no OCS fluxes were measured during the first 2 wk of May and again the first 2 wk of June, the mean uptake for both May and June was based only on measurements during the last half of each month.

There was less than 60 mm precipitation during June and July, and this precipitation was concentrated into four short events. Prolonged high temperature (>30 °C) affected the site in mid-July, resulting in low soil moisture in the area. Storms arrived in early August, bringing prolonged heavy precipitation and increasing soil moisture. Hurricane Irene on August 28 caused extensive flooding in the region. October was unseasonably warm, and leaves were still on trees when a snowstorm on October 29 brought the most snowfall for all of 2011. The last snowfall of the winter was on May 1, corresponding to a net uptake of 43.5 ± 0.5 g S per ha per yr (as OCS) per ha per yr or 16.3 ± 0.1 g C as (OCS) per ha per yr by the biosphere. The total CO uptake for the year, selected from times of valid OCS fluxes, was 0.26 ± 0.14 Mg C per yr or 2.7 Mg C per ha per yr. This size of the observed range of net CO uptake of 1.0–4.7 Mg C per ha per yr for the years 1992–2004 (25). Overall, the OCS fluxes had a greater relative uncertainty than fluxes of CO , reflecting a combination of both a less precise measurement of the OCS flux (the gradient flux calculated OCS flux has more uncertainty than the eddy covariance calculated OCS flux) and more variability of the actual daytime OCS fluxes.

The Simple Biosphere Model version 3 (SiB3), adapted to include OCS, was run for the 2011 meteorology data from Harvard Forest (41). SiB3 links stomatal conductance (both C3 and C4) to the energy budget (42, 43) and incorporates satellite-specified phenology (44). Stomatal conductance, determined by the Ball–Berry equation (45), has a direct dependence on relative humidity and CO concentration and indirect dependence on soil water, temperature, light, and humidity through the assimilation term. Both leaf and soil uptake of OCS are explicitly represented in SiB3 (4) independently from flux sampling. Carolina Sisco for analysis at National Oceanic and Atmospheric Administration (NOAA), Brad Hall for OCS standardization at NOAA, Ryan McGovern at
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Supporting Information

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Technical Details

**Instrument Description.** A TILDAS (Aerodyne Research Inc.) was used to measure atmospheric mixing ratios and derive gradients and fluxes of carbonyl sulfide and water vapor at 2,048.495 cm$^{-1}$ and 2,048.649 cm$^{-1}$, respectively. There were no CO$_2$ absorption lines in the spectral range of this laser. Mixing ratios of OCS and H$_2$O at a frequency of 4 Hz (eddy flux) or 1 Hz (gradient flux) were calculated using TDL Wintel software (Aerodyne Research Inc.). A background spectrum (30-s duration) was obtained every 10 min and interpolated and subtracted from the sample spectra to account for any temporal changes in instrument response. A diaphragm pump was used for gradient flux measurements, which resulted in a flow rate of ~3 standard liters per minute (slm) and cell response time of 15 s (90% response time). The first 60 s at each level were discarded to allow for equilibration of water vapor. The 1σ instrumental precision was 5 pptv (pmol·mol$^{-1}$) in 1-s averaging down to 0.9 pptv at 100 s. During eddy flux measurements, a TriScroll 600-slm pump resulted in a flow rate of 12 slm through the cell and a response time of 1 s. The 1σ instrument precision was typically 14 pptv at 4 Hz, likewise averaging down to <1 pptv at 60 s. The sensor is a further development of previous work (17, 18).

The combined water vapor dilution and pressure broadening correction factor was 1.27 at this wavelength, which, if not corrected, could have caused an underestimation of 7 pptv (in 400 standard liters per minute) at 60 s. During eddy flux measurements, the vertical concentration gradients of OCS and H$_2$O, respectively, measured simultaneously by the TILDAS at two heights (29.5 m and 24.1 m),

$$gX = \left[X\right]_{29.5\text{m}} - \left[X\right]_{24.1\text{m}}/(29.5 - 24.1),$$  

and the water vapor flux $F_{\text{H}_2\text{O}}$ is measured directly by eddy covariance at the EMS tower using a infrared gas analyzer [IRGA; Li-COR 6262 (24)]. The nominal TILDAS water vapor mixing ratios were 22% higher than the calibrated water vapor mixing ratios measured by the IRGA. The water vapor observed by the TILDAS was based on spectroscopic parameters and was not externally calibrated, so this correction was applied to the TILDAS water vapor mixing ratios before calculation of the gradient flux.

The OCS flux could not be calculated for 23% of the OCS measurements made during the May–August 2011 sampling period. This was due to a combination of rain events (when no water vapor flux was calculated) and unrealistic water vapor mixing ratios ($\Delta$H$_2$O outside the 95% quantiles of the total data), which resulted in equally unrealistic OCS fluxes. Fig. S4 shows the diel cycle of the measurements of OCS gradient (Fig. S4A) and H$_2$O gradient (Fig. S4B), the H$_2$O flux measured by eddy flux (Fig. S4C), and the calculated OCS flux using the gradient flux method (Fig. S4D) for June 14, 2011. The CO$_2$ flux measured by eddy covariance (Fig. S4E) is included for comparison. Negative fluxes indicate loss from the atmosphere and uptake by the biosphere.

The overall uncertainty of the gradient flux method was calculated for each point as the root-mean-square of the 95% confidence intervals of the gradient measurements (gOCS and gH$_2$O) and the mean error of the eddy covariance calculated water vapor [15% (24)]. As the instrument is optimized to OCS detection, the error in the water vapor gradient measurement, combined with the SD of the water vapor mixing ratio within a 30-min period, dominated the overall uncertainty. For the June–July period, the uncertainty in absolute fluxes ranged from 0.05 pmol·m$^{-2}$·s$^{-1}$ to 20 pmol·m$^{-2}$·s$^{-1}$ on rare occasions with a median of 0.43 pmol·m$^{-2}$·s$^{-1}$. For example, as shown in Fig. S4, this uncertainty reaches a maximum of 5.7 pmol·m$^{-2}$·s$^{-1}$ for an OCS flux of 1.1 pmol·m$^{-2}$·s$^{-1}$ on June 14, 2011.

For the gradient flux method, ambient air was alternatively sampled from the tower heights of 29.5 m and 24.1 m using 40 m diameter 4-wk to prevent artificial production of OCS on chemically aged or dirty surfaces (Artificial OCS Production). These filters resulted in a pressure drop through the tubing, which reduced the effects of adsorption/desorption on the tubing. The black Synflex tubing also reduced any sunlight effects on the sample. The air in each sampling line was tested after each background (10- or 30-min interval) to ensure no in situ production of OCS (short-lived increase in OCS). The materials in the instrument were carefully chosen to minimize any artifacts during sampling: clean Teflon filters, Synflex tubing, stainless steel solenoid valves, and the glass sampling cell were not found to scavenge or emit OCS. No pump was used upstream of sampling to prevent contamination of the sample gas. Some initial measurements were made at 25 m and 1 m during the winter of 2010–2011. The calculated fluxes for this winter 2011 period agreed with eddy fluxes for winter 2012, so these early data have been included in the seasonal cycle of $F_{\text{OCS}}$. For eddy covariance flux measurements, only the 29.5-m inlet was used.
The gradient flux method has been used successfully at Harvard Forest previously to measure fluxes of hydrogen (47), non-methane hydrocarbons (48, 49), and isoprene (50). In each of these studies, the use of CO$_2$, H$_2$O, and air temperature produced similar fluxes throughout the year with varying precision and accuracy. These methods were further validated by McKinney et al. (51), who found very similar fluxes of isoprene using a disjunct eddy covariance method, compared with Goldstein’s gradient flux method. Particularly relevant to the study here, Meredith et al. (47) found that the gradient flux method using either H$_2$O or CO$_2$ was valid throughout 2011.

To test scalar similarity using water vapor during the anomalous hot period in July, we calculated the OCS flux using the CO$_2$ gradient and flux from the eddy covariance system (Fig. S5). The CO$_2$ gradients are smaller than H$_2$O, and the top levels are measured less frequently, which introduces additional noise to the calculated flux compared with the flux calculated from H$_2$O. The OCS flux calculated based on CO$_2$ shows the same seasonal cycle for both day and night data as OCS from water vapor but with substantially more noise in the CO$_2$-based flux, as expected. Fig. S5 shows $F_{OCS}$ for the July 21–31 OCS emission period calculated from H$_2$O (FOCS$_{H2O}$) and CO$_2$ (FOCS$_{CO2}$). Regardless of the method used, statistically significant emission of OCS was observed throughout the day in the July period.

Eddy covariance method. The eddy covariance fluxes of OCS ($eF_{OCS}$) and H$_2$O ($eF_{H2O}$) were calculated from high-frequency (4 Hz) measurements of OCS and H$_2$O made by the TILDAS at 29.5 m. After subtracting a block average for the interval, the covariance of the residual of the vertical wind velocity ($w'$) and concentration (OCS’ or H$_2$O’) for each 30-min interval was calculated as in Gaulden et al. (21), e.g.,

$$F_{OCS} = \bar{w}OCS' eF_{OCS} = w'OCS eF_{H2O} = w'H2O'. \ [S3]$$

The instrument synchronization time lag was determined by maximizing the correlation between $w'$ and H$_2$O’. This lag also accounted for differences in computer clock times between the sonic and OCS data systems, which increased gradually after each synchronization reset (daily). The flux is rotated to the plane where the mean vertical wind is zero (52). The calibrated IRGA water vapor fluxes were used for all analysis. Accurate fluxes can be calculated even though high-frequency noise limits the precision of the OCS concentration at short times because the noise is not correlated with vertical wind velocity. The error in the eddy covariance was determined by calculating the root-mean-square combination of observed covariance for periods ±25 s from the lag time. This resulted in a mean SE in the eddy covariance calculated OCS flux of 14%.

Gradient flux and eddy covariance comparison. Both gradient measurements and eddy flux measurements were made for a limited time period, 6–12 August 2011, when additional measurements were made at a height of 24.1 m for 120 s every 30 min. This shorter sampling period at 24.1 m resulted in a greater error in the gradient flux ($gF_{OCS}$) for this period (47). In a comparison of the two methods, the composite diel cycle (2 hourly bins) of $gF_{OCS}$ (Fig. S6, black circles) and $eF_{OCS}$ (Fig. S6, red boxes) for periods of common measurements showed similar behavior but with slightly more variance in $gF_{OCS}$, as expected. The overall trend through the composite day compares well for both methods, with no statistical difference between the daily mean flux calculated by either method: daily mean OCS uptake of $-8.6 \pm 6.2$ (95% CI) pmol·m$^{-2}·s^{-1}$ for $gF_{OCS}$ and $-9.6 \pm 4.4$ pmol·m$^{-2}·s^{-1}$ for $eF_{OCS}$. The gradient flux of OCS underestimates the total daily flux ($gF_{OCS} = -174$ pmol·m$^{-2}·s^{-1}$) by 7% compared with the eddy flux ($eF_{OCS} = -187$ pmol·m$^{-2}·s^{-1}$). The signs and the diel patterns of the flux are consistent for both methods, except during transition periods near sunrise and sunset when fluxes, especially the water vapor flux used to calculate $gF_{OCS}$, are small and neither method is reliable.

OCS storage. The actual net uptake or emission of a trace gas by the ecosystem is the observed vertical flux plus any accumulation (or depletion) in the canopy space below the flux sensor (storage term). For CO$_2$, the storage term is significant compared with the vertical flux, especially around dawn and dusk transitions, disregarding nonideal conditions with significant horizontal advective fluxes. Although the storage term sums to nearly 0 over a daily interval, it must be included to interpret net CO$_2$ exchange on subdaily intervals. During summer 2012 (and when large CO$_2$ storage values were calculated), changes in storage of OCS calculated from OCS profile measurements were negligible. The physical processes affecting storage should not change from year to year, so the contribution to the ecosystem flux of OCS from 2012 should be applicable to 2011. Therefore, storage has been neglected in the OCS flux results that we report here.

Artificial OCS Production. Heterogeneous production of OCS on the surface of the contaminated Teflon filters was observed over 5 d after sampling an anthropogenically influenced air mass in February 2011 because unsafe climbing conditions prevented immediate replacement of the filter, which had been in place since late December. This OCS production was observed as large, short-lived pulses of OCS (up to 800 pptv) when sampling the line (and contaminated filter) after zero air background measurements. However, no evidence of OCS production from filter contamination was observed during the summer emission period described in the main text. Air mass trajectories for this February event indicate that the air was influenced by high sulfur emission from the copper and nickel smelters in Sudbury, ON, Canada, and SO$_2$ mixing ratios of greater than 60 ppbv were observed in the same air mass at a site 60 miles east of Harvard Forest (Aerodyne Research) on the same day. OCS dissolves, but is not hydrolyzed, in acidic water. Belviso et al. (53) measured supersaturated OCS in acidic rainwaters in France and suggested an in situ production of OCS from the acid catalyzed reaction of thiocyanate salts. No further studies have confirmed this suggested mechanism. However, the emission of high mixing ratios of OCS from Teflon filters could be related to a similar production mechanism because OCS production continued for a number of days and was increased in warmer, and slightly more humid, daylight conditions. There is limited literature on the heterogeneous production of OCS and potential mechanisms should be investigated in future studies. Data with contaminated filter production of OCS have been removed from further analysis and from Fig. S1.

Materials for the instrumental setup were carefully chosen to ensure no artificial production of OCS in the system. Testing showed that OCS was produced by rubber diaphragms in pumps and resulted in strong OCS production (pulses up to 24 ppb) in recirculating soil chambers at Harvard Forest. No soil chamber data were used in the analysis presented here. Neoprene and plastic tubing, which are often used in soil chambers, were particularly strong producers of OCS. Clean Synflex and Teflon tubing were not found to produce observable OCS. Although steps have been taken to minimize the effect of material contamination and to remove any data influences by atmospheric contamination, it is possible that the large OCS emission observed in July may be the result of some unknown physical production mechanism. In their studies of OCS in a wheat field, Maseyk et al. (8) observed OCS emission of 217 μg S per m$^2$ over the final 10 d of measurements (from a total of 657 μg S per m$^2$ over 7 wk). We estimate a comparable OCS emission of 207 μg S per m$^2$ over the 10 d of observed net OCS emission at Harvard Forest.

Soil warming and nitrogen fertilization experiments have been conducted in plots to the SW of the tower from 2006 to present, including during 2011 (54). These experiments use ammonium nitrate (NH$_4$NO$_3$) to fertilize 12 plots of size 3 x 3 m. The fertilizer
contains trace levels of sulfur (~0.002% sulfur as SO$_4^{−}$), which is equivalent to an application of 2.2 g S per ha per y, a less than 0.01% increase on the sulfur content of the soil. The periods of OCS emissions were not found to correlate with the application of the fertilizer. Although we cannot discount the possibility of an OCS artifact from the fertilizer, we suspect that the small area involved and the low levels of sulfur application are too small to contribute to the observed OCS signal. Nitrogen fertilization experiments also found increased OCS emission from soils (55), but we do not see a correlation with soil temperature and the related increase in microbial activity. It is possible that the sulfur present in the soils at Harvard Forest, like the soils of the wheat fields in Oklahoma (8), is a source of OCS through some unknown biophysical mechanism.

Site Description and Ancillary Measurements

**Site Description.** Measurements were made at the Environmental Measurement Site (EMS) at Harvard Forest, Petersham, MA (42.54°N, 72.17°W, elevation 340 m). The CO$_2$ flux into and out of the forest has been measured at this LTER site since 1990 (24). The 30-m meteorology tower extends about 5 m over the forest canopy and is located on moderately hilly terrain surrounded by several kilometers of relatively undisturbed forest; ~80% of the turbulent fluxes are produced within 0.7–1 km of the tower (56). The basal area (m$^2$ ha$^{−1}$) of various tree species within the footprint of the tower is tracked on plots established in 1993. In 2011, the southwest sector was dominated by deciduous species red oak (20.0% basal area) and red maple (11.8%) with some black oak (2.6%) and ash (2.1%). The northwest sector was more mixed with red oak (17.3%) and hemlock (13.2%) dominating and some red maple (9%), red pine (7.3%), and white pine (5.4%). A dried up pond, which is now an area of new tree growth, was also located in the northwest sector.

Soils at Harvard Forest are acidic and originate from sandy loam glacial till. The diversity and richness of the soil microbial community is somewhat reduced at low soil pH (57), but the soil at Harvard Forest contains representatives of the phyla typical in most soils, many of which can encode for one or more carbonic anhydrase enzymes (29).

**CO$_2$ Flux Measurements.** The CO$_2$ flux at the EMS tower was measured by eddy covariance as described extensively in previous work (21, 24) and in Figs. S5–S7. The CO$_2$ flux term accounts for storage of CO$_2$ within the canopy as determined from gradient measurements of the CO$_2$ concentration (58). The daytime respiration of CO$_2$ is projected from the observed temperature dependence of respiration at night. To estimate gross ecosystem productivity (GEP) from the measured CO$_2$ flux, we use the difference between the daytime CO$_2$ flux and the projected daytime respiration (21).

The Hemlock Tower is another flux tower at Harvard Forest located 500 m away from the EMS tower in a mature hemlock stand. The CO$_2$ uptake by confier species in 2011 was found to be greatest in April, May, and June (2.1–2.4 g C per m$^2$ per d) before being drastically reduced in July (0.5 g C per m$^2$ per d), recovering in August (1.5 g C per m$^2$ per d) and reducing in the fall (0.4–0.6 g C per m$^2$ per d; September–October). The confier uptake flux increased again in November (1.1 g C per m$^2$ per d) with higher air temperatures before essentially stopping in December (0.008 g C per m$^2$ per d).

**Sap Flow Measurements.** Ecosystem-scale flux observations cannot distinguish the canopy flux from the soil flux, because both sinks are located beneath the flux measurement point. Measurements of sap flow through trees (i.e., water uptake by trees) provide understanding of whole-tree transpiration with high temporal resolution when measured continuously throughout the growing season. Because both transpiration and photosynthesis are controlled by stomatal conductance, measurements of sap flow and eddy flux can be combined to understand patterns of canopy carbon uptake (59). We measured rates of sap flow (60) in the dominant (by mass) deciduous tree species *Quercus nigra* (northern red oak) and *Acer rubrum* (red maple) in a nearby site at Harvard Forest during a period that overlapped with OCS flux measurements (Figs. S3 and S7). These measurements provide an indication of tree activity that has been used to understand the observed OCS (and CO$_2$) fluxes. Two sensors were installed at breast height on six individual red oak red maple trees (24 sensors total).

Sap flow rates in both species began to increase on May 19, just after bud break. Senescence began around late October, with water uptake by the red oak continuing until about November 13. Elevated sap flow was generally observed before midnight throughout the growing season before reducing to minimal levels in the early hours of the morning. Fig. S7 shows the summer sap flow rates staying high into the late afternoon after both PAR and the water vapor flux began to decrease. The bulk tree activity, as observed by sap flow rates, showed that the red oaks continued to be active for up to 5 h into the night before reaching zero.

**OCS:CO$_2$ Atmospheric and Ecosystem Relative Uptake**

The effect of vegetative uptake on ambient OCS mixing ratios can be explored by looking at a ratio of OCS to CO$_2$. The atmospheric relative uptake (ARU) is the seasonal change in the OCS:CO$_2$ uptake ratio (1):

\[
ARU = \frac{[OCS]_{\text{max-min}}}{[OCS]_{\text{annual mean}}} \times \frac{[CO_2]_{\text{annual mean}}}{[CO_2]_{\text{max-min}}},
\]

where $[X]_{\text{max-min}}$ is the difference between spring maximum and autumn minimum ambient mixing ratios of OCS and CO$_2$ normalized by their annual mean. We calculate an ARU of 8.5 for 2011, which is similar to the ARU (~8 ± 2) calculated from a multiannual analysis of flask data collected at Harvard Forest for 2000–2005 (1).

The ratio of the ecosystem flux per molecule of OCS to the flux per molecule of CO$_2$ ($f_{OCS/CO_2}$) can also be compared with the ecosystem relative uptake (ERU) of OCS to CO$_2$, reported for studies on larger scales (1, 13). The ERU calculated from aircraft profile data [4.6–6.5 for the New England area in July–August 2004 (13)] was higher than the $f_{OCS/CO_2}$ ratio calculated for the Harvard Forest flux tower in this study: during summer months when photosynthesis was greatest (June–September, excluding July), the mean daily $f_{OCS/CO_2}$ ratio was 2.6 ± 0.7, and the mean daytime (PAR > 300 µE m$^{−2}$ s$^{−1}$) $f_{OCS/CO_2}$ ratio was 1.5 ± 0.3. The $f_{OCS/CO_2}$ ratio increased from August through October (Table S1). This difference between the regional scale and the tower (local scale) is likely due to the larger nonvegetative sources of CO$_2$ (including anthropogenic) than OCS (marine and anthropogenic) in the wider region not present within the tower footprint. This example illustrates the value of $F_{OCS}$ data for interpretation of large scale CO$_2$ signals.
Fig. S1. Comparison of OCS (pptv; pmol·mol⁻¹⁻¹) measured by the TILDAS [30-min average (black) with 1σ SDs shown in gray], NOAA flask pair means [red points; 1σ SDs shown as red line error bars (barely visible)], and cosampled TILDAS OCS [3-h average at the time of the flask sample (gray circles)]. The flasks were sampled weekly followed by analysis by GC-MS in Boulder as part of the NOAA flask sample network (1).

Fig. S2. The OCS (black circles; pmol·m⁻²·s⁻¹) flux, CO₂ flux (green squares; μmol·m⁻²·s⁻¹), and the air temperature (blue diamonds; °C) for given surface soil temperatures in December 2011. The data are partitioned to have equal numbers of data points for each temperature shown.
Fig. S3. Diel cycles of (A) F_{OCS} (black circles; pmol m^{-2} s^{-1}) and VPD (blue squares; Pa) and (B) bulk sap flow rate for maple trees (brown triangles; g H_{2}O m^{-2} s^{-1}) and stomatal conductance (g_{S}; green squares; cm s^{-1}) for the anomalous OCS emission period on July 19–31, 2011.

Fig. S4. Components of gradient flux calculated OCS flux for June 14, 2011. (A) g_{OCS}: OCS gradient (black; pptv m^{-1}), confidence intervals of the OCS gradient (gray bars, which are barely visible). (B) g_{H_{2}O}: H_{2}O gradient (dark blue; pptv m^{-1}), confidence intervals of H_{2}O gradient (gray bars). (C) F_{H_{2}O}: H_{2}O flux (blue; mmol m^{-2} s^{-1}), 15% error on eddy covariance measurements (gray bars). (D) g_{FOCS}: OCS gradient flux (pmol m^{-2} s^{-1}), 2-h average (black), and 30-min g_{FOCS} (gray points with SE as gray bars). (E) F_{CO_{2}}: CO_{2} flux (as NEE including storage contribution) (μmol m^{-2} s^{-1}), 30-min F_{CO_{2}} (small light green points), 15% error on eddy covariance measurements (green bars), 2-h mean (dark green points).
Fig. S5. Composite diel cycle of the gradient flux OCS derived from gradients of water vapor ($F_{OCS,H_2O}$; black points) and carbon dioxide ($F_{OCS,CO_2}$; red squares) for coincident data in 3 hourly time bins for July 19–31, 2011. The error bars indicate the 95% confidence intervals of the data within the composite 3-h period.

Fig. S6. Composite diel cycle of the gradient flux OCS ($g_{F_{OCS}}$; black points) and eddy covariance OCS flux ($e_{F_{OCS}}$; red squares) for coincident data in 2 hourly time bins for August 6–12, 2011. The error bars indicate the 95% confidence intervals of the data within the composite 2-h period.
Fig. S7. Diurnal composite of OCS (black) and CO$_2$ (green) fluxes (Eastern Standard Time) for the summer months of 2011: (A1) June, (B1) August, and (C1) September, with times of low turbulence ($u^* < 0.17$ m·s$^{-1}$) removed. The 95% confidence intervals for each species are shown as black error bars. The 95% confidence intervals for CO$_2$ are barely visible. Both columns show PAR (solid orange line; 10$^{-8}$ E·m$^{-2}$·s$^{-1}$) on two different scales. A2, B2, and C2 show the sap flow rates for oak (brown triangles; gH$_2$O·m$^{-2}$·s$^{-1}$), the vapor pressure deficit (magenta dashed line; Pa), and the water vapor flux [blue/navy circles; 5 mmol·m$^{-2}$·s$^{-1}$ (multiplied by 5 for graphing)] for June, August, and September.

Table S1. Monthly mean normalized flux

<table>
<thead>
<tr>
<th>Normalized flux</th>
<th>April</th>
<th>May</th>
<th>June</th>
<th>July</th>
<th>August</th>
<th>September</th>
<th>October</th>
<th>November</th>
<th>December</th>
<th>Year mean</th>
</tr>
</thead>
<tbody>
<tr>
<td>$f_{OCS}$ m·s$^{-1}$</td>
<td>−0.4</td>
<td>−0.7</td>
<td>−1.3</td>
<td>0.7</td>
<td>−1.6</td>
<td>−1.7</td>
<td>−0.8</td>
<td>−0.5</td>
<td>−0.4</td>
<td>−0.9</td>
</tr>
<tr>
<td>$f_{OCS}$ mmol·m$^{-2}$·s$^{-1}$</td>
<td>−9</td>
<td>−17</td>
<td>−31</td>
<td>18</td>
<td>−39</td>
<td>−43</td>
<td>−18</td>
<td>−10</td>
<td>−10</td>
<td>−22</td>
</tr>
<tr>
<td>$f_{OCS}/f_{CO2}$ (ERU)</td>
<td>−0.4</td>
<td>0.7</td>
<td>−2.0</td>
<td>0.8</td>
<td>2.4</td>
<td>3.7</td>
<td>−9.6</td>
<td>5.2</td>
<td>1.4$^\ast$</td>
<td></td>
</tr>
<tr>
<td>$f_{OCS}/f_{GEP}$ (LRU)</td>
<td>1.5</td>
<td>0.5</td>
<td>−0.4</td>
<td>1.2</td>
<td>1.2</td>
<td>0.7</td>
<td>6.6</td>
<td>9.4</td>
<td>1.0$^\ast$</td>
<td></td>
</tr>
</tbody>
</table>

Monthly mean of ecosystem flux per molecule of OCS ($f_{OCS}$, m·s$^{-1}$); ecosystem flux per mole of OCS ($f_{OCS}$, mmol·m$^{-2}$·s$^{-1}$); ratio of the flux per molecule of OCS to that of CO$_2$ ($f_{OCS}/f_{CO2}$), comparable to the ERU; and ratio of the flux per molecule of OCS to that of GEP ($f_{OCS}/f_{GEP}$), comparable to the LRU for daylight hours (PAR > 300 μE·m$^{-2}$·s$^{-1}$). Note: Period of net OCS emission in July. This has been removed from the mean year calculation.

$^\ast$Only the growing season mean (June–October 2011, excluding July) was calculated for $f_{OCS}/f_{CO2}$ and $f_{OCS}/f_{GEP}$ instead of an annual mean. Only data with $u^* > 0.17$ m·s$^{-1}$ have been used in the calculations.